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Scatarie Island and adjacent Hay Island, located 2 km east of the eastern tip of the Avalonian Mira terrane of southern Cape Breton Island, Nova Scotia, contain a succession of epiclastic and other sedimentary rocks of inferred Ediacaran to Cambrian age. The age assignment was based previously on lithological comparison with the Main-à-Dieu Group and overlying Bengal Road and MacCodrum formations of the Mira River Group. Detrital zircon grains from two sandstone samples from the Bengal Road Formation yielded typical Avalonian detrital zircon spectra with middle to late Neoproterozoic, Meso- to Paleoproterozoic (1300–2200 Ma) and Neoarchean ages. They indicate maximum depositional ages of 532.4 \pm 4.2 Ma and 525.4 \pm 2.4 Ma from essentially the same stratigraphic level, consistent with the interpretation that the rocks are Cambrian. The Bengal Road Formation also yielded scarce organic-walled microfossils including an acanthomorphic acritarch identified as Polygonium sp., also consistent with Cambrian age. The fine-grained siliciclastic succession on Hay Island, tentatively attributed to the MacCodrum Formation, yielded trace fossils, including Teichichnus isp. and Gyrolithes scintillus, that confirm Cambrian age. The Hay Island Gyrolithes scintillus expands the geographical distribution of this ichnospecies, previously known mainly from the Chapel Island Formation of Newfoundland, and represents a younger occurrence.

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Ediacaran and Cambrian rocks on Scatarie Island and nearby Hay Island, Avalonian Mira terrane, Cape Breton Island, Nova Scotia, Canada

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ABSTRACT

Scatarie Island and adjacent Hay Island, located 2 km east of the eastern tip of the Avalonian Mira terrane of southern Cape Breton Island, Nova Scotia, contain a succession of epiclastic and other sedimentary rocks of inferred Ediacaran to Cambrian age. The age assignment was based previously on lithological comparison with the Main-à-Dieu Group and overlying Bengal Road and MacCodrum formations of the Mira River Group. Detrital zircon grains from two sandstone samples from the Bengal Road Formation yielded typical Avalonian detrital zircon spectra with middle to late Neoproterozoic, Meso- to Paleoproterozoic (1300–2200 Ma) and Neoarchean ages. They indicate maximum depositional ages of 532.4 ± 4.2 Ma and 525.4 ± 2.4 Ma from essentially the same stratigraphic level, consistent with the interpretation that the rocks are Cambrian. The Bengal Road Formation also yielded scarce organic-walled microfossils including an acanthomorphic acritarch identified as *Polygonium* sp., also consistent with Cambrian age. The fine-grained siliciclastic succession on Hay Island, tentatively attributed to the MacCodrum Formation, yielded trace fossils, including *Teichichnus* isp. and *Gyrolithes scintillus*, that confirm Cambrian age. The Hay Island *Gyrolithes scintillus* expands the geographical distribution of this ichnospecies, previously known mainly from the Chapel Island Formation of Newfoundland, and represents a younger occurrence.

RÉSUMÉ

L'île Scatarie et l'île Hay voisine , situées à deux kilomètres à l'est de la pointe du terrane avalonien Mira dans le sud de l'île du Cap-Breton en Nouvelle-Écosse, abritent une succession de roches épiclastiques et d'autres roches sédimentaires remontant présumément à la période de l'Édiacarien au Cambrien. L'âge attribué était antérieurement basé sur une comparaison lithologique avec le groupe Main-à-Dieu et les formations sus-jacentes de Bengal Road et de MacCodrum du groupe de la rivière Mira. Des grains de zircon détritique provenant de deux échantillons de grès de la Formation de Bengal Road ont affiché des spectres de zircon détritique avaloniens typiques des périodes du Néoprotozoïque moyen à tardif, du Mésoprotozoïque au Paléoprotérozoïque (1300 à 2200 Ma) et du Néoarchéen. Ils signalent des âges de sédimentation maximaux de 532,4 \pm 4,2 Ma et de 525,4 \pm 2,4 Ma d'un niveau stratigraphique essentiellement identique, ce qui correspond à l'interprétation situant les roches au Cambrien. La Formation de Bengal Road a de plus révélé la présence de microfossiles palynomorphes, notamment un acritarche acanthomorphe identifié en tant que l'espèce Polygonium, ce qui correspond également à l'époque du Cambrien. La succession silococlastique à grains fins sur l'île Hay, rattachée à la Formation de MacCodrum, a présenté des ichnofossiles, par exemple, l'ichnoespèce Teichichnus et le Gyrolithes scintillus, qui confirment l'âge du Cambrien. Le Gyrolithes scintillus de l'île Hay élargit la distribution géographique de cette ichnoespèce, antérieurement reconnue comme une espèce principalement présente dans la Formation de Chapel Island de Terre-Neuve, et il représente une occurrence plus récente.

[*Traduit par la redaction*]

INTRODUCTION

REGIONAL BACKGROUND

Scatarie Island is a Nova Scotia Wilderness Protected area (https://novascotia.ca/nse/protectedareas/wa_scatarie.asp) located in the Cabot Strait 2 km from the southeastern tip of Cape Breton Island (Fig. 1). The island has a long history as a fishing settlement, and archaeological research has provided evidence for 18th century fishing activity on the island; tales abound of life there in the Nineteenth and Twentieth centuries. The area is also known for its long history of shipwrecks, most recently the Great Lakes freighter *M.V. Miner*, which ran aground when a towing cable broke in September 2011 and which was removed from the island in 2015 (https://novascotia.ca/news/release/?id=20150622003).

Although designated a protected area because of its rare or unusual flora, the bedrock geology on Scatarie Island is also remarkable because it includes a rare occurrence of rocks that formed during a time interval that spans the boundary between the Ediacaran and Cambrian periods of geological time (Barr *et al.* 1992, 1996). Although much of the interior of the island consists of dense spruce and fir forest, barrens, and bogs and hence lacks outcrop, the Ediacaran–Cambrian rocks are well exposed along parts of the shoreline. The purpose of this paper is to present the first direct paleontological and U–Pb zircon evidence that rocks formed during the Ediacaran/Cambrian transition are exposed on the eastern part of Scatarie Island and adjacent Hay Island (Fig. 2).

Southeastern Cape Breton Island, geologically comprising the Mira terrane (Barr and Raeside 1989), is part of Avalonia in the northern Appalachian orogen (Fig. 1, inset). Like other parts of Avalonia, the Mira terrane is characterized by Neoproterozoic volcanic, sedimentary, and plutonic rocks, overlain by lower Paleozoic sedimentary rocks (Fig. 1). Weeks (1954) included all the Precambrian volcanic and sedimentary rocks in a single unit called the Fourchu Group; but based on mapping, petrological studies, and U-Pb zircon dating, Barr et al. (1992, 1996) subdivided the rocks into groups of three different ages, ca. 680 Ma, ca. 620 Ma, and ca. 575 Ma (Fig. 1). The name Fourchu Group was retained only for the ca. 575 Ma volcanic and minor sedimentary rocks in the coastal part of the terrane south of Louisbourg. The new name Main-à-Dieu Group was assigned to rocks west, north, and east of Louisbourg that were considered equivalent in age to and/or younger than the Fourchu Group on the basis of field relations, petrological characteristics, and a poorly constrained U-Pb zircon age of <560 Ma from a rhyolite sample (Bevier et al. 1993). Other group names were assigned to volcanic, volcaniclastic, and sedimentary rocks farther south and west in the terrane, based mainly on U-Pb zircon ages from volcanic rocks and associated plutons (Fig. 1). Subsequently, formation names were introduced for lithological units within each of these groups by Barr and White (2017a, b).



Figure 1. Simplified geological map of the northern part of the Mira terrane, southeastern Cape Breton Island, modified from Barr *et al.* (1996), showing the distribution of Neoproterozoic, Cambrian, and Carboniferous rocks and the location of Scatarie Island. Inset map shows the location of the Mira terrane and map area in the context of northern Appalachian terranes, modified from Hibbard *et al.* (2006). A = Avalonia, G = Ganderia, L = Laurentian margin, M = Meguma, PL = peri-Laurentian arcs.



Figure 2. Geological map of the southeastern part of Scatarie Island and nearby Hay Island modified from Barr and White (2017a, b, c). Locations of dated samples SMB17-234 and 235 are shown in member 3 of the Bengal Road Formation, as well as sample SMB17-232 from the underlying Northwest Cove Formation which lacked zircon. Also shown are locations of photographed specimens in Figs. 6–8. Location for sample SMB17-232 = Grid Zone 21T (WGS84) 289998 5099504

The Cambrian sedimentary rocks that locally overlie the Neoproterozoic volcanic, sedimentary, and plutonic rocks form a broadly synclinal structure in the Mira River area, and Barr and White (2017a, b) referred to them collectively as the Mira River Group. Hutchinson (1952) and Weeks (1954) divided Cambrian sedimentary rocks into six formations: Morrison River, MacCodrum, Canoe Brook, Trout Brook, MacLean, and McNeil. Barr *et al.* (1992, 1996) retained these formation names except for Morrison River, which they replaced by the Bengal Road and Sgadan Lake formations because they considered the latter to be mappable units distinct from the red beds in the underlying Mainà-Dieu Group, all of which had been included previously in the Morrison River Formation (Weeks 1954).

As defined by Barr *et al.* (1992, 1996), the Bengal Road Formation is a dominantly red clastic sedimentary unit that overlies volcanic and sedimentary rocks of the Main-à-Dieu Group. Contacts are mainly faults or are unexposed, but an originally disconformable relationship is most likely. The Bengal Road Formation differs from sedimentary units of the underlying Main-à-Dieu Group in that the former contains abundant detrital muscovite and lacks volcanic and volcaniclastic rocks. It also contains a distinctive quartzitequartz-pebble conglomerate unit at or near its base, and is interlayered with, and overlain by, red sandstone and siltstone, commonly with well-developed cross-bedding and graded bedding. Landing (1991) correlated this red bed unit of the Mira terrane with the late Ediacaran Rencontre Formation of Newfoundland; but Barr *et al.* (1992, 2003, 2012) and Reynolds *et al.* (2009) considered it to be Cambrian based mainly on the presence of detrital muscovite and on other lithological contrasts with underlying non-micaceous red beds. Maximum depositional ages of ca. 544 and 537 Ma were reported by Willner *et al.* (2013) based on U–Pb dating of detrital zircon from the Mira River area.

The overlying Sgadan Lake Formation is a distinctive white, rarely maroon, cross-bedded quartz arenite which is locally conglomeratic with abundant quartz pebbles. The Sgadan Lake Formation in its type area near Sgadan Lake was correlated with the Random Formation of Newfoundland by Landing (1991). Landing (1991) also recognized a separate unit of shale between the red bed and quartz arenite units that he correlated with the Chapel Island Formation, which in Newfoundland occurs between the Rencontre and Random formations. Barr et al. (1996) also noted the presence of shale and siltstone but did not consider those rocks to constitute a separate mappable unit and so included them in their Bengal Road Formation. Barr et al. (2012) reported a single concordant detrital zircon age of ca. 529 Ma from the Sgadan Lake Formation in its type area, although a single analysis cannot be considered to provide a reliable maximum depositional age.

Barr *et al.* (1992, 1996) also assigned white quartz arenite and conglomerate in MacCodrum Brook to their Sgadan Lake Formation, although Landing (1991) had placed those rocks in the Rencontre Formation. The Sgadan Lake Formation in MacCodrum Brook is overlain by the type section of the MacCodrum Formation of Hutchinson (1952). It consists of grey and green siltstone and shale, with minor sandstone intervals. Carbonate nodules are present close to the base of the formation and also higher in the section. This section is the only good exposure of this unit, because many outcrops farther south in the Mira River area originally attributed to the MacCodrum Formation by Hutchinson (1952) were later shown to belong to younger units (Landing 1991; Barr et al. 1996). By comparison with Newfoundland, Landing (1991) equated the type section of the MacCodrum Formation with the Bonavista Group (Cambrian Stage 2-3) and suggested that an unconformity is present between it and the underlying quartz arenite. However, Barr et al. (1996) observed interlayered quartz arenite and shale in some locations and hence, like Hutchinson (1952), considered that the units are conformable. Later Landing (2004) revised his earlier correlation of the type section of the MacCodrum Formation with Newfoundland units because of the higher sandstone content than found in the Bonavista Group, and instead attributed it to the Chapel Island Formation. The details of rocks in the MacCodrum Brook section are further complicated by the fact that Willner et al. (2013) reported a maximum depositional age of 517 ± 3 Ma for the Sgadan Lake quartz arenite there, based on results from an eightgrain detrital zircon population. Results of work in progress to evaluate the implications of this young age and the details in the still-enigmatic MacCodrum Brook section will be reported in a subsequent paper.

The MacCodrum Formation is the oldest unit on Cape Breton Island that has yielded trace fossils (Landing 2004) and acritarchs (Palacios et al. 2015). Acritarchs from the type section on MacCodrum Brook include Granomarginata and Asteridium (Palacios et al. 2015). Acritarchs and trace fossils from the type section are consistent with attribution to the Fortunian or Cambrian Stage 2. The overlying Canoe Brook Formation consists of red-brown, carbonate-rich mudstone and siltstone, maroon siltstone containing grey-green reduction spots, and minor pink to red limestone. According to Landing (1991), the Canoe Brook Formation, as well as the upper part of the underlying MacCodrum Formation (as originally mapped by Hutchinson 1952), are equivalent to the Bonavista Group and Brigus Formation in Newfoundland based on lithological characteristics and skeletal fossils. The contact between the MacCodrum and Canoe Brook formations is nowhere exposed. Landing (1991) recovered skeletal fossils from the Canoe Brook Formation along the Louisbourg Highway, which he attributed to the Camenella baltica Zone of Cambrian Stage 3 (Geyer 2019). From the upper part of the formation, Landing (1991) reported the trace fossil Teichichnus and trilobite hash. Trilobites were also discovered by Hutchinson (1952), who reported the Cambrian Stage 3 Callavia Zone-trilobite Strenuella strenua from the Victoria Bridge area (Fig. 1) in rocks originally attributed to the MacCodrum Formation but now part of the Canoe Brook Formation (Barr et al. 1996).

The Trout Brook Formation (Hutchinson 1952), dom-

inantly dark-grey to rust-brown, well cleaved shale and siltstone, overlies the Canoe Brook Formation. Toward the stratigraphic top of the Trout Brook Formation, the shale becomes locally maroon, with thin, graded beds of finegrained sandstone. The Trout Brook Formation contains Miaolingian trilobites (Hutchinson 1952) and acritarchs (Palacios et al. 2009). The MacLean Brook Formation (Hutchinson 1952) overlies the Trout Brook Formation, and consists of interbedded grey quartz sandstone, siltstone, and shale with minor light-grey quartz sandstone and maroon shale. It contains Miaolingian trilobites (Hutchinson 1952) and Miaolingian and Furongian acritarchs (Palacios et al. 2009). The MacLean Brook Formation appears to be conformable with the underlying shales of the Trout Brook Formation and is overlain by grey shale, siltstone, and limestone of the Furongian McNeil Formation (Hutchinson 1952).

GEOLOGY OF SCATARIE ISLAND AND HAY ISLAND

The geology of Scatarie Island was first described by Fletcher (1879), who provided a vivid description of the island and recognized various types of "felsites" from coastal sections. He also reported a conglomerate visible at low tide on eastern Scatarie Island that he thought to be of probable Carboniferous age, subsequently assigned to the Cambrian Bengal Road Formation (Barr *et al.* 1996, 2003, this paper). Weeks (1954) included all of Scatarie Island in his Fourchu Group, based on mapping by Hayes *et al.* (1938). However, Barr and White (1989) and Barr *et al.* (1996) re-assigned most of the Scatarie Island rocks to the Main-à-Dieu Group and also recognized for the first time the presence of probable Cambrian rocks overlying the Main-à-Dieu Group on the eastern part of the island (Fig. 2).

Barr *et al.* (1996) divided the Main-à-Dieu Group on Scatarie Island into map units based on rock type, and those units were assigned formation names by Barr and White (2017a, b, c). The inferred oldest unit on the island, the Scatarie Island Formation, forms the northeastern part of the island, in faulted contact to the south with the Northwest Cove Formation and Cambrian rocks of the Bengal Cove Formation (Fig. 2). The Northwest Cove Formation is the uppermost formation in a conformable stratigraphic succession that youngs consistently to the east across Scatarie Island (Barr *et al.* 1996; Barr and White 2017c). The formation consists mainly of amygdaloidal basaltic flows interlayered with red volcanogenic conglomerate, epiclastic sandstone, and tuff.

The overlying mainly red to maroon sedimentary rocks were assigned to the Cambrian Bengal Road Formation based on lithology (Barr *et al.* 1992, 1996). The distribution of rock types, structural orientations, and well-preserved younging directions indicate that the Bengal Road Formation occurs in a faulted synclinal structure (Fig. 2). At the time of a visit in August 1991, a gap of about 10 m separated the uppermost amygdaloidal basalt flow of the Northwest Cove Formation from an outcrop of quartzite- and quartz-pebble conglomerate, characteristic of the base of the Bengal Road Formation elsewhere in the Mira terrane (Barr et al. 1992). The conglomerate is repeated in three outcrops in the well-exposed section, although during our most recent visit in 2017, the southernmost outcrops observed in 1991 were not exposed. The rocks are divided into four members. The quartzite- and quartz-pebble conglomerate (member 1) grades into red to maroon sandstone with minor interbedded grey laminated siltstone and conglomerate of member 2, and then into a unit of red to maroon sandstone and grey laminated siltstone which lacks conglomerate (member 3), overlain by red to grey sandstone and siltstone with minor red nodular limey siltstone (member 4). Younging direction is clear throughout this section, and the first three members are repeated north of an east-northeast-trending fault (Fig. 2). All four members contain abundant detrital muscovite, absent from the underlying Northwest Cove Formation and other formations of the Main-à-Dieu Group.

Some control from mainly submerged rocky shoals enables the synclinal structure to be traced offshore toward Hay Island, but red beds do not occur on the island. There, the rocks are grey laminated siltstone, sandstone, and shale that appear lithologically similar to the MacCodrum Formation in the Mira River area (Barr *et al.* 1992, 1996). However, based on some lithological similarities to the Trout Brook Formation, in their more recent compilation maps Barr and White (2017c) assigned the Hay Island rocks to the Trout Brook Formation. In either case, no evidence for the white quartz arenite of the Sgadan Lake Formation, or for other intervening formations, was observed, and if present, it is hidden under water.

U-PB GEOCHRONOLOGY

Sample descriptions and methods

Three samples (SMB17-232, 234, and 235) were collected for U–Pb dating of detrital zircon but only the latter two contained zircon grains. Sample SMB17-232 is red pebble conglomerate from the upper part of the Northwest Cove Formation (Fig. 2). The conglomerate overlies and underlies amygdaloidal basalt flows. It contains abundant plagioclase and epidote clasts and varied lithic fragments (including volcanic glass, basalt, dacite, and rhyolite) but minor quartz and no detrital muscovite or zircon.

In contrast, samples SMB17-234 and 235, both from the red to maroon sandstone and grey laminated siltstone of member 3, contain abundant detrital muscovite, as well as zircon. SMB17-234 is red arkosic sandstone with abundant angular quartz and less abundant plagioclase clasts. Lithic clasts include both volcanic and clastic sedimentary material. Sample SMB17-235 is greywacke, finer grained than sample 234 and with more abundant muscovite. Many of the muscovite fragments had been deformed and crenulated prior to deposition. They occur with quartz and feldspar

clasts in a muddy matrix that forms about 25% of the rock.

The samples were sent to Overburden Drilling Management (Ottawa, Ontario) for electro-pulse disaggregation and zircon separation. Zircon grains in samples SMB17-234 and 235 were then handpicked, mounted on an epoxy- covered thin section, polished to expose the centres of the zircon grains, and imaged using cold cathodoluminescence to identify internal zoning and inclusions. These images were used to select ablation points (30 µm diameter), avoiding any visible inclusions, cracks, or other imperfections. Grains were analyzed by laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS) at the Department of Earth Sciences, University of New Brunswick (Appendix A: Table A1 and A2 — two runs for each sample are reported and all data from both runs are analyzed and discussed together). U and Pb isotopic compositions were measured using a Resonetics S-155-LR 193 nm Excimer laser ablation system connected to an Agilent 7700× quadrupole in-ductively coupled plasmamass spectrometer, following the procedure outlined by McFarlane and Luo (2012) and Archibald et al. (2013). Data reduction was done in-house using Iolite software (Paton et al. 2011) to process the laser output into data files, and further reduced for U-Pb geochronology using VizualAge (Petrus and Kamber 2012).

Corrections are applied as follows: for grains with <100 counts/s ²⁰⁴Pb, data are uncorrected; for grains where the percentage error on the ²⁰⁴Pb counts per second was <20%, we used a ²⁰⁴Pb-based correction (Andersen 2002), and for grains where the percentage of radiogenic Pb (Pb* in file) is less than 98.5% we used a ²⁰⁸Pb-based correction (Petrus and Kamber 2012). Data were sorted by % concordance (²⁰⁶Pb/²³⁸U versus ²⁰⁷Pb/²³⁵U for grains <1000 Ma and ²⁰⁶Pb/²³⁸U versus ²⁰⁷Pb/²⁰⁶Pb for grains <1000 Ma), and by the % of radiogenic Pb in the grains as calculated using VizualAge (Appendix A: Tables A1 and A2). Concordia and weighted mean ages as well as probability distribution histograms were calculated using Isoplot version 4.15 (Ludwig 2012).

Probability distribution histograms are based on ²⁰⁶Pb/²³⁸U dates for grains <1000 Ma and ²⁰⁷Pb/²⁰⁶Pb dates for >1000 Ma and show all grains that are between 90 and 102% concordant. To determine the youngest age represented in each sample we used only clusters of more than 3 grains with ages that overlap within error and are 98–101% concordant. Using only near-concordant grains that overlap within error is a conservative approach, which serves to reduce the possibility of misrepresenting the maximum depositional age as too young by using single grains that may have experienced Pb loss (Dickinson and Gehrels 2010). Data for reference materials FC-1 and Plesovice are presented in Appendix A, Table A2, with standards for Run 1 and Run 2 shown separately.

Results

Sample SMB17-234 contains zircon grains with a wide range of sizes between 50–300 $\mu m,$ and grains of all sizes

were analyzed. Some of the larger grains are subhedral and rectangular and some are rounded and anhedral; the smaller grains are mostly subhedral to anhedral and rounded. All analyzed grains had very weak fluorescence in CL but some of the larger grains showed faint oscillatory zoning. Of 132 grains analyzed, 80 are between 90 and 102% concordant (Appendix A: Table A1). The major peak in the cumulative probability distribution is around 520-530 Ma, with minor peaks at 1.5 Ga, 1.8 Ga and 2.0 Ga and a few grains at 3.0 Ga (Figs. 3a, b). The gaps in ages at around 1000 Ma and around 2500 Ma are typical of detrital zircon signatures in Avalonia, as is the scatter of ages between about 1300 and 2100 Ma (e.g., Barr et al. 2012). Among the Neoproterozoic to Cambrian grains, it is possible to calculate a concordia age for 12 grains in the youngest peak at 526.3 ± 2.2 Ma, but the mean square of weighted deviation (MSWD) is very high at 7.0 and has a correspondingly low probability of concordance at 0.008 (Fig. 3c). The weighted mean age for the same 12

grains is 525.4 ± 2.4 Ma with a much lower MSWD of 1.11 (Fig. 3d). In this case the weighted mean age at ~525 Ma is likely the most robust estimate of the maximum depositional age for this sample.

Zircon grains in sample SMB17-235 range in size from 50–200 µm and grains of all sizes were analyzed. Some of the larger grains are subhedral and acicular to rectangular whereas the smaller grains are mostly subhedral to anhedral and rounded. All of the grains had weak fluorescence in CL but some of the larger grains showed faint oscillatory zoning. Of 128 grains analyzed, 86 are between 90 and 102% concordant (Appendix A: Table A1). The major peak in the cumulative probability distribution is around 530 Ma, with minor peaks at 1.5 Ga, 1.8 Ga and 2.0 Ga and a few grains around 2.7 Ga and 3.4 Ga (Figs. 4a, b). Like sample SMB17-234, this sample also displays the gaps at ca. 1000 and 2500 Ma and a spread of ages between about 1300 and 2200 Ma, typical of Avalonia. For Neoproterozoic to Cambrian grains,



Figure 3. U–Pb zircon diagrams for sample SMB17-234 (data from Table A1). (a) Probability plot for all data between 90 and 102% concordant. (b) Expanded view of probability plot in (a) for ages less than 750 Ma. (c) Concordia diagram for youngest statistically valid age population. (d) Weighted mean diagram for the same 12 grains shown in (c).

three separate peaks occur on the cumulative probability distribution (Fig. 4b), but the calculated concordia age for the youngest group of 5 grains is 534.4 ± 3.9 Ma with a high MSWD of 4.5 and a low probability of concordance at 0.034. The weighted mean age of the same 5 grains is 532.4 ± 4.2 Ma with a much lower MSWD of 0.15 (Fig. 4d). In this case the weighted mean age is likely the most robust estimate of the maximum depositional age for this sample. The grains with ages scattered between 550 Ma and 780 Ma could all have sources within Avalonia (e.g., van Staal *et al.* 2020).

TRACE FOSSILS AND ORGANIC-WALLED MICROFOSSILS

Material and methods

Coastal sections on eastern Scatarie Island and on Hay

Island were examined for their fossil contents and documented by digital photography. More than 100 m of continuous outcrop of the Bengal Road Formation is accessible at low tide. As exposed in July 2017, the section commenced with about 30 m of laminated and thin-bedded, yellow- and brown-weathering grey siltstone and minor sandstone of member 2. Small-scale truncations and bedding rupture is commonly seen (Figs. 5a–c). The proportion of sandstone increases up-section into member 3 and include beds with large cubes of pyrite (Fig. 5d). Three samples of darkgrey silt-stone were collected for organic-walled microfossils from this section. Member 3 is overlain by red sandstone/siltstone of member 4 (Fig. 5e) which contains minor red nodular limey siltstone (Fig. 5f).

Outcrops tentatively attributed to the MacCodrum Formation were examined at three coastal sections on Hay Island, with a total of five samples collected for organic-walled microfossils (Fig. 2). Samples collected for organic-walled



Figure 4. U–Pb zircon diagrams for sample SMB17-235 (data from Table A1). (a) Probability plot for all data between 90 and 102% concordant. (b) Expanded view of probability plot in (a) for ages less than 950 Ma. (c) Concordia diagram for youngest statistically valid age population. (d) Weighted mean diagram for the same 5 grains shown in (c).



Figure 5. Field images of coastal exposures of the Bengal Road Formation on eastern Scatarie Island. Scale bars represent 10 mm. (a-c) Characteristic lamination and bedding in lower part of measured section of Bengal Road Formation (members 2 and 3). Samples SC17-5 and -6 were collected from fine-grained intervals in this type of rock. Deformed bedding give rise to trace fossil-like structures, particularly well seen in (c). (d) Fine-grained sandstone with large pyrite crystals. (e) Red sandstone and siltstone (member 4). (f) Red nodular limey siltstone (member 4).

microfossils were prepared and examined at Área de Paleontologia, University of Extremadura, Badajoz, following palynological procedures outlined in Vidal (1988). See Appendix B for details on locations and samples. Palynological slides containing figured and representative material are stored with the collections in Nova Scotia Museum of Natural History, Halifax (museum numbers added in the explanation of the figures).

Results: organic-walled microfossils

All samples from Hay Island had no or only small amounts of dispersed organic material and no identifiable organic-walled microfossils (see Appendix B). Two samples from the Bengal Road Formation yielded poorly preserved organic-walled microfossils (Fig. 6). This material includes probable cyanobacterial filamentous sheaths (Fig. 6a) and organic fragments of uncertain origin. A single poorly preserved acanthomorphic acritarch (Fig. 6b) is identified as *Polygonium* sp. This acritarch does not provide biostratigraphic information beyond that of a post-Ediacaran age. Sarjeant and Stancliffe (1996) restricted *Polygonium* to the Cambrian to Devonian interval, but younger occurrences have been reported. These samples provide the first records of organic-walled microfossils from the Bengal Road Formation.

Results-trace fossils

No definitive trace fossils were observed in the Bengal Road Formation. Bedding-plane exposures are not well developed, which severely limits the possibility of observing any delicate bedding-plane parallel trace fossils. Angular or rod-shaped bodies seen in vertical and oblique section (Figs. 5b-c) have similarity to trace fossils but their interpretation



Figure 6. Organic-walled microfossils from Bengal Road Formation, Scatarie Island. Scale bar in c is equivalent to 20 µm for a-d. Sample number, the Nova Scotia Museum of Natural History collection number, and England Finder coordinates (for position of microfossils on the palynological slide; <u>https://www.graticulesoptics.com/</u> <u>products/stage-micrometers-calibration-scales-grids/coordinate-graticules/s7-england-finder</u>) are provided. (a) Possible cyanobacterial filament, SC17-6, NSM020GF14.2, J-43-1 (c) Organic-walled fragment, SC17-6, NSM020GF14.2, M-41-1. (d) Degraded organic-walled fragment, SC17-6, NSM020GF14.2, B-42-4. is complicated by frequent small-scale (syndepositional?) disturbance, fracturing and rotation of laminae and beds. It is probable that structures such as those shown in Fig. 5c represent rotated pieces of primary sedimentary structures.

Trace fossils were observed on several coastal sections on Hay Island (Fig. 2). On the southern part of the island, outcrops of mainly grey-green and dark-grey siltstone contain examples of starved ripples of fine sandstone (Fig. 7b) and teichichnid trace fossils with clear evidence for spreite (Fig. 7a). The spreiten show at least 5 lamellae, with downward-oriented convexity, which indicates a retrusive development of the spreite, although no causative burrow was clearly identified. These trace fossils are assigned to *Teichichnus* isp.

Trace fossils were observed on weathered outcrop of cleaved grey-green siltstone on the northern tip of the island, with the most notable being small, vertically oriented, spiral trace fossils (Figs. 8a-c). These spirals are developed within silty material and filled with fine sand related to thin sandstone event beds. The best exposed specimen (Fig. 8a) has two whorls (whorl height approximately 6 mm), a burrow diameter of 0.7 mm and a spiral radius of 1.6 mm. Another less well-exposed specimen with comparable dimensions shows three whorls of the spiral (Fig. 8b). Both in dimensions and spiral geometry this material is like Gyrolithes scintillus from the Chapel Island Formation, Burin Peninsula, Newfoundland, as described by Laing et al. (2018). Other trace fossils from the Hay Island outcrop consist of short plug-shaped cones (Fig. 8d). Similar trace fossils from the Chapel Island Formation in Newfoundland



Figure 7. Field images of trace fossils and sedimentary structures in the MacCodrum? Formation, southern Hay Island. (a) Siltstone and fine-grained sandstone with transverse sections through *Teichichnus* spreiten. Scale bar represents 10 mm. (b) Current ripple cross-lamination, scale bar represents 10 mm.

have been identified as *Conichnus*, but both the material described here and that from Newfoundland could alternatively be the fill of the funnel-shaped top of an otherwise not preserved, short, vertical trace fossil. Additional nondescript trace fossils are also present (Fig. 8e). Trace fossils in this outcrop exhibit "floating" and "adhering" preservation, similar to that commonly seen in the Chapel Island Formation of Newfoundland (Droser *et al.* 2002), in which sand-filled burrows are floating in a finer-grained matrix (Fig. 8e) or secondarily adhered to a later sand bed. A loose sample along the same stretch of outcrop show *Teichichnus* isp. and more strongly developed sediment mixing (Fig. 8f). The slab is sedimentologically similar and is interpreted as a less weathered sample from the same interval of the succession.

DISCUSSION

Both the detrital zircon ages and fossils described here confirm a Cambrian age for the siliciclastic successions on eastern Scatarie Island and on Hay Island, a conclusion previously based on lithological correlation with the Bengal Road and MacCodrum formations, respectively, on Cape Breton Island (Barr et al. 1996). The maximum depositional ages of 532 and 526 Ma and the presence of the acritarch Polygonium sp. from the Bengal Road Formation on eastern Scatarie Island demonstrate that the upper part of this formation is definitely younger than the late Ediacaran Rencontre Formation in Newfoundland, with which the Bengal Road Formation was compared by Landing (1991, 1996). The new data reported here, combined with detrital zircon age constraints from Cape Breton Island (Barr et al. 2012; Willner et al. 2013), indicate that the Bengal Road Formation spans much of the Fortunian and Cambrian Stage 2. The age of the base of the formation is not known, nor is the duration of the inferred break in sedimentation between the Main-à-Dieu and Mira River groups. The stratigraphic position of the Bengal Road Formation could in part be equivalent to that of the Ratcliffe Brook Formation in New Brunswick, as previously suggested by Barr et al. (2012); if so, the base of the formation is intra-Fortunian.

The geological mapping information (Barr et al. 1992, 1996; this paper) indicates that the eastern coast of Scatarie Island and Hay Island are both part of the same east-plunging synclinorium (Fig. 2). Based on lithological characteristics Barr et al. (1996) attributed the Hay Island succession to the MacCodrum Formation, in which case the sandstone facies of the Sgadan Lake Formation either is not exposed above sea level or was never developed in this region. The trace fossils reported here provide ichnostratigraphic evidence for a post-Ediacaran age of the Hay Island succession, especially the presence of Teichichnus. The earliest Teichichnus appear in late Fortunian and Stage 2 rocks, and more generally spreite-burrows have been used as evidence for a post-Ediacaran age in the absence of other evidence (e.g., Bland and Goldring 1995; Jensen et al. 2016). Teichichnus is a common element in lower Cambrian strata of Avalonia



Figure 8. Field images of trace fossils from MacCodrum? Formation, northern Hay Island. (a-e) Images from vertical surfaces of cleaved siltstone. (a-c) Vertical spiral trace fossil *Gyrolithes scintillus*. Scale bars represent 5 mm. (d) Small plug-shaped sand-filled structure, which may be a *Conichnus* or a funnel-shaped top of a vertical tube. Scale bar represents 5 mm. (e) Several short trace fossils in "floating" preservation along a bedding plane marked by gentle change in grain size. A larger burrow is a possible *Gyrolithes*. Scale bar represents 5 mm. (f) Vertical view of alternation of sandstone and mudstone in loose block. Along mid-line three discrete *Teichichnus* are seen in, from right to left, transverse (t), oblique (o), and near-longitudinal (n) section. Scale bar represents 5 mm.

and Baltica (Loughlin and Hillier 2010). In the Chapel Island Formation Teichichnus occurs sparsely in member 2, more commonly so in members 3 and 4, where it is the dominant trace fossil (Landing et al. 1989; Droser et al. 2002; Gougeon et al. 2018). It is also prominent in the Bonavista Group, in which Landing et al. (1989) named a Teichichnus Interval. In the Mira River area Teichichnus is common in red mudstone of the Canoe Brook Formation (Landing 1991). The earliest Gyrolithes straddle the Ediacaran/Cambrian boundary in Newfoundland (Gehling et al. 2001; Laing et al. 2018), and northern Norway (Jensen et al. 2018). Laing et al. (2018) documented Gyrolithes scintillus from close to the basal Cambrian GSSP through 400 m of section of member 2 of the Chapel Island Formation. If the Hay Island succession overlies that on Scatarie Island, these Gyrolithes scintillus are younger than occurrences in Newfoundland.

Finally it is noted that the maximum depositional ages from zircon in the Bengal Road Formation on Scatarie Island are close to the age for the base of the Canoe Brook Formation on Cape Breton Island as inferred from fossils. This suggests that future work is needed to evaluate possible regional differences between the sedimentary successions on Cape Breton and Scatarie islands. Furthermore, although attribution of the Hay Island succession to the MacCodrum Formation is tentatively maintained here, further studies are needed to confirm this assignment, and to evaluate the possibility that the Hay Island succession is younger than the MacCodrum Formation, and perhaps consistent with the assignment by Barr and White (2017c) to the Miaolingian Trout Brook Formation.

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REFERENCES

- Andersen, T. 2002. Correction of common lead in U–Pb analyses that do not report 204Pb. Chemical Geology, 192, pp.59–79. <u>https://doi.org/10.1016/S0009-2541(02)00195-X</u>
- Archibald, D.B., Barr, S.M., Murphy, J.B., White, C.E., MacHattie, T.G., Escarraga, E.A., Hamilton, M.A., and McFarlane, C.R.M. 2013. Field relationships, petrology, age, and tectonic setting of the Ordovician West Barneys River Plutonic Suite, southern Antigonish Highlands, Nova Scotia, Canada. Canadian Journal of Earth Sciences, 50, pp. 727–745. <u>https://doi.org/10.1139/cjes-2012-0158</u>
- Barr, S.M. and Raeside, R.P. 1989. Tectono-stratigraphic terranes in Cape Breton Island, Nova Scotia: implications for the configuration of the northern Appalachian orogen. Geology, 17, pp. 822–825. <u>https://doi. org/10.1130/0091-7613(1989)017<0822:TSTICB>2.3.</u> <u>CO;2</u>
- Barr, S.M. and White, C.E. 1989. The Main-à-Dieu sequence: an extensive late Precambrian volcanic-sedimentary package in southeastern Cape Breton Island. Nova Scotia Department of Mines and Energy, Report 89-3, pp. 149–152.
- Barr, S.M. and White, C. E. 2017a. Overview map showing locations of bedrock geology maps for Cape Breton Island, Nova Scotia. Nova Scotia Department of Natural Resources, Geoscience and Mines Branch, Open File Map ME 2017-006, scale 1:220 000.
- Barr, S. M. and White, C. E. 2017b. Bedrock geology legend for Cape Breton Island, Nova Scotia; Nova Scotia Department of Natural Resources, Geoscience and Mines Branch, Open File Illustration ME 2017-001.
- Barr, S. M. and White, C. E. 2017c. Bedrock geology map of Glace Bay Area, NTS 11J/04, Cape Breton County, Nova Scotia, Nova Scotia Department of Natural Resources, Geoscience and Mines Branch, Open File Map ME 2017-021, scale 1:50.000.
- Barr, S.M., White, C.E., and Macdonald, A.S. 1992. Revision of upper Precambrian–Cambrian stratigraphy southeastern Cape Breton Island, Nova Scotia; in Current Research, Part D; Geological Survey of Canada, Paper 92-1 D, pp. 21–26. https://doi.org/10.4095/132875
- Barr, S.M., White, C.E., and Macdonald, A.S. 1996. Stratigraphy, tectonic setting, and geologic history of Late Precambrian volcanic-sedimentary-plutonic belts in southeastern Cape Breton Island, Nova Scotia. Geological Survey of Canada Bulletin 468, 84 p. https://doi.

org/10.4095/208235

- Barr, S.M., Davis, D.W., Kamo, S., and White, C.E. 2003. Significance of U–Pb ages of detrital zircon in quartzite from peri-Gondwanan terranes, New Brunswick and Nova Scotia, Canada. Precambrian Research, 126, pp. 123–145. <u>https://doi.org/10.1016/S0301-9268(03)00192-X</u>
- Barr, S.M., Hamilton, M.A., Samson, S.D., Satkoski, A., and White, C.E. 2012. Provenance variations in northern Appalachian Avalonia based on detrital zircon age patterns in Ediacaran and Cambrian sedimentary rocks, New Brunswick and Nova Scotia, Canada. Canadian Journal of Earth Sciences, 49, pp. 533–546. <u>https://doi.org/10.1139/ e11-070</u>
- Bevier, M.L., Barr, S.M., White, C.E., and Macdonald, A.S. 1993. U–Pb geochronologic constraints on the volcanic evolution of the Mira (Avalon) terrane, southeastern Cape Breton Island, Nova Scotia. Canadian Journal of Earth Sciences, 30, pp. 1–10. <u>https://doi.org/10.1139/e93-001</u>
- Bland, B.H. and Goldring, R. 1995. *Teichichnus* Seilacher 1955 and other trace fossils (Cambrian?) from the Charnian of central England. Neues Jahrbuch für Geologie und Paläontologie, Abhandlungen, 195, pp. 5–23. <u>https:// doi.org/10.1127/njgpa/195/1995/5</u>
- Dickinson, W.R. and Gehrels, G.E. 2010. Insights into North American paleogeography and paleotectonics from U–Pb ages of detrital zircons in Mesozoic strata of the Colorado Plateau, USA. International Journal of Earth Sciences, 99, pp. 1247–1265. <u>https://doi.org/10.1007/s00531-009-0462-0</u>
- Droser, M.L., Jensen, S., Gehling, J.G., Myrow, P.M., and Narbonne, G.M. 2002. Lowermost Cambrian ichnofabrics from the Chapel Island Formation, Newfoundland: implications for Cambrian substrates. Palaios, 17, pp. 3–15. <u>https://doi.org/10.1669/0883-1351(2002)017<0003:L-CIFTC>2.0.CO;2</u>
- Fletcher, H. 1879. Report of explorations and surveys in Cape Breton, Nova Scotia. Geological Survey of Canada Report of Progress for 1877–78, Part F, 11, 32 p.
- Gehling, J.G., Jensen, S. Droser, M.L., Myrow, P.M., and Narbonne, G.M. 2001. Burrowing below the basal Cambrian GSSP, Fortune Head, Newfoundland. Geological Magazine, 138, pp. 213–218. <u>https://doi.org/10.1017/ S001675680100509X</u>
- Geyer, G. 2019. A comprehensive Cambrian correlation chart. Episodes, 42, pp. 321–332. <u>https://doi.org/10.18814/epiiugs/2019/019026</u>
- Gougeon, R.C., Mangano, M.G., Buatois, L.A., Narbonne, G.M., and Laing, B.A. 2018. Early Cambrian origin of the shelf sediment mixed layer. Nature Communications, 9(1909), 8 p. <u>https://doi.org/10.1038/s41467-018-04311-8</u>
- Hayes, A.O., Bell, W.A., and Goranson, E.A. 1938. Glace Bay sheet; Department of Mines and Resources, Map 362A, scale 1:63 360.
- Hibbard, J.P., van Staal, C.R., Rankin, D.W., and Williams,H., 2006. Lithotectonic map of the Appalachian Orogen,Canada-United States of America; Geological Survey

of Canada, Map 2096A, scale 1:1 500 000. <u>https://doi.org/10.4095/221912</u>

- Hutchinson, R.D. 1952. The stratigraphy and trilobite faunas of the Cambrian sedimentary rocks of Cape Breton Island, Nova Scotia; Geological Survey of Canada, Memoir 263, 124 p. <u>https://doi.org/10.4095/101599</u>
- Jensen, S., Harper, D.A.T., and Stouge, S. 2016. Trace fossils from the lower Cambrian Kløftelv Formation, Ella Ø, North-East Greenland. GFF, 138, pp. 369–376. <u>https:// doi.org/10.1080/11035897.2015.1076029</u>
- Jensen, S., Högström, A.E.S., Almond, J., Taylor, W.L., Meinhold, G., Høyberget, M., Ebbestad, J.O.R., Agić, H., and Palacios, T. 2018. Scratch circles from the Ediacaran and Cambrian of Arctic Norway and southern Africa, with a review of scratch circle occurrences. Bulletin of Geosciences, 93, pp. 287–304. <u>https://doi.org/10.3140/bull.geosci.1685</u>
- Laing, B.A., Buatois, L.A., Mángano, M.G., Narbonne, G.M., and Gougeon, R.C. 2018. *Gyrolithes* from the Ediacaran–Cambrian boundary section in Fortune Head, Newfoundland, Canada: exploring the onset of complex burrowing: Palaeogeography, Palaeoclimatology, Palaeoecology, 405, pp. 171–185. <u>https://doi.org/10.1016/j.palaeo.2018.01.010</u>
- Landing, E. 1991. Upper Precambrian through Lower Cambrian of Cape Breton Island: faunas, paleoenvironments, and stratigraphic revision; Journal of Paleontology, 65, pp. 570–595. https://doi.org/10.1017/S0022336000030675
- Landing, E. 1996. Avalon: insular continent in the Cambrian. *In* Avalonian and related peri-Gondwanan terranes of the circum-North Atlantic. *Edited by* D. Nance and M.D. Thompson. Geological Society America Special Paper, 304, pp. 29–63. <u>https://doi.org/10.1130/0-8137-2304-3.29</u>
 Landing, E. 2004. Precambrian–Cambrian boundary interval deposition and the marginal platform of the Avalon microcontinent. Journal of Geodynamics, 37, pp. 411–435. <u>https://doi.org/10.1016/j.jog.2004.02.014</u>
- Landing, E., Myrow, P.M., Benus, A.P., and Narbonne, G.M. 1989. The Placentian Series: appearance of the oldest skeletalized faunas in southeastern Newfoundland. Journal of Paleontology, 63, pp. 739–769. <u>https://doi.org/10.1017/</u> <u>S0022336000036465</u>
- Loughlin, N.J.D. and Hillier, R.D. 2010. Early Cambrian *Teichichnus*-dominated ichnofabrics and palaeoenvironmental analysis of the Caerfai Group, southwest Wales, UK. Palaeogeography, Palaeoclimatology, Palaeoecology, 297, 239–251. <u>https://doi.org/10.1016/j.palaeo.2010.07.030</u>
- Ludwig, K.R. 2012. Isoplot 4.15: A geochronological toolkit for Microsoft Excel. Berkeley Geochronological Center. URL <<u>http://www.bgc.org/isoplot_etc/isoplot/Isoplot4_15files.zip</u>>, March 2020.
- McFarlane, C.R.M. and Luo, Y. 2012. Modern analytical facilities: U–Pb geochronology using 193 nm Excimer LA-ICP-MS optimized for in-situ accessory mineral dating in thin sections. Geoscience Canada, 39(3), pp. 158–172.

Palacios, T., Jensen, S., Barr, S.M., and White, C.E. 2009.

Acritarchs from the MacLean Brook Formation, southeastern Cape Breton Island, Nova Scotia, Canada: new data on Middle Cambrian–Lower Furongian acritarch zonation. Palaeogeography, Palaeoclimatology, Palaeoecology, 273, pp. 123–141. <u>https://doi.org/10.1016/j.palaeo.2008.12.006</u>

- Palacios, T., Jensen, S., Barr, S.M., and White, C.E. 2015. Stratigraphic constraints on Cambrian stratigraphy. Mira and Bras d'Or terranes, Cape Breton Island, Nova Scotia, Canada. Atlantic Geology, 51, pp. 127–128. <u>https://doi. org/10.4138/atlgeol.2015.005</u>
- Paton, C., Hellstrom, J., Paul, B., Woodhead, J., and Hergt, J.M. 2011. Iolite: freeware for the visualisation and processing of mass spectrometric data. Journal of Analytical Atomic Spectrometry, 26, pp. 2508–2518. <u>https://doi. org/10.1039/c1ja10172b</u>
- Petrus, J.A. and Kamber, B.S. 2012. VizualAge: A novel approach to laser ablation ICP-MS U- Pb geochronology data reduction. Geostandards and Geoanalytical Research, 36, pp. 247–270. <u>https://doi.org/10.1111/j.1751-908X.2012.00158.x</u>
- Reynolds, P.H., Barr, S.M., and White, C.E. 2009. Provenance of detrital muscovite in Cambrian Avalonia of Maritime Canada: ⁴⁰Ar/³⁹Ar ages and chemical compositions. Canadian Journal of Earth Sciences, 46, pp. 169– 180. https://doi.org/10.1139/E09-013
- Sarjeant, W.A.S. and Stancliffe, R.P.W. 1996. The acritarch genus *Polygonium* Vavrdová emend Sarjeant and Stancliffe 1994: a reassessment of its constituent species. Annales de la Société géologique de Belgique, 117, pp. 355– 369. <u>https://doi.org/10.2307/1485867</u>
- van Staal, C.R., Barr, S.M., McCausland, P.M., Thompson, M.D., and White, C.E. 2020. Tonian-Ediacaran tectonomagmatic evolution of West Avalonia and its Ediacaran–Early Cambrian interactions with Ganderia: an example of complex terrane transfer due to arc–arc collision? *In* Pannotia to Pangaea: Neoproterozoic and Paleozoic orogenic cycles in the circum-Atlantic region. Edited by J.B. Murphy, R.A. Strachan, and C. Quesada. Geological Society, London, Special Publications, 503. <u>https://doi. org/10.1144/SP503-2020-23</u>
- Vidal, G. 1988. A palynological preparation method. Palynology, 12, pp. 215–220. <u>https://doi.org/10.1080/019161</u> 22.1988.9989345
- Weeks, F.J. 1954. Southeast Cape Breton Island, Nova Scotia. Geological Survey of Canada, Memoir 277, 112 p. <u>https://doi.org/10.4095/101500</u>
- Willner, A.P., Barr, S.M., Gerdes, A., Massonne, H.-J., and White, C.E. 2013. Origin and evolution of Avalonia: evidence from U–Pb and Lu–Hf isotopes in zircon from the Mira terrane, Canada, and the Stavelot–Venn Massif, Belgium. Journal of the Geological Society, London, 170, pp. 769–784. <u>https://doi.org/10.1144/jgs2012–152</u>

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APPENDIX A: U-PB DATA TABLES

Corrections: 1 = th	treshold ²⁰⁴	Pb no cor	rection (100 cps); 2 = th	reshold	% ²⁰⁴ Pb-bas	ed correctio	n(20%	error)	; 3 = thr	eshold % fc	r ²⁰⁸ Pb	-based	correct	ion (98.	5 %Pb*).							
													Is	otopic	ratios				ũ	lculate	d ages		╞	
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ardrino	cps	(mqq)	(mqq)	0/111	cps ¹ ²	⁰⁴ Pb cps	²⁰⁴ Pb cps	cps	Pb^{\star^2}		1000	²³⁵ U [–] –	²³⁸ U	-) T	206 Pł	-	$^{206}\mathrm{Pb}$	2σ	²³⁵ U	2σ	²³⁸ U	2σ ci	onc
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SMB17-234 - 1 ^a	1.86E + 08	469	266.7	1.76	6-	11	-122.2	120050	99.83	1 1	ALSE	3.282 0.08	7 0.25	6 0.00	l 0.26	5 0.093	4 0.0013	1496	26	1476	20	1470	21	98.2
SMB17-234 - 2 ^{a,b}	1.95E+08	210.2	121	1.74	4	14	200.0	2566	99.74	1 1	ALSE	0.687 0.02	5 0.08	6 0.002	0.20	5 0.058	8 0.0018	560	67	530	15	529	10	9.6
SMB17-234 - 3 ^a	1.86E + 08	222	206.1	1.08	9	14	233.3	3243	<u>99.69</u>	1 F	ALSE	0.668 0.02	7 0.08	1 0.00	0.48	5 0.059	9 0.0020	600	72	518	17	504	8	97.2
SMB17-234 - 4 ^a	1.85E+08	103	73.6	1.40	14	12	85.7	714	99.74	1 F	ALSE	0.789 0.03	3 0.09	7 0.002	0.27	1 0.059	8 0.0021	596	76	587	19	594	11 10	01.2
SMB17-234 - 5 ^a	2.01E+08	427.1	101	4.23	4	15	375.0	30050	99.74	1 F	ALSE	3.569 0.09	2 0.26	∠ 0.00	ł 0.45	8 0.097	8 0.0013	1583	25	1542	21	1526	21	96.4
SMB17-234 - 6 ^a	1.85E+08	122.2	60.8	2.01	1	11	1100.0	10560	99.75	1 F	ALSE	0.683 0.03	2 0.08	6 0.002	0.07	0.058	3 0.0025	541	94	525	20	531	9 1(01.2
SMB17-234 - 7	2.13E+08	207.7	67.7	3.07	16	17	106.3	2624	98.38	1 F	ALSE	2.438 0.09	8 0.19	3 0.00	l 0.46	4 0.091	7 0.0028	1461	58	1251	29	1139	22	78.0
SMB17-234 - 8 ^a	1.77E+08	111.7	70.9	1.58	ς	12	-240.0	38810	99.72	1 F	ALSE	5.738 0.16	0 0.35	0 0.00	6 0.48	0.120	1 0.0022	1958	33	1935	25	1931	29	98.6
SMB17-234 - 9	1.78E+08	459	84.2	5.45	1	16	1600.0	101000	98.13	1 1	ALSE	3.022 0.08	5 0.22	1 0.00	0.56	9 0.100	0 0.0016	1624	30	1412	21	1287	19	79.2
SMB17-234 - 10 ^a	1.68E+08	268.2	37.66	7.12	ŝ	11	220.0	5146	99.87	1 F	ALSE	0.883 0.02	7 0.10	5 0.002	0.11	3 0.061	4 0.0014	: 653	49	641	15	645	10 10	00.7
SMB17-234 - 11 ^{a,b}	1.84E + 08	109.5	119	0.92	-	13	-1300.0	9500	99.53	1 1	ALSE	0.688 0.03	2 0.08	5 0.002	0.23	8 0.059	5 0.0024	585	88	531	20	524	10	98.7
SMB17-234 - 12 ^a	1.89E+08	101.3	86.64	1.17	ή	13	-260.0	31580	99.52	1 1	ALSE	4.435 0.13	0 0.30	2 0.00	5 0.39	3 0.107	8 0.0021	1763	36	1717	24	1698	26	96.3
SMB17-234 - 13	2.19E+08	547	34.73	15.75	4	21	525.0	22375	99.57	1 1	ALSE	1.636 0.08	1 0.15	9 0.00	0.87	1 0.075	3 0.0019	1077	51	981	30	951	27	38.3
SMB17-234 - 14 ^{a,b}	1.78E+08	274	165	1.66	8	14	175.0	2785	99.74	1 1	ALSE	0.668 0.02	3 0.08	4 0.002	0.17	5 0.058	3 0.0017	541	64	520	14	520	9 1(0.00
SMB17-234 - 15 ^{a,b}	1.76E+08	113.1	54.1	2.09	б	10	346.4	3243	99.63	1 1	ALSE	0.683 0.02	8 0.08	5 0.002	0.22	5 0.059	2 0.0020	574	73	526	17	524	6	9.6
SMB17-234 - 16 ^a	1.90E+08	108.6	60	1.81	ς	13	-433.3	9810	99.81	1 1	ALSE	0.710 0.03	5 0.09	0 0.00	0.15	9 0.057	8 0.0026	522	66	545	20	555	11 10	01.8
SMB17-234 - 17 ^a	1.87E+08	164.6	90.9	1.81	ю	10	333.3	4803	99.89	1 F	ALSE	0.674 0.02	8 0.08	6 0.002	0.03	5 0.057	4 0.0019	507	73	521	17	531	9 1(01.8
SMB17-234 - 18 ^a	1.82E+08	241.5	145.1	1.66	9-	12	-200.0	19870	99.72	1 1	ALSE	0.661 0.02	3 0.08	3 0.00	0.19	8 0.058	5 0.0016	549	60	516	13	512	8	99.3
SMB17-234 - 20 ^{a,b}	1.85E+08	134.8	79.6	1.69	1	11	1100.0	11490	99.61	1 1	ALSE	0.679 0.02	5 0.08	4 0.002	0.14	2 0.059	1 0.0018	571	66	524	15	520	6	99.3
SMB17-234 - 21 ^a	1.69E+08	175.4	91.25	1.92	6	15	166.7	5514	<u>99.69</u>	1 1	ALSE	4.004 0.12	0 0.28	5 0.00	0.42	8 0.102	4 0.0019	1668	34	1634	24	1618	24	97.0
SMB17-234 - 22	1.96E+08	168.7	127.3	1.33	9	11	183.3	8210	98.74	1 F	ALSE	4.020 0.11	0 0.27	2 0.00	ł 0.54	3 0.107	6 0.0018	1759	31	1636	23	1552	22	38.2
SMB17-234 - 23 ^a	1.84E+08	152.6	72.3	2.11	4	12	300.0	3120	99.57	1 F	ALSE	0.662 0.02	8 0.08	0 0.00	0.22	5 0.060	5 0.0022	622	78	513	17	496	6	9.6
SMB17-234 - 24 ^a	1.90E+08	209.7	182.9	1.15	5	13	260.0	3690	99.80	1 1	ALSE	0.690 0.02	4 0.08	7 0.00	0.15	3 0.058	0 0.0017	530	64	532	14	537	9 1(9.00
SMB17-234 - 25 ^{a,b}	1.75E+08	134	114	1.18	8	11	137.5	1349	99.85	1 F	ALSE	0.668 0.03	1 0.08	4 0.002	0.27	0.057	6 0.0020	515	76	519	18	520	9 1(0.2
SMB17-234 - 26 ^a	1.92E+08	818	13.1	62.44	1	13	1300.0	85700	99.82	1 1	ALSE	0.852 0.02	5 0.10	3 0.002	0.39	9 0.060	9 0.0012	634	42	625	14	629	10 10	0.6
SMB17-234 - 27 ^a	1.90E+08	96.7	103.6	0.93	4	14	350.0	2445	99.68	1 F	ALSE	0.856 0.03	9 0.10	2 0.002	0.13	5 0.060	9 0.0025	636	88	625	21	629	11 10	0.6
SMB17-234 - 28	1.76E+08	1237	238.8	5.18	31	13	41.9	5548	98.89	1 1	ALSE	1.564 0.04	3 0.14	6 0.00	8 0.77	7 0.077	6 0.0010	1136	26	955	17	881	14	77.5
SMB17-234 - 29 ^a	2.05E+08	195.7	184	1.06	3	15	500.0	6080	99.75	1 F	ALSE	0.725 0.03	3 0.09	0 0.00	0.08	0.058	6 0.0025	552	93	552	19	554	10 10	0.4
SMB17-234 - 30 ^a	1.62E+08	757	590	1.28	6	15	166.7	6233	99.51	1 1	ALSE	0.639 0.02	2 0.07	7 0.00	0.54	0.060	4 0.0014	618	50	501	13	478	6	95.5
SMB17-234 - 31 ^{a,b}	1.94E+08	188	87.6	2.15	8	19	237.5	2071	99.71	1 F	ALSE	0.686 0.03	2 0.08	5 0.00	0.08	2 0.058	9 0.0025	563	92	528	19	523	10	99.1
SMB17-234 - 32 ^a	1.94E+08	322	148.5	2.17	б	19	633.3	9063	99.58	1 F	ALSE	0.678 0.03	0 0.08	1 0.002	0.12	8 0.060	6 0.0024	625	85	524	18	502	6	95.9
SMB17-234 - 33 ^a	1.91E+08	188.8	85.6	2.21	2	12	600.0	7975	99.62	1 F	ALSE	0.682 0.02	8 0.08	4 0.00	0.20	7 0.059	1 0.0021	571	77	526	17	519	6	98.7
SMB17-234 - 34 ^a	1.83E+08	620	32.5	19.08	9	11	183.3	13117	99.89	1 1	ALSE	1.144 0.03	2 0.12	7 0.002	0.62	2 0.065	5 0.0010	789	32	774	15	769	12	97.4
SMB17-234 - 35	2.18E+08	446	32.1	13.89	1	18	1800.0	60940	99.35	1 1	ALSE	1.298 0.04	9 0.13	3 0.00	9.0.69	4 0.071	1 0.0016	960	46	848	19	805	17	33.8

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Table A1. LA-ICP-MS U-Pb isotopic analyses of detrital zircon samples analyzed at the University of New Brunswick.

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Corrections: 1 = tl	ıreshold ²⁰⁴ 1	2b no cori	ection (100 cps); 2 = th	rreshold %	²⁰⁴ Pb-base	ed correcti	on(20%	error);	3 = three	shold % 1	for ²⁰⁸ P	b-based	correct ratios	tion (98.	5 %Pb*)		C	lculate	d ages		_	
Samule	90 Zr	Ŋ	Th	Th/II	$^{204}\mathrm{Pb}$	±2σ	%± ²⁽	⁾⁶ Pb/ ²⁰⁴ Pb	%	Correcti	207	Pb/ + 2	, ²⁰⁶ Ρ	b/ + 24	ц Ц Ч	3 ²⁰⁷ Pb	/ + 26	207 Pb	+	²⁰⁷ Pb/	م ۲	/qd ₉₀	_ ∎+	%
authre	cps	(mqq)	(mqq)	0 /11 1	cps ¹ ²	⁰⁴ Pb cps ²	²⁰⁴ Pb cps	cps	Pb^{\star^2}	COLLECT	23	³⁵ U - 1	238	n L L	2	206 Pf		²⁰⁶ Pb	2σ	²³⁵ U	2σ 2	²³⁸ U 2	cσ co	nc
SMB17-234 - 35a ^a	1.50E+08	216.5	50.8	4.26	-9	19	-316.7	40100	99.67	I FA	LSE 2	.235 0.0	86 0.2	00 0.00	5 0.54	180.0 6	1 0.002	5 122	4 61	1190	26	1175	27 9	6.0
SMB17-234 - 36a	2.23E+08	1601	21.7	73.78	7	36	514.3	26900	99.69	1 FA.	LSE 1	.025 0.0	38 0.1	14 0.00	3 0.37	8 0.065	1 0.001	9 77	8 61	716	19	698	15 9	7.5
SMB17-234 - 36a ^a	1.56E+08	170.2	48.1	3.54	-15	13	-86.7	51500	99.95	1 FA.	LSE 4	.812 0.1	50 0.3	22 0.00	6 0.31	5 0.108	1 0.002	4 176	8 41	1785	25	1799	29 10	1.8
SMB17-234 - 37 ^a	1.83E+08	132.9	93.8	1.42	б	13	433.3	3903	99.66	1 FA.	LSE 0	.733 0.0	33 0.0	90 0.00	0.40	7 0.055	1 0.002	1 57	1 77	555	19	554	10 9	9.8
SMB17-234 - 38	1.81E+08	124.3	72.7	1.71	ŝ	10	200.0	9120	99.99	1 FA.	LSE 6	.076 0.1	70 0.3	66 0.00	6 0.53	5 0.120	3 0.002	0 196	1 30	1985	24	2010	29 10	2.5
SMB17-234 - 39	1.57E+08	460	85	5.41	99	28	42.4	574	98.64	1 FA.	LSE 1	.111 0.0	53 0.1	11 0.00	3 0.24	ł0 0.072	9 0.003	2 101	1 89	758	25	676	18 8	9.2
SMB17-234 - 40	2.21E+08	892	28.19	31.64	15	26	173.3	6460	98.64	1 FA.	LSE 1	.069 0.0	37 0.1	07 0.00	0.58	6 0.072	3 0.001	2 99	4 48	738	19	656	13 8	8.9
SMB17-234 - 41 ^a	1.78E+08	181	155	1.17	1	12	1200.0	17300	99.68	1 FA.	LSE 0	.782 0.0	30 0.0	94 0.00	0.17	0 0.055	6 0.001	9 58	69 6	584	17	582	10 9	9.6
SMB17-234 - 42	1.40E+08	624	211	2.96	-2	16	-800.0	135000	97.46	1 FA.	LSE 3	.806 0.1	10 0.2	48 0.00	5 0.67	0 0.110	8 0.001	9 181	3 31	1593	23	1428	24 7	8.8
SMB17-234 - 43	1.55E+08	1114	1020	1.09	26	15	57.7	2854	98.99	1 FA.	LSE 0	.639 0.0	19 0.0	72 0.00	1 0.30	0 0.063	6 0.001	2 72	8 40	502	12	451	7 8	6.6
SMB17-234 - 44 ^a	1.87E+08	227.6	131	1.74	16	16	100.0	1174	99.12	1 FA.	LSE 0	.668 0.0	26 0.0	76 0.00	1 0.03	4 0.062	9 0.002	1 70	5 71	518	16	475	8	1.7
SMB17-234 - 45 ^a	1.81E+08	175.6	104.9	1.67	4	10	250.0	3625	99.60	1 FA.	LSE 0	.702 0.0	24 0.0	84 0.00	0.32	8 0.060	0 0.001	6 60	4 58	538	14	522	66	7.1
SMB17-234 - 46	1.50E+08	38.8	5.92	6.55		18	-600.0	8540	99.73	1 FA.	LSE 2	.980 0.1	90 0.2	48 0.00	9 0.35	32 0.086	6 0.005	0 1353	2 111	1390	49	1429	45 10	5.7
SMB17-234 - 47 ^{a,b}	1.80E+08	56.6	118.4	0.48	13	11	84.6	362	99.66	1 FA.	LSE 0	.701 0.0	42 0.0	86 0.00	0.15	6 0.055	1 0.003	3 57	1 121	533	25	529	11 9	9.3
SMB17-234 - 48	1.95E+08	335	38.2	8.77	13	14	107.7	3600	98.85	1 FA.	LSE 1	.436 0.0	51 0.1	38 0.00	4 0.67	4 0.075	4 0.001	7 1079	9 45	902	21	830	20 7	6.9
SMB17-234 Run 2	2 (in italics	to disting	quish fro	un Run	(1)																			
SMB17-234 - 1	2.52E+08	813	147.8	5.50	<i>I</i> -	28	-2800.0	156700	98.02	I FA	LSE 1.	669 0.0	36 0.1-	14 0.00	3 0.71	9 0.083	9 0.001	4 1290	32	966	14	870	14 6;	7.4
SMB17-234 - 2 ^a	1.44E+08	240.9	148.4	1.62	5	13	260.0	3570	99.28	I FAI	LSE 0.	699 0.02	20 0.0	32 0.00	1 0.22	5 0.061	9 0.001	7 671	59	537	12	505	8	4.0
SMB17-234 - 3	2.29E+08	462	81.3	5.68	35	19	54.3	3323	98.33	I FA	LSE 2.	479 0.04	45 0.1:	0.00	3 0.76	3 0.091	9 0.001	4 1465	29	1265	13	1159	17 79	9.1
SMB17-234 - 4 ^a	2.10E+08	145.7	82.7	1.76	6-	15	-166.7	17090	99.56	I FA	LSE 0.	720 0.02	28 0.0	88 0.00	1 0.02	5 0.059	1 0.002	4 571	88	549	16	543	7 98	8.9
SMB17-234 - 5	2.13E+08	720	157	4.59	98	22	22.4	2929	97.61	I FA	LSE 5.	117 0.00	50 0.31	00.0 00	4 0.57.	3 0.122	8 0.001	1997 C	7 14	1839	10	1692	18 84	4.7
SMB17-234 - 6 ^a	2.44E+08	1361	29.8	45.67	70	25	35.7	2820	99.27	I FA	LSE 1.	0.0 600	16 0.11	00.0	1 0.24	7 0.067	4 0.000	9 851	28	708	8	663	8	3.7
SMB17-234 - 7 ^a	2.15E+08	1912	198	9.66	213	39	18.3	1047	97.64	,	3 0.	801 0.04	41 0.0	00.0 It	1 0.33.	5 0.063	7 0.002	5 732	83	595	24	559	7 93	3.9
SMB17-234 - 8 ^a	2.08E+08	352.1	17.2	20.47	13	15	115.4	12708	99.36	I FA	LSE 6.	411 0.08	83 0.31	50 0.00	4 0.66	0 0.128	1 0.001	9 2072	2 13	2033	11	1983	20 95	5.7
SMB17-234 - 9 ^a	2.15E+08	871	43.5	20.02	~	12	171.4	16657	99.58	I FA	LSE 0.	867 0.0,	14 0.10	00.0 OC	1 0.71	6 0.062	6 0.000	9 694	1 29	634	8	612	7 96	5.7
SMB17-234 - 10	2.21E+08	266	128.7	2.07	IJ.	17	340.0	24280	91.11	I FA	LSE 10.	530 0.1;	70 0.3.	76 0.00	6 0.72	1 0.203	0 0.001	9 2850	15	2482	15	2057	26 72	2.2
SMB17-234 - 11	2.37E+08	543	48.71	11.15	35	17	48.6	2820	99.17	I FA	LSE 1.	414 0.0	30 0.1.	39 0.00	2 0.19	7 0.073	5 0.001.	5 1028	41	894	13	841	11 8	1.9
SMB17-234 - 12 ^a	2.02E+08	254.1	129.9	1.96	Э	16	533.3	8760	99.25	I FA	LSE 0.	670 0.02	23 0.0.	78 0.00	1 0.00	4 0.061	6 0.002	I 660) 73	520	14	486	7 93	3.4
SMB17-234 - 13	2.38E+08	805	190.5	4.23	22	15	68.2	9945	98.56	I FA	LSE 2.	539 0.0	35 0.21	0.00	2 0.54	4 0.091	2 0.000	8 1450	16	1283	10	1184	13 8	1.6
SMB17-234 - 14	1.86E+08	339.9	73.1	4.65	10	17	170.0	6182	98.36	I FA	LSE 1.	727 0.04	40 0.1.	51 0.00	2 0.44	9 0.081	0.001	5 1243	38	1017	15	907	14 73	3.0
SMB17-234 - 15 ^a	2.16E+08	416	299.2	1.39	11	11	100.0	4100	99.19	I FA	LSE 0.	730 0.0	15 0.0	33 0.00	1 0.10	0 0.063	3 0.001	1 718	37	556	6	515	7	2.7
SMB17-234 - 16 ^a	2.21E+08	292	156	1.87	9	13	144.4	3533	99.21	$I FA_{I}$	LSE 0.	757 0.02	20 0.0i	36 0.00	1 0.01	7 0.063	6 0.001	7 728	57	572	11	533	6 9	3.1
SMB17-234 - 17 ^a	2.00E+08	308.1	185	1.67	10	19	190.0	3222	99.60	I FA_{I}	LSE 0.	641 0.02	21 0.0	30 0.00	1 0.10	9 0.057	5 0.001	9 511	73	502	13	497	7	9.0
SMB17-234 - 18 ^a	2.48E+08	1297	42.4	30.59	31	27	87.1	5845	99.43	I FA	LSE 0.	996 0.02	25 0.10	00.0 60	2 0.51	1 0.066	4 0.001.	3 819	41	701	13	668	96	5.2
SMB17-234 - 19	2.19E+08	162.1	74.4	2.18	18	12	66.7	1017	98.76	I FA	LSE 0.	792 0.02	23 0.0,	35 0.00	1 0.30	6 0.067	0 0.001	8 836	56	591	13	527	7 89	9.2
SMB17-234 - 20	2.05E+08	2230	763	2.92	70	15	21.4	6543	98.72	I FA	LSE 1.	812 0.02	23 0.In	51 0.00	2 0.81	7 0.081	0 0.000	5 1221	11	1049	8	963	12 78	8.9

BARR ET AL. – Ediacaran and Cambrian Rocks on Scatarie Island and nearby Hay Island, Avalonian Mira terrane, Cape Breton Island, Nova Scotia, Canada

	II COHOIR			ed a not	n – 7 ()		1 0-043		0/ 07/110				Isotop	oic ratio	SO SO	0 TO/ C'O		0	Calculat	ed age	8		Γ
Sample	⁹⁰ Zr cps	U (ppm)	Th (ppm)	Th/U	²⁰⁴ Pb	± 2σ ²⁰⁴ Pb cps ²	$\% \pm ^2$	⁰⁶ Pb/ ²⁰⁴ Pb cps	$\%$ Pb^{\star^2}	Correction	1S 235	$\frac{Pb}{U} \pm 2c$	$\sum_{238U}^{206} \frac{bb/}{\pm}$	2σ E	C^{3} C^{207} I	^b / ±2 Ph	α ²⁰⁷ Ρ	b/ \pm b/ b/ b/ b/ b/ b/ b/ b/ 2σ	²⁰⁷ Pb/	2 _σ +	²⁰⁶ Pb/ ²³⁸ U	2α c	% onc
SMB17-234 - 21 ^a	1.35E+08	1856	1321	1.40	151	29	19.2	897	97.79	- 3	0.6	560 0.04	5 0.082 0.0	0 1 0.4	443 0.05	80 0.00	35 53	30 132	511	28	506	~	98.9
SMB17-234 - 22 ^a	1.94E+08	270.4	160.9	1.68	Ι	15	1500.0	26700	98.97	I FALS	E 0.6	596 0.02	4 0.078 0.0	0.1 0.3	325 0.06	39 0.00	21 73	38 70	535	14	487	~	91.0
SMB17-234 - 23	2.17E+08	254	155.9	1.63	9	11	122.2	3124	98.89	I FALS	E 0.7	749 0.01	7 0.082 0.0	0 I O.4	417 0.06	58 0.00	13 8(00 41	567	10	508	~	89.7
SMB17-234 - 24	2.03E+08	598.1	168	3.56	14	15	107.1	11743	98.82	I FALS	E 2.6	538 0.03	8 0.210 0.0	03 0.2	236 0.09	04 0.00	09 143	33 19	1311	11	1229	14	85.8
SMB17-234 - 25	1.97E+08	986	116	8.50	21	14	66.7	14571	97.52	I FALS	E 3.8	829 0.04	5 0.250 0.0	03 0.2	11.0 602	03 0.00	08 18()4 13	1598	10	1436	16	79.6
SMB17-234 - 26	2.19E+08	1130	248	4.56	97	21	21.6	1344	97.46	I FALS	E 0.9	957 0.02	4 0.088 0.0	0.0 0.0	573 0.07	80 0.00	13 114	47 33	681	12	545	12	80.0
SMB17-234 - 27	1.97E+08	925	147.9	6.25	7	14	700.0	115300	98.83	I FALS	E 2.4	447 0.03	9 0.200 0.0	03 0.2	770 0.08	81 0.00	08 138	84 17	1256	12	1175	15	84.9
SMB17-234 - 28	1.56E+08	1284	476	2.70	94	33	35.1	3768	97.89	I FALS	E 4.1	180 0.05	8 0.269 0.0	03 0.4	149 0.11	19 0.00	12 183	81 19	1670	11	1533	17	83.7
SMB17-234 - 29 ^a	2.00E+08	362.8	79.3	4.58	11	13	118.2	9373	99.37	I FALS	E 2.7	784 0.04	9 0.225 0.0	03 0.0	506 0.08	92 0.00	11 14(38 24	1350	13	1307	17	92.8
SMB17-234 - 30 ^{a,t}	2.23E+08	246.6	167.8	1.47	-4	14	-350.0	27270	99.65	I FALS	E 0.7	700 0.02	0 0.086 0.0	0.1 0.2	212 0.05	90 0.00	17 56	57 63	538	12	531	~	98.6
SMB17-234 - 31 ^{a,t}	2.41E+08	316.9	201.7	1.57	1-	20	-2000.0	36090	99.63	I FALS	E 0.6	587 0.01	8 0.084 0.0	0.0 1.00	0.0 200	92 0.00	16 57	74 59	530	11	521	~	98.3
SMB17-234 - 32 ^a	1.99E+08	152.5	13.15	11.60	I	15	1500.0	24200	99.54	I FALS	E 1.1	174 0.03	1 0.129 0.0	02 0.	252 0.06	59 0.00	16 8(<i>33 51</i>	787	14	780	11	97.1
SMB17-234 - 33	1.88E+08	218.4	94.9	2.30	12	17	141.7	6342	97.87	I FALS	E 4.6	512 0.08	4 0.285 0.0	05 0.0	538 0.11	66 0.00	15 190)5 <i>2</i> 3	1750	15	1615	23	84.8
SMB17-234 - 34	2.15E+08	406	225	1.80	7	10	500.0	66100	98.28	I FALS	E 3.3	381 0.04	5 0.239 0.0	03 0.5	545 0.10	17 0.00	10 165	56 18	1499	10	1383	16	83.5
SMB17-234 - 35	2.47E+08	965	44	21.93	215	31	14.4	1774	96.17	ع	5.5	209 0.08	7 0.295 0.0	0.4 0.6	570 0.12	81 0.00	13 207	72 18	1854	14	1665	19	80.4
SMB17-234 - 36	2.48E+08	778	32.8	23.72	9	28	466.7	23500	99.46	1 FALS	E 1.3	325 0.08	5 0.136 0.0	05 0.9	957 0.07	08 0.00	22 95	52 64	853	36	819	27	86.1
SMB17-234 - 37 ^a	2.25E+08	399	127	3.14	7-	13	-650.0	175600	98.77	1 FALS	E 5.8	867 0.06	7 0.337 0.0	0.4 0.5	564 0.12	58 0.00	09 204	40 12	1956	10	1874	19	91.9
SMB17-234 - 38	2.00E+08	423	109.1	3.88	71	19	26.8	2146	97.04	1 FALS	E 4.7	740 0.06	9 0.279 0.0	0.4 0.6	533 0.12	22 0.00	361 11	39 16	1774	12	1587	18	79.8
SMB17-234 - 39 ^a	2.13E+08	59.9	53.88	1.11	8	11	137.5	3304	99.21	I FALS	E 5.9	950 0.11	0 0.349 0.0	05 0.4	487 0.12	35 0.00	18 200	07 26	1965	16	1928	23	96.0
SMB17-234 - 40 ^a	2.03E+08	161.8	98.3	1.65	12	13	108.3	1429	99.18	I FALS	E 0.7	731 0.02	2 0.084 0.0	0.0 1.00	94 0.06	33 0.00	20 71	18 67	558	13	520	~	93.2
SMB17-234 - 41	2.02E+08	644.6	555	1.16	9	18	300.0	42733	98.20	I FALS	E 5.1	144 0.06	5 0.308 0.0	0.4 0.4	413 0.12	07 0.00	11 196	57 16	1843	11	1729	18	87.9
SMB17-234 - 42 ^a	2.13E+08	142.6	97.8	1.46	10	13	130.0	1566	98.89	1 FALS	E 0.7	734 0.02	5 0.082 0.0	0 I O.4	442 0.06	51 0.00	19 77	78 61	557	15	507	~	91.0
SMB17-234 - 43 ^a	2.13E+08	172.3	90.7	1.90	Э	11	366.7	15570	99.24	1 FALS	E 2.5	543 0.04	9 0.211 0.0	03 0.3	30.0 868	:76 0.00	13 137	74 29	1283	12	1236	14	90.0
SMB17-234 - 44 ^a	2.11E+08	149.2	105.9	1.41	8	11	137.5	8850	99.78	I FALS	E 6.5	540 0.09	0 0.374 0.0	05 0.5	544 0.12	68 0.00	12 205	54 17	2050	12	2047	21	99.7
SMB17-234 - 45 ^a	2.14E+08	182.2	98.3	1.85	I	12	1200.0	19070	99.38	I FALS	E 0.6	584 0.01	9 0.081 0.0	0.0 0.0	947 0.06	07 0.00	16 62	29 57	528	11	505	9	95.5
SMB17-234 - 46 ^a	2.17E+08	1280	31	41.29	16	10	62.4	10580	99.75	I FALS	E 0.8	826 0.01	1 0.098 0.0	0 I O	427 0.06	14 0.00	06 65	53 21	611	6	601	~	98.3
SMB17-234 - 47	2.22E+08	491	126.4	3.88	22	13	59.1	7423	97.79	I FALS	E 3.7	747 0.04	3 0.250 0.0	03 0.5	550 0.10	89 0.00	07 178	82 12	1581	9	1436	15	80.6
SMB17-234 - 48	2.31E+08	690.4	48.2	14.32	35	25	71.4	3580	97.91	I FALS	E 1.7	721 0.05	5 0.146 0.0	03 0.5	90. 0.0e	58 0.00	14 133	34 32	1015	21	880	17	<i>56.0</i>
SMB17-234 - 49 ^a	2.23E+08	314.1	149.8	2.10	~	13	185.7	4967	99.72	I FALS	E 0.6	10.0 265	4 0.086 0.0	0.1 0.2	223 0.05	88 0.00	11 56	50 41	537	8	533	~	99.3
SMB17-234 - 50	1.99E+08	135.7	53.6	2.53	12	12	100.0	3246	97.76	I FALS	E 3.3	358 0.06	5 0.232 0.0	0.4 0.5	561 0.10	48 0.00	15 171	11 26	1493	15	1347	18	78.7
SMB17-234 - 51	2.18E+08	189.9	109.1	1.74	25	12	48.0	815	98.78	I FALS	E 0.7	758 0.01	8 0.083 0.0	0.1 0.	129 0.06	68 0.00	15 83	32 47	572	10	511	9	89.4
SMB17-234 - 52	2.45E+08	773	118.9	6.50	33	25	75.8	3645	98.40	I FALS	E 1.2	245 0.03	4 0.119 0.0	0.2 0.5	588 0.07	68 0.00	11 91	16 42	821	15	722	11	87.9
SMB17-234 - 53 ^{a,t}	2.14E+08	196.2	107.6	1.82	~	11	157.1	3050	99.64	I FALS	E 0.6	10 ^{.0} 069	9 0.086 0.0	0.1 0.7	279 0.05	84 0.00	15 54	45 56	532	11	531	~	9.66
SMB17-234 - 54 ^a	2.05E+08	969	47.6	14.62	37	22	59.5	2381	99.26	I FALS	E 1.1	128 0.03	7 0.118 0.0	02 0.5	598 0.06	98 0.00	18 92	22 53	766	18	720	12	94.0
SMB17-234 - 55 ^a	2.13E+08	687.1	98.1	7.00	12	15	125.0	20008	99.26	I FALS	E 3.6	529 0.05	3 0.262 0.0	03 0.0	501 0.10	04 0.00	10 163	31 18	1555	12	1502	16	92.1
SMB17-234 - 56	2.08E+08	415.8	243.6	1.71	9	15	166.7	12722	98.52	I FALS	E 2.7	737 0.04	0 0.212 0.0	03 0.5	589 0.0	40 0.00	10 150	38 20	1338	11	1239	16	82.2

Corrections: 1 = t	hreshold ²⁰⁴ 1	Ph no cor	rection (100 cps	s); 2 = tl	nreshold %	6 ²⁰⁴ Pb-base	ed correctio	on(20%	error);	3 = thre	shold %	for ²⁰⁸ F	b-based	corre	tion (9	8.5 %Pb'	.(
														Isotopi	c ratio				0	alculat	ed age	Se		
Sample	90 Zr	D	Th	Th/II	^{204}Pb	$\pm 2\sigma$	% ± 21	$^{6}\mathrm{Pb}/^{204}\mathrm{Pb}$	%	Correct	ions 20	⁷ Pb/ +	رہ ²⁰⁶ ا	2 + 7	ע EC	$^{-3}$ 207 P	b/ + 2	ر ²⁰⁷ Pl	P/ ±	207 Pb/	+1	²⁰⁶ Pb/	+1	%
ardumo	cps	(mqq)	(mqq)	0	cps ¹ 2	²⁰⁴ Pb cps	²⁰⁴ Pb cps	cps	Pb^{\star^2}	1001100	2.	³⁵ U ⁻ .	238	U -	3	, ²⁰⁶ I	₽ P	²⁰⁶ P	$b 2\sigma$	235 U	2σ	238 U	2σ 6	conc
SMB17-234 - 57	2.17E+08	382.7	137.8	2.78	56	16	28.6	2825	97.69	I FA	LSE 5	.233 0.0	84 0.3	0.0 0.00	4 0.8	1 0.12	48 0.00	11 202	6 16	1859	15	1720	22	84.9
SMB17-234 - 58 ^a	2.04E+08	209.7	118.4	1.77	6	12	200.0	3578	99.15	I FA	LSE 0	.681 0.0	19 0.0	178 0.00	1.0.1	56 0.06	29 0.00	17 76	5 58	526	11	487	~	92.6
SMB17-234 - 59 ^a	2.20E+08	625	596	1.05	21	12	57.1	2919	99.36	I FA	LSE 0	.638 0.0	11 0.0	76 0.00	1 0.40	0.06	16 0.00	08 65	9 28	501	\sim	470	9	93.8
SMB17-234 - 60	2.10E+08	388.8	150.9	2.58	36	15	41.7	3575	94.59	I FA	LSE 4	.831 0.1	00 0.2	56 0.00	5 0.8.	9 0.13	73 0.00	16 219	3 20	1788	61	1467	27	60.9
SMB17-234 - 61 ^a	1.96E+08	162.8	94.5	1.72	8	17	212.5	2066	99.48	I FA	LSE 0	.632 0.0	29 0.0	178 0.00	2 0.0	22 0.05	87 0.00	28 55	6 104	496	18	486	10	97.9
SMB17-234 - 62 ^a	2.46E+08	814	20.9	38.95	18	20	111.1	6350	99.49	I FA	LSE 0	.917 0.0	21 0.1	03 0.00	1 0.2	27 0.06	53 0.00	16 78	84 51	663	12	633	8	95.5
SMB17-234 - 63	2.46E+08	477.9	439.7	1.09	33	19	57.6	7045	95.52	I FA	LSE 7	760 0.1	60 0.3	56 0.00	5 0.70	64 0.15	95 0.00	20 245	0 21	2201	19	1962	22	80.1
SMB17-234 - 64	2.59E+08	843.4	87.4	9.65	I	23	2300.0	390300	95.86	I FA	LSE 6	.923 0.1	10 0.3	36 0.00	5 0.6	0 0.15	08 0.00	15 235	5 17	2101	14	1868	24	79.3
SMB17-234 - 65 ^a	2.21E+08	510	219	2.33	<i>I</i> -	10	-1000.0	163200	99.51	I FA	LSE 3	.110 0.0	44 0.2	43 0.00	3 0.7	60.09	32 0.00	07 145	2 14	1434	11	1402	17	94.0
SMB17-234 - 66 ^a	1.96E+08	156	97.6	1.60	9	14	233.3	3527	99.16	I FA	LSE 1	.020 0.0	36 0.1	00 0.00	2 0.1	<i>t8 0.06</i>	75 0.00	22 85	3 68	711	18	668	10	93.9
SMB17-234 - 67	2.11E+08	525	80.5	6.52	6	10	166.7	13583	98.14	I FA	LSE 1	.348 0.0	21 0.1	23 0.00	2 0.4	55 0.07	95 0.00	10 118	34 24	866	9	751	9	63.4
SMB17-234 - 69	2.44E+08	1631	765	2.13	65	21	32.3	5858	97.19	I FA.	LSE 2	.223 0.0	35 0.1	69 0.00	2 0.7	36 0.09	61 0.00	08 154	9I 6i	1188	11	1006	13	64.9
SMB17-234 - 70	2.18E+08	1136	109.3	10.39	ŝ	13	433.3	89100	98.87	I FA	LSE 2	.343 0.0	28 0.1	94 0.00	2 0.52	25 0.08	81 0.00	07 138	84 15	1225	9	1145	13	82.8
SMB17-234 - 71 ^a	2.20E+08	602	32.3	18.64	29	16	55.2	3052	98.90	I FA	LSE 1	.126 0.0	26 0.1	14 0.00	2 0.4	15 0.07	17 0.00	14 97	7 40	765	13	698	6	91.2
SMB17-234 - 72	2.12E+08	299.6	82	3.65	S	12	240.0	8940	98.51	I FA	LSE 1	.188 0.0	29 0.1	15 0.00	2 0.48	30 0.07	49 0.00	14 106	6 38	794	13	702	6	88.4
SMB17-234 - 73	1.94E+08	944	200.3	4.71	36	15	41.7	0006	98.77	I FA	LSE 4	.174 0.0	49 0.2	79 0.00	3 0.7(0 0.10	87 0.00	07 177	7 11	1669	10	1586	17	89.2
SMB17-234 - 74	2.30E+08	270.1	70.8	3.81	19	15	78.9	2253	97.79	I FA	LSE 1	379 0.0	35 0.1	22 0.00	2 0.3(0.08	24 0.00	19 125	5 45	879	15	744	10	84.6
SMB17-234 - 75	1.84E+08	751	31.3	23.99	64	16	25.0	1630	98.57	I FA	LSE 1	.315 0.0	18 0.1	26 0.00	2 0.4	12 0.07	63 0.00	08 110	3 22	853	8	762	10	69.1
SMB17-234 - 76 ^a	2.04E+08	300	91.4	3.28	9	13	216.7	20533	98.62	I FA	LSE 5	.158 0.0	82 0.3	13 0.00	5 0.7	35 0.11	95 0.00	11 194	61 I G	1845	14	1756	24	90.1
SMB17-234 - 77 ^a	2.13E+08	163.4	102.1	1.60	\mathcal{O}	10	333.3	5850	99.47	I FA	LSE 0	.703 0.0	21 0.0	85 0.00	1 0.38	35 0.05	99 0.00	16 60	0 58	538	12	527	~	97.9
SMB17-234 - 78 ^a	1.52E+08	223.9	190.8	1.17	6	17	283.3	2963	99.22	I FA	LSE 0	.735 0.0	24 0.0	84 0.00	1.0.1	94 0.06	33 0.00	20 71	8 67	558	14	522	6	93.6
SMB17-234 - 79 ^a .	^b 2.13E+08	224.1	131.7	1.70	-9	10	-176.4	24200	99.57	I FA	LSE 0	.695 0.0	16 0.0	85 0.00	1 0.0	14 0.05	95 0.00	14 58	85 51	535	10	526	6	98.5
SMB17-234 - 80	1.93E+08	186.2	152.3	1.22	-4	10	-250.0	55100	99.85	I FA	LSE 2	.905 0.0	49 0.2	42 0.00	3 0.3	80 <i>.</i> 08	68 0.00	13 135	6 29	1382	13	1397	17 1	03.0
SMB17-234 - 81 ^a	2.34E+08	135	88	1.53	11	13	118.2	9018	98.56	I FA	LSE 16	580 0.2	20 0.5	53 0.00	8 0.7	53 0.21	88 0.00	20 297	2 15	2910	13	2838	33	95.5
SMB17-234 - 82	2.29E+08	1251	405.6	3.08	28	16	57.1	9925	98.56	I FA	LSE 1	.963 0.0	23 0.1	69 0.00	2 0.50	9 0.08	49 0.00	06 131	4 14	1103	8	1004	11	76.4
SMB17-234 - 83 ^a	2.10E+08	208.7	218	0.96	9	9	164.9	4198	99.07	I FA	LSE 0	.834 0.0	21 0.0	92 0.00	1 0.2	12 0.06	60 0.00	15 80	6 48	615	11	568	~	92.4
SMB17-234 - 84 ^a	2.06E+08	261	209	1.25	б	13	433.3	31933	99.03	I FA	LSE 4	.249 0.0	65 0.2	86 0.00	4 0.70	52 0.10	78 0.00	13 176	3 22	1683	13	1620	21	91.9
SMB17-234 - 85 ^a	2.08E+08	276.8	198.8	1.39	I	11	1100.0	28610	99.38	I FA	LSE 0	.702 0.0	14 0.0	82 0.00	1 0.37	77 0.06	21 0.00	11 67	78 38	539	6	511	~	94.8
SMB17-235 Run	1 Grid Zon	e 21T (W	GS84) 28	86668	509950	4																		
SMB17-235 - 7	2.32E+08	721	92.4	7.80	45	31	68.9	2111	97.75	I FA	LSE 1	.418 0.0	50 0.	125 0.00	0.7	30.0 86	30 0.00	16 120	59 38	896	21	757	16	59.6
SMB17-235 - 1	2.34E+08	595	83.4	7.13	-15	32	-213.3	106330	98.19	1 FA	LSE 1	.995 0.0	72 0.	166 0.00	33 0.6	0.0 40	73 0.00	21 130	57 46	1113	25	992	19	72.6
SMB17-235 - 2 ^a	1.95E+08	468.3	10.98	42.65	4	10	152.3	7771	99.90	1 FA	LSE C	.864 0.0	0.25	104 0.00	0.3	84 0.06	04 0.00	11 6.	18 39	632	14	636	6	00.7
SMB17-235 - 3 ^a	2.10E+08	149.4	101.8	1.47	18	18	100.0	781	99.68	1 FA	LSE C	0.725 0.0	34 0.0	0.0 060	0.2	49 0.05	92 0.00	26 57	74 95	555	22	554	10	6.66
SMB17-235 - 4 ^a	1.85E+08	287.2	133	2.16	ю	14	466.7	31200	99.70	1 FA	LSE 4	1.969 0.1	30 0.	318 0.00	0.5	39 0.11	32 0.00	17 185	51 27	1813	23	1780	25	96.1
SMB17-235 - 5 ^a	1.92E+08	138.6	117	1.18	ή	14	-466.7	13800	99.71	1 FA	LSE C	0.762 0.0	31 0.(0.0	0.2 0.2	17 0.05	0.0 66	20 6(00 72	574	19	571	6	99.4
SMB17-235 - 6 ^a	1.94E+08	223.3	312.1	0.72	ŝ	13	260.0	17156	99.39	I FA	LSE 6	674 0.1	80 0.3	372 0.00	0.7	44 0.13	05 0.00	17 210	05 23	2068	23	2037	31	96.8

Corrections: 1 = tl	rreshold ²⁰⁴ 1	p no cori	ection (]	100 cps)); 2 = th	reshold %	²⁰⁴ Pb-based	correction	(20% e	error);3 = tl	reshol	d % foi	. ²⁰⁸ Pb-bas	ed corr	ection	98.5 %Pb*).						
													Isotoj	oic rati	so			Ű	alculate	ed age			
Sample	90 Zr	U (maa)	Th (mon)	Th/U	²⁰⁴ Pb	$\pm 2\sigma$	$\% \pm 206_{10}$	b/ ²⁰⁴ Pb	% h*2	Corrections	²⁰⁷ Pb/	± 2σ	²⁰⁶ Pb/ ±	2σ F	C ³ 20	⁷ Pb/ ± 20	3 ²⁰⁷ Pl	o/ ± - 2σ	²⁰⁷ Pb/ ²³⁵ 11	+ 2α	⁰⁶ Pb/ ²³⁸ 11	+ 2α c	%
SMB17-235 - 8 ^a	2 04F+08	1683	7 38	22.80	-2	13	-650.0	20680 9	9 58	1 FALSE	1 056	0.037	0116 0	0.200	295 0	ru 0660_000	17 F	и 54	730	18	707	12	6 96
SMB17-235 - 9	2.15E+08	419.2	145	2.89	78	22	28.2	1098 9	8.22	1 FALSE	2.372	0.065	0.189 0.	003 0	330 0.	0916 0.00	15 145	59 31	1233	20	1114	19	76.4
SMB17-235 - 10 ^a	1.91E+08	111	87.9	1.26	22	12	54.5	479 9	9.57	1 FALSE	0.762	0.034	0.092 0.	002 0	080 0	0604 0.003	24 6.	18 86	572	19	566	10	0.66
SMB17-235 - 11 ^a	1.90E+08	516	274.2	1.88	-2	13	-650.0	158800 9	9.93	1 FALSE	4.416	0.110	0.305 0.	005 0	819 0	1054 0.00	13 172	21 23	1714	22	1713	25	9.5
SMB17-235 - 12 ^a	1.85E+08	355.3	198.9	1.79	6	12	133.3	3418 9	9.55	1 FALSE	0.728	0.024	0.086 0.	001 0	173 0.	0611 0.00	16 6	ł3 56	555	14	534	~	96.3
SMB17-235 - 13	1.84E+08	245.3	64.3	3.81	~	13	185.7	4776 9	9.39	1 FALSE	1.336	0.042	0.136 0	002 0	325 0.	0717 0.00	16 97	77 45	862	19	820	12	33.9
SMB17-235 - 14 ^a	1.76E+08	144.1	21.54	69.9	16	14	87.5	1185 9	9.68	1 FALSE	1.163	0.045	0.130 0.	003 0	170 0.	0654 0.00	23 78	87 74	784	23	787	14 10	0.00
SMB17-235 - 15 ^a	1.69E+08	164.3	94.6	1.74	19	15	78.9	629 9	8.92	1 FALSE	0.711	0.040	0.080 0.	002 0	135 0.	0647 0.003	34 70	55 111	543	24	495	Π	91.2
SMB17-235 - 16a	2.31E+08	865	73	11.85	42	30	71.4	2462 9	9.36	1 FALSE	1.033	0.037	0.111 0	002 0	521 0.	0675 0.00	17 85	53 52	720	18	680	13	94.5
SMB17-235 - 16a ^a	1.69E+08	204.7	113.2	1.81	14	15	107.1	3143 9	9.56	1 FALSE	2.558	0.078	0.217 0.	004 0	541 0.	0853 0.00	18 132	22 41	1287	22	1265	21	95.7
SMB17-235 - 17 ^{a,b}	1.92E+08	193.9	122.1	1.59	-1	11	-1100.0	16950 9	9.67	1 FALSE	0.700	0.024	0.086 0.	001 0	154 0.	0592 0.00	17 5.	74 62	537	15	530	6	98.8
SMB17-235 - 18 ^a	1.81E+08	366.9	180.1	2.04	-5	13	-260.0	100120 9	9.79	1 FALSE	3.787	0.098	0.277 0.	004 0	536 0.	00.0 0660	14 16()6 26	1591	20	1575	22	98.1
SMB17-235 - 19	1.95E+08	225	131.8	1.71	20	13	65.0	6 696	8.09	1 FALSE	0.842	0.029	0.084 0.	002 0	194 0.	0728 0.00	21 100	8 59	620	17	521	6	34.0
SMB17-235 - 20a	2.20E+08	1171	8.93	131.13	27	17	63.0	4896 9	9.39	1 FALSE	0.997	0.028	0.108 0.	002 0	197 0.	0665 0.00	13 82	22 41	702	14	664	10	94.6
SMB17-235 - 20a	1.51E+08	590	41.5	14.22	45	22	48.9	1789 9	8.37	1 FALSE	1.715	0.047	0.151 0	003 0	396 0.	0823 0.00	13 125	52 31	1014	18	904	16	72.2
SMB17-235 - 21 ^a	1.89E+08	246	246.7	1.00	ю	14	466.7	7743 9	9.71	1 FALSE	0.764	0.025	0.092 0.	002 0	152 0.	0597 0.00	15 59	93 54	575	14	567	6	98.6
SMB17-235 - 22 ^{a,b}	1.86E+08	227.5	226.4	1.00	-5	12	-240.0	19600 9	99.66	1 FALSE	0.705	0.023	0.086 0.	001 0	005 0	0293 0.00	16 53	78 59	543	14	533	6	98.2
SMB17-235 - 23 ^a	1.67E+08	208.6	89.8	2.32	22	14	63.6	2702 9	8.84	1 FALSE	4.647	0.130	0.297 0.	005 0	334 0.	1123 0.00	22 183	37 35	1756	24	1675	24	91.2
SMB17-235 - 24 ^a	1.89E+08	199.8	67.1	2.98	2	12	600.0	8970 9	9.76	1 FALSE	0.719	0.027	0.089 0.	001 0	209 0.	0581 0.00	18 53	34 68	548	16	549	8 10	0.2
SMB17-235 - 25 ^a	1.93E+08	329	196	1.68	6	11	122.2	3100 9	9.71	1 FALSE	0.702	0.024	0.085 0.	002 0	442 0	0592 0.00	14 57	74 51	539	14	528	6	97.9
SMB17-235 - 26 ^a	1.89E+08	128.3	134	0.96	12	11	91.7	978 9	9.56	1 FALSE	0.764	0.029	0.090 0.	002 0	400 0.	0607 0.00	19 62	29 67	576	17	558	6	96.9
SMB17-235 - 27	2.20E+08	131.9	23.56	5.60	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	18	-225.0	21470 9	7.98	1 FALSE	1.867	0.085	0.156 0.	004 0	437 0.	0860 0.000	32 133	38 72	1067	31	935	19	6.65
SMB17-235 - 28 ^a	1.60E+08	707	251.8	2.81	216	45	20.8	790 9	7.36	- 2	3.350	0.310	0.252 0	0 900	741 0.	00.0 956 0.00	35 154	40 69	1495	69	1449	33	94.1
SMB17-235 - 29 ^a	1.75E+08	113.6	43.92	2.59	1	10	1000.0	34400 9	9.87	1 FALSE	4.946	0.140	0.322 0.	005 0	365 0.	1103 0.00	20 18()4 33	1808	24	1798	26	9.66
SMB17-235 - 30 ^a	1.73E+08	192.2	154.5	1.24	17	13	76.5	6 696	9.71	1 FALSE	0.722	0.029	0.089 0.	002 0	521 0.	0583 0.00	19 5-	H 71	550	17	548	6	99.5
SMB17-235 - 31 ^a	1.72E+05	213000	<lod< td=""><td>na</td><td>4</td><td>12</td><td>300.0</td><td>5438 9</td><td>9.72</td><td>1 FALSE</td><td>0.967</td><td>0.032</td><td>0.112 0.</td><td>002 0</td><td>307 0.</td><td>0626 0.00</td><td>15 69</td><td>95 51</td><td>685</td><td>16</td><td>684</td><td>Ξ</td><td>6.66</td></lod<>	na	4	12	300.0	5438 9	9.72	1 FALSE	0.967	0.032	0.112 0.	002 0	307 0.	0626 0.00	15 69	95 51	685	16	684	Ξ	6.66
SMB17-235 - 32 ^a	1.86E+08	141.1	57.39	2.46	Ŋ	11	220.0	7466 9	9.61	1 FALSE	3.605	0.100	0.268 0.	004 0	386 0.	00.0 9960	17 155	59 33	1548	23	1532	22	98.2
SMB17-235 - 33 ^{a,b}	2.09E+08	79.75	41.79	1.91	2	15	750.0	3611 9	69.60	1 FALSE	0.705	0.044	0.086 0.	002 0	309 0	0586 0.003	33 55	52 123	537	26	531	10	98.9
SMB17-235 - 34 ^a	1.80E+08	220	86.3	2.55	9	14	233.3	9717 9	09.60	1 FALSE	3.459	0.097	0.262 0	004 0	378 0.	00.0 0.00	17 152	28 34	1518	23	1497	21	98.0
SMB17-235 - 35 ^a	1.93E+08	160.5	191.9	0.84	14	21	150.0	1071 9	9.70	1 FALSE	0.822	0.039	0.098 0.	002 0	008 0	0602 0.003	27 6.	11 97	607	22	604	12	99.5
SMB17-235 - 36	2.31E+08	769	108	7.12	112	37	33.0	1028 9	6.28	- 2	1.608	0.120	0.139 0	003 0	500 0.	0829 0.00	41 120	57 97	971	47	841	17	56.4
SMB17-235 - 37	2.02E+08	200	40.1	4.99	25	14	56.0	1140 9	98.6	1 FALSE	1.269	0.051	0.139 0.	003 0	149 0.	0658 0.00	20 8(00 64	830	23	836	16 10	04.5
SMB17-235 - 38 ^a	1.77E+08	98.8	82.8	1.19	4	13	185.7	1214 9	69.6	1 FALSE	0.733	0.034	0.090 0.	002 0	117 0.	0586 0.003	24 55	52 89	555	20	554	Ξ	8.66
SMB17-235 - 39 ^{a,b}	1.76E+08	81.1	59.22	1.37	1	12	1200.0	6853 9	9.50	1 FALSE	0.702	0.037	0.087 0.	002 0	045 0.	0587 0.00	30 55	56 111	535	23	535	11 10	0.00
SMB17-235 - 40 ^a	1.78E+08	175.9	132.1	1.33	1	14	1400.0	14540 9	9.52	1 FALSE	0.725	0.030	0.085 0.	001 0	317 0.	0612 0.003	21 64	1 6 74	552	18	527	8	95.4

eshold ²⁰⁴ Pb r	Pbr	10 COT.	rection (]	00 cps)	m = 7 :	resnoia %	21 Pb-base	l correctio	n(20%	error);3	i = three	shold % f	or ^{∠uo} Pl I	o-based sotopic	correcti ratios	on (98.5 ⁹	6Pb*).		Calcul	ated a	ges		
	$ \begin{array}{ccc} U & Th & Th \\ (ppm) & (ppm) & Th/U & ^{204}Pb & \pm 2\sigma & \% \\ & & & & & \\ ppm) & & & & & \\ ppm & & & \\ $	$ \begin{array}{ccc} Th & Th \\ (ppm) & Th/U & {}^{204}Pb & \pm 2\sigma & \% \\ cps^1 & cps^1 & {}^{204}Pb & cps & {}^{204}Pl \end{array} $	Th/U ²⁰⁴ Pb $\pm 2\sigma$ % cps ¹ ²⁰⁴ Pb cps ²⁰⁴ Pl	$^{04}\text{Pb} \pm 2\sigma \qquad \%$ $\text{cps}^{1 \ 204}\text{Pb} \text{ cps}^{204}\text{Pl}$	± 2σ % ⁴ Pb cps ²⁰⁴ Pl	4 Pl	± ²⁰⁶ 206	Pb/ ²⁰⁴ Pb cps	$\%$ Pb^{\star^2}	Correctio	ons 23	$\frac{Pb}{U} \pm 2$	ισ ²⁰⁶ Ρ΄ 2 ³⁸ ι	b/ ± 2σ J	EC ³	²⁰⁷ Pb/ ²⁰⁶ Pb	$\pm 2\sigma$	⁷ Pb/ : ⁰⁶ Pb 2	± ²⁰⁷ Pl	b/ ± J 2σ	²⁰⁶ Pb/ ²³⁸ U	2 H	% con
1.89E+08 176.2 120.34 1.46 4 12	176.2 120.34 1.46 4 12	120.34 1.46 4 12	1.46 4 12	4 12	12		300.0	29125	100.00	1 FAI	SE 23.	.300 0.59	90 0.6	51 0.010	0.756	0.2583	0.0031	3236	19 323	39 24	1 3232	39	6.66
2.26E+08 333.1 6.78 49.13 7 28	333.1 6.78 49.13 7 28	6.78 49.13 7 28	49.13 7 28	7 28	28		400.0	8986	98.12	1 FAI	SE 2	.332 0.1	10 0.13	85 0.00	5 0.536	0.0913	0.0031	1453	65 122	21 33	1094	25	75.3
1.80E+08 257.4 22 11.70 14 19 1	257.4 22 11.70 14 19 1	22 11.70 14 19 1	11.70 14 19 1	14 19 1	19 1	1	35.7	5439	98.91	1 FAL	SE 4	.526 0.1	40 0.2	93 0.00	5 0.688	0.1113	0.0021	1821	34 173	34 27	7 1657	31	91.0
1.77E+08 98.3 27.43 3.58 1 12 11	98.3 27.43 3.58 1 12 11	27.43 3.58 1 12 11	3.58 1 12 13	1 12 13	12 13	Ц	200.0	21000	99.64	1 FAI	SE 2	.435 0.08	82 0.2	13 0.00	4 0.296	0.0824	0.0021	1255	50 125	50 24	l 1247	20	99.4
1.50E+08 136.3 170.9 0.80 321 45	136.3 170.9 0.80 321 45	170.9 0.80 321 45	0.80 321 45	321 45	45		14.0	44	58.00	ب		.020 1.10	00 0.0	80 0.00	8 0.933	0.0590	0.0770	567 28	41 79	90 54(549	49	69.5
1.82E+08 169.9 103.33 1.64 -4 14 -3	169.9 103.33 1.64 -4 14 -3	103.33 1.64 -4 14 -3	1.64 -4 14 -3	-4 14 -3	14 -3	9	50.0	59770	99.43	1 FAI	SE 5	.831 0.10	60 0.3	45 0.00	5 0.497	0.1224	0.0020	1992	29 195	50 24	ł 1909	26	95.9
7.39E+07 113.2 72.4 1.56 26 15	113.2 72.4 1.56 26 15	72.4 1.56 26 15	1.56 26 15	26 15	15		57.7	169	95.30	1 FAI	SE 1	.151 0.08	82 0.0	89 0.00	4 0.385	0.0957	0.0068	1542 1	34 78	80 43	549	21	70.4
1.89E+08 125.9 72.8 1.73 4 12 3	125.9 72.8 1.73 4 12 3	72.8 1.73 4 12 3	1.73 4 12 3	4 12 3	12 3	Э	0.00	3118	99.64	1 FAL	SE 0	.842 0.0	34 0.0	00.0 66	2 0.397	0.0618	0.0020	667	69 62	20 18	909 8	10	97.7
1.85E+08 87.46 41.54 2.11 5 12 2	87.46 41.54 2.11 5 12 2	41.54 2.11 5 12 2	2.11 5 12 2	5 12 2	12 2	0	40.0	6226	99.78	1 FAI	SE 6	.023 0.13	70 0.3	58 0.00	5 0.466	0.1219	0.0022	1984	32 197	77 25	1977	30	9.66
1.71E+08 89.2 76.9 1.16 6 12 20	89.2 76.9 1.16 6 12 20	76.9 1.16 6 12 20	1.16 6 12 20	6 12 20	12 20	5	0.00	1165	99.70	1 FAI	SE 0	.707 0.03	37 0.0	86 0.00	2 0.174	0.0598	0.0029	596 1	05 53	38 22	534	Ξ	99.2
1.99E+08 176.2 92.2 1.91 3 17 56	176.2 92.2 1.91 3 17 56	92.2 1.91 3 17 56	1.91 3 17 56	3 17 56	17 56	56	6.7	5260	99.20	1 FAI	SE 0	.734 0.03	32 0.0	84 0.00	2 0.028	0.0639	0.0026	738	86 55	58 18	519	10	93.1
1.53E+08 122.1 48.5 2.52 16 17 10	122.1 48.5 2.52 16 17 10	48.5 2.52 16 17 10	2.52 16 17 10	16 17 10	17 10	10	6.3	604	98.51	1 FAI	SE 0	.832 0.0	47 0.03	87 0.00	2 0.037	0.0700	0.0040	928 1	17 61	12 26	537	12	87.7
(in italics to distinguish from Run 1)	to distinguish from Run 1)	guish from Run 1)	m Run 1)	1)																			
2.30E+08 326.8 235.1 1.39 28 14 50	326.8 235.1 1.39 28 14 50	235.1 1.39 28 14 50	1.39 28 14 50	28 14 50	14 50	5(0.0	1429	99.53	I FAL	SE 0.	745 0.01	16 0.08	88 0.001	0.220	0.0614	0.0012	653 4	12 56	5 9	546	\sim	96.6
2.19E+08 457.8 58.9 7.77 15 17 11.	457.8 58.9 7.77 15 17 11.	58.9 7.77 15 17 11.	7.77 15 17 11.	15 17 11.	17 11.	11.	3.3	12593	98.65	I FAL	SE 4.	890 0.07	71 0.30	4 0.004	0.711	0.1169	0.0010	606	5 180	0 12	1710	20	89.6
2.49E+08 134.9 144.4 0.93 13 15 11	134.9 144.4 0.93 13 15 11	144.4 0.93 13 15 11	0.93 13 15 11	13 15 11	15 11	11	5.4	1248	97.82	1 FAL	SE 0.	857 0.02	30 0.08	84 0.001	0.205	0.0752	0.0023	074 6	61 62	27 14	520	8	82.9
2.01E+08 428.4 103.8 4.13 54 25 4	428.4 103.8 4.13 54 25 4	103.8 4.13 54 25 4	4.13 54 25 4	54 25 4	25	4	1 6.3	1133	99.45	I FAL	SE 1.	007 0.02	3 0.11	1 0.002	0.129	0.0656	0.0017	794	4 70	17 12	680	10	96.2
2.27E+08 136.6 61.3 2.23 30 16 5	136.6 61.3 2.23 30 16 5	61.3 2.23 30 16 5	2.23 30 16 5	30 16 5	16 5	Ś	3.3	3153	98.77	1 FAL	SE 12.	880 0.15	0 0.45	7 0.002	0.361	0.1875	0.0022	2720 j	9 267	70 14	2600	31	95.6
2.21E+08 264.4 118.5 2.23 4 18 45	264.4 118.5 2.23 4 18 45	118.5 2.23 4 18 45	2.23 4 18 45	4 18 45	18 45	45	0.0	7575	99.20	1 FAL	SE 0.	719 0.02	30.0 IS	3 0.001	0.069	0.0631	0.0019	712 0	54 54	13 I I	512	8	93.2
2.04E+08 563.3 389 1.45 66 17 2	563.3 389 1.45 66 17 2	389 1.45 66 17 2	1.45 66 17 2	66 17 2	17 2		25.8	3258	94.46	I FAL	SE 6.	023 0.05	0 0.25	4 0.005	0.737	0.1481	0.0013	324	5 197	78 13	1661	22	71.5
2.17E+08 189.3 122.6 1.54 -1 12 -12	189.3 122.6 1.54 -1 12 -12	122.6 1.54 -1 12 -12	1.54 -1 12 -12	-1 12 -12	12 -12	-12	00.00	24730	99.18	I FAL	.SE 0.	893 0.02	25 0.05	100.0 6	0.253	0.0659	0.0018	803 5	57 64	14 i	606	8	93.3
1.95E+08 332.6 134.8 2.47 16 20 1	332.6 134.8 2.47 16 20 1.	134.8 2.47 16 20 1.	2.47 16 20 1.	16 20 1.	20 1.	1.	25.0	16500	96.30	1 FAL	SE 24.	670 0.37	70 0.63	1 0.005	0.808	0.2820	0.0021	3374]	2 329	5 15	3152	36	93.4
2.24E+08 79 46.2 1.71 1 13 130	79 46.2 1.71 1 13 130	46.2 1.71 1 13 130	1.71 1 13 130	1 13 130	13 130	130	0.0	9940	98.71	1 FAL	SE 0.	885 0.03	88 0.05	4 0.002	0.206	0.0683	0.0028	878 8	85 64	i0 20	576	9	90.0
1.94E+08 770 528 1.46 378 43 1	770 528 1.46 378 43 1.	528 1.46 378 43 1.	1.46 378 43 1	378 43 1.	43 1.	1	1.4	359	94.08	ى	<i>1</i> .	398 0.11	0 0.13	3 0.003	0.606	0.0741	0.0047	044 12	8 87	78 47	807	15	77.3
2.44E+08 905 254 3.56 937 54	905 254 3.56 937 54	254 3.56 937 54	3.56 937 54	937 54	54		5.8	369	91.85	۔ ئ	ŗ.	000 0.13	30 0.27	74 0.004	0.416	0.1320	0.0024	2125	182	1 21	1559	19	73.4
2.12E+08 313.7 246.8 1.27 203 26 i	313.7 246.8 1.27 203 26	246.8 1.27 203 26	1.27 203 26 1	203 26 1	26 1		2.8	526	95.81	- 3	ω.	390 0.18	80 0.24	16 0.005	0.669	0.0985	0.0036	1596 6	8 149	3 41	1416	28	88.7
2.16E+08 539 243 2.22 -1 12 -120	539 243 2.22 -1 12 -120	243 2.22 -1 12 -120	2.22 -1 12 -120	-1 12 -120	12 -120	-120	0.0	184700	99.82	I FAL	SE 3.	302 0.03	37 0.25	6 0.003	0.549	0.0929	0.0007	1486	3 148	6 I8	1472	15	99.0
2.27E+08 349 375 0.93 26 13 50	349 375 0.93 26 13 50	375 0.93 26 13 50	0.93 26 13 50	26 13 50	13 5(5(0.0	5496	99.69	1 FAL	SE 4.	707 0.05	59 0.31	1 0.004	0.560	0.1094	0.0009	1200	4 176	8 11	1744	18	97.4
2.68E+08 1164 312.6 3.72 337 37 11	1164 312.6 3.72 337 37 11	312.6 3.72 337 37 11	3.72 337 37 11	337 37 11	37 11	[]	0.	1060	95.46	- 3	ω.	094 0.05	2 0.21	1 0.004	1 0.766	0.1062	0.0016	1735 2	8 143	80 23	1233	20	71.1
2.18E+08 445 12.33 36.09 19 13 68	445 12.33 36.09 19 13 68	12.33 36.09 19 13 68	36.09 19 13 68	19 13 68	13 68	66	8.4	3284	99.51	1 FAL	SE 0.	944 0.01	6 0.10	0.001	0.140	0.0643	0.0010	752 3	82 67	74 9	647	\sim	95.9
2.24E+08 465 241.6 1.92 46 16 3	465 241.6 1.92 46 16 3	241.6 1.92 46 16 3	1.92 46 16 3	46 16 3	16 <i>3</i>	(1)	84.8	4787	98.80	I FAL	SE 5.	90.0 666	30 0.34	13 0.005	0.806	0.1262	0.0009	046	2 197	5 12	1899	22	92.8
2.43E+08 245.2 93.2 2.63 16 13	245.2 93.2 2.63 16 13	93.2 2.63 16 13	2.63 16 13	16 13	13		81.3	7713	99.60	I FAL	SE 6.	244 0.08	32 0.36	1 0.004	0.514	0.1254	0.0012	2034	7 201	0 12	1986	21	97.6
2.24E+08 2099 501.1 4.19 121 27	2099 501.1 4.19 121 27	501.1 4.19 121 27	4.19 121 27	121 27	27		22.3	6149	99.43		б.	387 0.04	17 0.25	6 0.003	0.786	0.0951	0.0005	1231	0 150	11 10	1468	17	95.9
2.31E+08 255.8 23.37 10.95 17 14	255.8 23.37 10.95 17 14	23.37 10.95 17 14	10.95 17 14	17 14	14	-	82.4	13476	98.67	1 FAL	SE 25.	840 0.27	70 0.66	1 0.008	8 0.679	0.2821	0.0014	3375	8 334	01 O	3272	30	97.0
2.25E+08 313.1 166.9 1.88 2 9 59	313.1 166.9 1.88 2 9 59	166.9 1.88 2 9 59	1.88 2 9 59	2 9 59	9 59	59	3.3	89933	99.86	I FAL	SE 4.	864 0.05	6 0.32	1 0.004	t 0.579	0.1097	0.0007	1.24	2 179	8	1794	18	100.0

BARR ET AL. – Ediacaran and Cambrian Rocks on Scatarie Island and nearby Hay Island, Avalonian Mira terrane, Cape Breton Island, Nova Scotia, Canada

Corrections: 1 = t	hreshold ²⁰⁴	Pb no cor	rection (100 cp	s); 2 = t	hreshold %	6 ²⁰⁴ Pb-basε	d correctic	on(20%	error);	3 = thre	shold % 1	or ²⁰⁸ P	b-based	correct	ion (98	.5 %Pb*)		C	-	-		╞	
														lsotopi	c ratios				ű	lculate	d ages			
Sample	90 Zr	D	Th	Th/U	204 Pb	±2σ	%± ²⁰	⁶ Pb/ ²⁰⁴ Pb	۔ %	Correc	tions 20	$^{7}Pb/ \pm 2$	²⁰⁶ Γ	'b/ ± 2	σ EC ³	207 Pf	ν/ ± 2σ	²⁰⁷ Pb	+	²⁰⁷ Pb/	≈ +	/qd ₉₀	+1	%
-	cps	(mqq)	(mqq)		cps ¹	²⁰⁴ Pb cps	²⁰⁴ Pb cps	cps	Pb^{*2}		2	³⁵ U	238	D	1	206 PI	0	^{206}Pb	2σ	²³⁵ U	5σ	²³⁸ U	2σ cc	onc
SMB17-235 - 27 ^a	2.24E+08	177.3	155.4	1.14	30	17	56.7	815	98.83	I FA	LSE 0	.963 0.03	32 0.1	0.0 IC	2 0.27.	5 0.068	9 0.002	2 896	99	683	17	619	9	0.7
SMB17-235 - 28 ^a	2.25E+08	112.3	91.1	1.23	4	11	275.0	20173	98.47	I FA	LSE 14	.480 0.20	0 0.5	21 0.00	17 0.75	5 0.200	6 0.001	6 2831	13	2780	13	2706	30 9	5.6
SMB17-235 - 29	2.15E+08	128	207.7	0.62	99	18	27.3	896	97.51	I FA	LSE 6	600 0.10	0 0.3	46 0.00	5 0.54	0.137	8 0.001	8 2200	23	2058	14	1913	22 8	7.0
SMB17-235 - 32 ^a	2.24E+08	105.2	71.8	1.47	~	11	157.1	7914	99.70	I FA	LSE 7	195 0.0	97 0.3	93 0.00	5 0.48	7 0.132	0.001	3 2137	17	2136	12	2139	23 10	0.1
SMB17-235 - 34 ^a	2.32E+08	951	158.1	6.02	524	40	7.6	327	94.69	ī	3 1	.092 0.03	73 0.1	23 0.00	2 0.53	4 0.063	6 0.003	2 728	: 107	745	36	748	12 10	0.3
SMB17-235 - 35	2.50E+08	566.1	120.6	4.69	174	26	14.9	1417	97.04	ı	3 4	.751 0.1	10 0.2	92 0.00	4 0.54	0.117	7 0.001	9 1922	29	1775	19	1653	20 8	6.0
SMB17-235 - 37 ^a	2.19E+08	128.8	62.4	2.06	I	11	1100.0	40700	99.73	I FA	LSE 2	.949 0.04	48 0.2	42 0.00	3 0.37.	3 0.088	5 0.001.	2 1393	26	1393	12	1396	16 10	0.2
SMB17-235 - 38 ^a	2.02E+08	163.2	105.9	1.54	31	17	54.8	3339	98.29	I FA	LSE 12	.840 0.22	20 0.4	91 0.00	17 0.65	0.189	1 0.002	1 2734	18	2667	16	2574	31 9	4.1
SMB17-235 - 39 ^a	2.53E+08	757	290	2.61	88	42	47.7	6318	98.99	I FA	LSE 15	.290 0.13	70 0.5.	34 0.00	7 0.54	4 0.208	3 0.001	2 2892	6	2833	11	2759	28 9	5.4
SMB17-235 - 41	2.17E+08	973	159.5	6.10	173	31	17.9	1343	97.41	ı	3	.094 0.00	52 0.1	77 0.00	2 0.40	7 0.085	5 0.002	0 1327	45	1145	21	1053	13 7	9.3
SMB17-235 - 42 ^a	2.19E+08	125.3	33.37	3.75	14	10	71.4	4514	99.14	I FA	LSE 7	024 0.10	0 0.3	78 0.00	5 0.59	7 0.135	0 0.001	4 2164	18	2113	13	2066	25 9	5.5
SMB17-235 - 43 ^a	2.26E+08	198.2	105.1	1.89	61	16	26.2	1625	98.61	I FA	LSE 6	870 0.12	20 0.3	56 <i>0.0</i> 0	5 0.60	1 0.136	7 0.001	9 2186	24	2095	16	2010	22 9	2.0
SMB17-235 - 44	2.29E+08	490	154.9	3.16	395	40	10.1	353	92.91		3	.660 0.14	40 0.1	94 0.00	3 0.64	2 0.09	0 0.003	7 1605	. 70	1309	40	1140	18 7	1.0
SMB17-235 - 45 ^a	2.46E+08	1049	81.5	12.87	105	23	21.9	1769	98.93		- I	.119 0.04	43 0.I.	24 0.00	2 0.37	9 0.065	4 0.002	0 787	64	765	19	756	10 9	6.0
SMB17-235 - 46	2.01E+08	354.1	49.2	7.20	17	14	82.4	12271	96.19	I FA	ILSE 11	.890 0.14	40 0.4	55 0.00	6 0.64	7 0.189	6 0.001	4 2739	12	2596	11	2417	26 8	8.3
SMB17-235 - 47 ^a	2.16E+08	276.3	115.1	2.40	384	39	10.2	320	94.67		35	.170 0.25	50 0.3	19 0.00	6 0.69	2 0.116	8 0.004	3 1908	99	1838	42	1786	29 9	3.6
SMB17-235 - 48	2.35E+08	687	40.7	16.88	27	15	55.6	4344	98.56	I FA	ITSE 1	.348 0.02	21 0.1	27 0.00	2 0.14	4 0.077	0 0.001	0 1121	26	866	9	773	9 6	8.9
SMB17-235 - 49 ^a	2.26E+08	181.4	121.2	1.50	23	14	60.9	4035	99.83	I FA	TSE 6	381 0.0	33 0.3	70 0.00	5 0.61.	3 0.125	5 0.001.	2 2036	17	2029	13	2028	22 9	9.6
SMB17-235 - 50 ^a	2.61E+08	518	34.8	14.89	15	27	180.0	5200	99.20	I FA	ITSE I	.059 0.04	<i>11 0.1</i>	12 0.00	2 0.39	2 0.069	3 0.002	5 908	74	733	20	684	6 11	3.3
SMB17-235 - 51 ^a	2.10E+08	826	49.7	16.62	7	16	228.6	47314	99.50	I FA	LSE 4	.611 0.05	59 0.3	0.00	4 0.68	7 0.110	0 0.000	9 1800	15	1751	11	111	20 9	5.1
SMB17-235 - 54	1.98E+08	698	317	2.20	159	24	15.1	1928	94.21		3	466 0.12	20 0.3	41 0.00	5 0.69	8 0.158	7 0.001.	5 2442	16	2168	15	1889	22 7	7.4
SMB17-235 - 55	1.72E+08	258.6	227.3	1.14	175	28	16.0	441	93.44		3 4	.280 0.22	20 0.2	59 0.00	5 0.68	5 0.118	8 0.004	5 1938	68	1680	43	1482	26 7	6.5
SMB17-235 - 56	2.12E+08	312.3	230	1.36	192	23	12.0	623	96.37	ī	3 4	.010 0.16	50 0.2	74 0.00	4 0.43	3 0.106	2 0.003	5 1735	60	1630	34	1558	19 8	9.8
SMB17-235 - 57	2.18E+08	356	57.4	6.20	156	22	14.1	1011	95.89	ī	3	.720 0.13	30 0.3	16 0.00	5 0.54	5 0.131	1 0.002	2 2113	29	1932	20	1768	23 8	3.7
SMB17-235 - 58	2.11E+08	475.3	67.2	7.07	73	18	24.7	1056	97.77	I FA	LSE I	.334 0.02	24 0.1	19 0.00	1 0.24	1 0.081	2 0.001.	2 1226	29	860	11	726	8 8	4.5
SMB17-235 - 59	2.19E+08	40	53.3	0.75	14	17	121.4	1557	94.83	I FA	LSE 10	.980 0.50	00 0.4	20 0.01	7 0.94	0.189	4 0.002	8 2737	24	2508	49	2256	81 8	2.4
SMB17-235 - 60 ^a	1.93E+08	228.3	73.8	3.09	ċ	14	-280.0	85600	99.72	I FA	LSE 4	.273 0.10	0 0.2	97 0.00	6 0.71	0.103	8 0.001.	5 1693	27	1686	20	1682	29 9	9.3
SMB17-235 - 62 ^a	2.13E+08	543	121.2	4.48	118	21	17.8	747	97.53	ī	3 I	.007 0.00	54 0.1	15 0.00	2 0.45	5 0.062	8 0.003	5 701	611	102	33	702	10 10	0.2
SMB17-235 - 63 ^a	2.14E+08	132	75.4	1.75	13	10	76.9	2623	99.30	I FA	LSE 2	.243 0.04	12 0.1	97 0.00	3 0.30	3 0.082	5 0.001	4 1257	33	1193	13	1159	16 9	2.2
SMB17-235 - 64 ^a	2.05E+08	561	14.8	37.91	42	18	42.9	6026	10.66	I FA	LSE 5	30.0 066.	81 0.3	44 0.00	5 0.78	8 0.125	7 0.000	9 2039	12	1974	12	1905	23 9	3.4
SMB17-235 - 65	1.91E+08	419	260	1.61	55	18	32.7	955	98.23	I FA	LSE 1	0.0 640	1.0 61	<i>32 0.00</i>	2 0.18	1 0.074	3 0.001	6 1050	43	728	10	627	9 8	6.1
SMB17-235 - 67	2.07E+08	55.69	37.86	1.47	19	15	78.9	1240	98.05	I FA	TSE 2	.620 0.10	50 0.3	25 0.00	6 0.66	7 0.125	8 0.002	5 2040	35	1913	25	1812	30 8	8.8
SMB17-235 - 68 ^a	1.86E + 08	1085	242.9	4.47	307	41	13.4	536	96.07	ī	3 1	.119 0.03	1.0 61	18 0.00	2 0.47	9 0.062	8 0.004	0 862	122	759	38	721	11 9	5.0
SMB17-235 - 69 ^a	2.16E+08	653.9	87.76	7.45	11	11	100.0	20255	99.66	I FA	LSE 3	.300 0.03	37 0.2	53 0.00	3 0.56	2 0.094	3 0.000	6 1513	: 13	1482	8	1452	15 9	6.0
SMB17-235 - 70 ^a	2.02E+08	327.6	174.1	1.88	74	18	24.3	1682	98.75	I FA	LSE 4	.887 0.00	54 0.3	<i>95 0.00</i>	4 0.45	5 0.115	7 0.001	1891 0	15	1799	11	1716	18 9	0.7
SMB17-235 - 71	2.45E+08	513	53	9.68	174	39	22.4	1484	92.85	ı	2	300 0.22	20 0.3	74 0.00	7 0.88	4 0.180	5 0.001	8 2657	17	2366	22	2046	31 7	7.0

													Isc	topic 1	atios				Calo	culated	ages		
Samule	90 Zr	D	Th	Th/IT	$^{204}\mathrm{Pb}$	$\pm 2\sigma$	% ± 2	^{:06} Pb/ ²⁰⁴ Pb	%	Correction	²⁰⁷ Pł	0/ + 7g	²⁰⁶ Pb/	νC +	EC3	²⁰⁷ Pb/	νC +	²⁰⁷ Pb/	± 20	Pb/	\pm ²⁰⁶ F	b/ ±	%
arduno	cps	(mqq)	(mqq)	0/111	cps ¹ 2	²⁰⁴ Pb cps ²	²⁰⁴ Pb cps	cps	Pb^{\star^2}		²³⁵ ر	FO	238 U	-	D'	$^{206}\mathrm{Pb}$	-	^{206}Pb	2σ ²	³⁵ U 2	(σ 238	U ^{2c}	conc
SMB17-235 - 72 ^a	1.98E+08	255.3	109.3	2.34	-5	11	-220.0	27690	99.30	I FALSI	E 0.74	6 0.020	0.086	0.001	0.067	0.0624	0.0017	688	58	564	12 5.	33	7 94.5
SMB17-235 - 73 ^a	2.26E + 08	38.6	19.34	2.00	ŗ.	10	-208.7	20420	98.81	1 FALS	3 7.52	0 0.170	0.392	0.007	0.489	0.1390	0.0024	2215	30	1213	21 21	33 3.	3 96.3
SMB17-235 - 74 ^a	2.32E+08	456	95.7	4.76	8	10	125.0	38563	99.42	1 FALSI	3 12.68	1 0.130	0.499	0.006	0.668	0.1842	0.0009	2691	8	656	10 26	38 2:	5 96.9
SMB17-235 - 76 ^a	2.19E+08	183.6	95.7	1.92	-9	14	-233.3	20310	99.24	I FALSI	E 0.65	0 0.020	0.081	0.001	0.139	0.0617	0.0017	664	59	532	12 5	32	7 94.3
SMB17-235 - 77	2.35E+08	455.5	119.5	3.81	160	29	18.1	1138	96.97	- 3	4.53	0 0.130	0.286	0.004	0.481	0.1143	0.0025	1869	39	738	22 16.	23 2.	86.8
SMB17-235 - 78	2.50E+08	1195	110.6	10.80	229	36	15.7	1937	92.13	۔ ع	5.60	0 0.120	0.265	0.004	0.841	0.1533	0.0017	2383	I 6I	915	19 15	13 1	63.5
SMB17-235 - 79 ^a	2.36E+08	946	13.9	68.06	21	16	76.2	6800	99.50	1 FALS	E 0.97	1 0.013	0.107	0.001	0.858	0.0654	0.0007	786	21	689	7 6.	58	7 95.5
SMB17-235 - 80	2.41E+08	426	83.2	5.12	60	19	31.7	2038	97.38	1 FALSI	3.05	3 0.061	0.214	0.004	0.823	0.1036	0.0011	1690	20	420	15 12	47 2.	2 73.8
SMB17-235 - 81 ^a	2.02E+08	168.1	121.9	1.38	10	12	120.0	6970	98.91	I FALSI	3 5.35	8 0.088	0.323	0.004	0.486	0.1197	0.0015	1952	22	877	14 18	<u> </u>	92.3
SMB17-235 - 82 ^a	2.11E+08	90.7	62.95	1.44	28	13	46.4	1450	98.91	I FALSI	5.83	8 0.100	0.340	0.005	0.558	0.1239	0.0016	2013	23	950	15 18	86 2,	t 93.7
SMB17-235 - 83	2.40E+08	426	189	2.25	374	43	11.5	452	93.67	- 3	4.99	0 0.200	0.286	0.006	0.688	0.1260	0.0031	2043	43	813	34 16	21 23	3 79.3
SMB17-235 - 84	1.90E+08	586.2	465	1.26	462	48	10.4	435	93.93	- 33	4.20	0 0.150	0.263	0.005	0.598	0.1146	0.0031	1874	49 j	. 671	30 15	<u> </u>	3 80.3
SMB17-235 - 85 ^a	2.27E+08	432.8	39.41	10.98	6-	13	-144.4	203100	99.69	1 FALS	3 5.78	1 0.065	0.346	0.004	0.548	0.1207	0.0009	1967	13	943	0 19	14 1	97.3
SMB17-235 - 86	2.20E+08	312.7	264.9	1.18	104	22	21.2	1886	96.08	- 2	11.54	0 0.200	0.454	0.006	0.475	0.1832	0.0022	2682	20	567	16 24	12 23	89.9
SMB17-235 - 87 ^a	2.09E+08	535.4	13.56	39.48	15	17	113.3	4911	99.22	I FALS	0.92	3 0.015	0.101	0.001	0.792	0.0658	0.0012	800	38	663	10 6	61	7 93.3
NOTES: ^a grains l second.	oetween 90 a	nd 102%	concord	lant; ^b gr	ains use	id in weigh	nted mean	calculation	s for yo	ungest clus	ters; ¹ af	ter Hg c	orrectio	n; ² radi	iogenic	Pb; ³ err	or correc	tion. Ab	brevia	tions: c	ps = co	unts p	er

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Corrections: 1 =	threshold 20	Pb no co	rrection	(100 cJ	os); 2 =	thresho	-qd ₁₀₂ % pl	based corre	ction(20	% error) ;	3 = threshold %	o for TVb	-based	correction	(98.5 %Pb	*).					-	ſ
	- 06		Ē		204	-	è	40 <i>C</i> 90C	2		207	²⁰⁶	opic rat	ios 207		207	ິບິ	Iculate	d ages		٦.	č
Sample	cps	(mqq)	(mqq)	Th/U	cps ¹	²⁰⁴ Pb	²⁰⁴ Pb cps	Pb cps	$\frac{70}{Pb^{*2}}$	Correction	s $\frac{PD}{235}$ $\pm 2a$	²³⁸ U	±2σ	^{3C³ ²⁰⁶P}	$b' \pm 2\sigma$	²⁰⁶ Pb	- 5α	²³⁵ U	5α H	³⁸ U	2α (-	% onc
Run 1 for both s	amples																					
FC-1 - 1	1.85E+08	317.2	191.6	1.66	12	12	100.0	4988	99.91	1 FALSE	1.950 0.05	3 0.186	0.003 (301 0.07	59 0.0013	1092	34	1097	18	1101	15	00.8
FC-1 - 2	1.83E+08	260.3	164	1.59		, 13	185.7	6950	99.79	1 FALSE	1.955 0.05	5 0.186	0.003 (.343 0.07	55 0.0014	1108	37	1099	19	1097	16	0.66
FC-1-3	1.84E+08	307.3	183.6	1.67	14	ال 12	85.7	4061	99.93	1 FALSE	1.961 0.05	4 0.186	0.003 (.302 0.07	53 0.0013	1103	34	1101	18	1102	15	9.66
FC-1 - 4	1.84E+08	336.7	159.83	2.11	13	13	100.0	4722	99.86	1 FALSE	1.933 0.05	1 0.185	0.003 (.105 0.07	58 0.0013	1090	34	1092	18	1093	15	00.2
FC-1 - 5	1.81E+08	408.4	260.8	1.57	-2	11	-550.0	73640	99.94	1 FALSE	1.969 0.05	2 0.187	0.003 (.305 0.07	57 0.0012	1113	31	1104	18	1103	15	99.1
FC-1 - 6	1.82E+08	178	77.75	2.29	4	ļ 12	300.0	8040	99.83	1 FALSE	1.957 0.05	8 0.187	0.003 (.284 0.07	58 0.0015	1090	40	1099	20	1102	16	01.1
FC-1 - 7	1.81E+08	157.4	62.75	2.51	-	12	-1200.0	27770	99.82	1 FALSE	1.930 0.05	8 0.185	0.003 (.284 0.07	53 0.0016	1077	43	1089	20	1092	16	01.4
FC-1 - 8	1.80E+08	284.9	169.9	1.68	ŝ	12	400.0	16860	16.66	1 FALSE	1.941 0.05	2 0.185	0.003 (.246 0.07	53 0.0012	1075	32	1096	19	1096	15	01.9
FC-1 - 9	1.79E+08	201.3	111.56	1.80	ų	10	-333.3	35900	99.82	1 FALSE	1.974 0.05	7 0.186	0.003 (.439 0.07	72 0.0014	1126	36	1105	20	1102	16	97.8
FC-1 - 10	1.93E+08	128.8	53.3	2.42	4	t 13	325.0	5843	99.73	1 FALSE	1.933 0.06	8 0.184	0.003 (0.057 0.07	70 0.0024	1121	62	1090	24	1090	18	97.2
FC-1 - 11	1.70E+08	473.6	294.2	1.61	4-	11	-275.0	82400	99.92	1 FALSE	1.960 0.05	1 0.187	0.003 (.290 0.07	57 0.0012	1112	31	1101	18	1106	15	99.4
FC-1 - 12	1.71E+08	473.7	292.6	1.62	-2	13	-650.0	82600	99.94	1 FALSE	1.948 0.05	2 0.186	0.003 (.312 0.07	59 0.0012	1001	32	1097	18	1102	15	01.0
FC-1 - 14	1.70E+08	101.69	38.29	2.66	1	12	1200.0	17300	99.84	1 FALSE	1.936 0.06	5 0.185	0.003 (.221 0.07	54 0.0020	1079	53	1090	23	1095	17	01.5
FC-1 - 15	1.70E+08	273.1	151.4	1.80		, 11	157.1	6613	99.88	1 FALSE	1.958 0.05	3 0.185	0.003 (.331 0.07	55 0.0013	1108	34	1101	19	1094	15	98.7
FC-1 - 16	1.64E+08	288.7	164.1	1.76	ή	13	-433.3	49110	66.66	1 FALSE	1.953 0.05	7 0.187	0.003 (.248 0.07	53 0.0015	1077	40	1100	19	1105	16	02.7
FC-1 - 17	1.75E+08	181.5	77.98	2.33	6	12	133.3	3429	99.85	1 FALSE	1.925 0.05	9 0.184	0.003 (.265 0.07	57 0.0016	1087	42	1088	20	1088	16	00.1
FC-1 - 18	1.67E+08	321.5	187.6	1.71	16	13	81.3	3376	99.88	1 FALSE	1.965 0.05	2 0.186	0.003 (.134 0.07	57 0.0013	1113	34	1103	18	1102	15	98.9
FC-1 - 19	1.65E+08	283	128.8	2.20	-2	11	-550.0	47040	99.92	1 FALSE	1.946 0.05	5 0.186	0.003 (.378 0.07	50 0.0014	1095	37	1095	19	1101	15	00.5
Plesovice - 1	1.87E+08	702.1	65.47	10.72	9	12	200.0	6380	99.86	1 FALSE	0.394 0.01	2 0.053	0.001 (.218 0.05	38 0.0012	363	50	337	6	333	S	98.7
Plesovice - 2	1.87E+08	406.1	32.58	12.46	16	11	68.8	1358	99.87	1 FALSE	0.380 0.01	3 0.052	0.001 (.486 0.05	27 0.0014	316	60	326	10	328	ŝ	00.6
Plesovice - 3	1.84E+08	466.2	40.14	11.61	4	t 14	350.0	6035	99.76	1 FALSE	0.381 0.01	3 0.052	0.001 (.111 0.05	37 0.0014	1 358	59	328	10	325	S.	98.9
Plesovice - 4	1.86E+08	395.7	35.1	11.27	1	Ξ	1100.0	20550	99.83	1 FALSE	0.379 0.01	3 0.052	0.001 (0.041 0.05	29 0.0015	325	64	325	10	326	ŝ	00.2
Plesovice - 5	1.86E+08	664.6	60.13	11.05	œ	12	150.0	4385	99.77	1 FALSE	0.398 0.01	2 0.053	0.001 (.266 0.05	42 0.0012	379	50	340	6	335	5	98.5
Plesovice - 6	1.81E+08	813	69.8	11.65	ŝ	12	400.0	13747	99.84	1 FALSE	0.389 0.01	1 0.053	0.001 (.286 0.05	29 0.0010	322	43	333	8	334	ŝ	00.2
Plesovice - 7	1.83E+08	547.8	47.84	11.45	6-	II	-122.2	28220	99.83	1 FALSE	0.396 0.01	2 0.053	0.001 (.286 0.05	33 0.001	342	47	338	6	334	S	98.6
Plesovice - 8	1.81E+08	685.6	74.99	9.14	11	II	100.0	3171	99.76	1 FALSE	0.392 0.01	2 0.053	0.001 (0.043 0.05	37 0.0012	358	50	335	6	330	Ŋ	98.4
Plesovice - 9	1.80E+08	629.2	56.86	11.07	1	12	1200.0	32270	99.80	1 FALSE	0.392 0.01	2 0.053	0.001 (0.011 0.05	40 0.0013	371	54	336	6	333	Ŋ	99.1
Plesovice - 10	1.79E+08	538.7	47.07	11.44	ή	10	-366.7	27390	99.81	1 FALSE	0.384 0.01	3 0.052	0.001 (.193 0.05	36 0.0014	l 354	59	329	10	329	ŝ	00.1
Plesovice - 11	1.79E+08	640.3	58.45	10.95	6	11	122.2	3600	99.87	1 FALSE	0.383 0.01	2 0.052	0.001 (0.057 0.05	36 0.0013	354	55	329	6	329	ŝ	0.00
Plesovice - 12	1.75E+08	598.8	53.39	11.22	-21	11	-52.4	29650	99.79	1 FALSE	0.385 0.01	1 0.052	0.001 (.294 0.05	10.001	373	46	330	8	326	Ŋ	98.8
Plesovice - 13	1.77E+08	829.5	84.84	9.78	ή	11	-366.7	41960	99.84	1 FALSE	0.388 0.01	1 0.053	0.001 (.113 0.05	33 0.001	339	47	332	8	331	ŝ	99.4
Plesovice - 14	1.84E+08	678.9	62.81	10.81	1	13	1300.0	34640	99.80	1 FALSE	0.394 0.01	3 0.053	0.001 (.122 0.05	35 0.0013	350	55	337	6	333	5	0.66
Plesovice - 15	1.70E+08	563.9	49.19	11.46	10	12	120.0	2781	99.89	1 FALSE	0.388 0.01	2 0.054	0.001 (.184 0.05	22 0.0012	294	52	332	6	336	ŝ	01.2
Plesovice - 16	1.77E+08	863.1	114.8	7.52	-	. 13	-1300.0	43570	99.84	1 FALSE	0.396 0.01	1 0.053	0.001 (.162 0.05	36 0.0010	352	42	339	8	335	Ŋ	98.8
Plesovice - 17	1.88E+08	846.2	90.5	9.35	0	14	700.0	21775	99.93	1 FALSE	0.387 0.01	3 0.053	0.001 (.287 0.05	23 0.0013	299	57	332	6	335	ŝ	01.1
Plesovice - 18	1.89E+08	962.4	109.37	8.80	9	15	250.0	8222	99.90	1 FALSE	0.394 0.01	2 0.054	0.001 (.289 0.05	36 0.0012	354	51	337	6	336	ŝ	9.66
Plesovice - 19	1.83E+08	725.5	69.89	10.38	11	. 15	136.4	3352 1	00.00	1 FALSE	0.384 0.01	2 0.053	0.001 0	.287 0.05	25 0.0013	307	56	329	6	334	ŝ	01.5

Corrections: 1 =	threshold ²⁰⁴	¹ Pb no co	rrection ((100 cp:	s); 2 = tl	hreshold	d % ²⁰⁴ Pb−	based correc	tion(20)% error) ; ;	3 = threshold	% for ²⁰⁸	Pb-base	ed corr	ection (9	8.5 %Pb	*).						
									F			Is	otopicı	ratios				Ca	lculated	d ages		_	
Sample	90 Zr	U (Th	Th/U	^{204}Pb	$\pm 2\sigma$	% 土 204~1	$^{206}\text{Pb}/^{204}$	%	Corrections	$\frac{^{207}\text{Pb/}}{^{235}} \pm 2$	λσ ²⁰⁶ Pb	/ ± 2σ	EC ³	²⁰⁷ Pb/	± 2σ	²⁰⁷ Pb/	+ 2	²⁰⁷ Pb/	7 ²⁰	⁶ Pb/	+	%
17-1u	cbs	(mdd)	(Inidd)		cps	9.4	The cps	Pb cps	r*d4		Ω_{cc7}	Ω_{oc7}			qdonz		qdanz	07	Ω_{cc7}	07	, U°	3	olic
FC-1 - 1	sampies 2.34E+08	236.3	86.18	2.74	15	13	86.7	3925	6 <i>2</i> .66	1 FALSE	1.959 0.0	31 0.180	§ 0.002	0.438	0.0768	0.0009	1111	24	1100	11	1101	12 1	0.00
FC-1 - 2	2.19E+08	289.9	174.8	1.66	8	12	150.0	8615	99.77	1 FALSE	1.947 0.0	31 0.186	5 0.002	0.353	0.0760	0.0010	1088	26	1096	11	1098	12 1	0.00
FC-1 - 3	2.14E+08	167.8	76.28	2.20	2	11	550.0	19850	99.70	1 FALSE	1.947 0.0	36 0.180	5 0.002	0.174	0.0756	0.0013	1075	34	1096	12	1100	12 1	0.00
FC-1 - 4	2.09E+08	118.82	39.87	2.98	-4	12	-300.0	27460	99.71	1 FALSE	1.945 0.0	38 0.18	5 0.002	0.100	0.0757	0.0015	1080	38	1095	13	1096	13 1	0.00
FC-1 - 5	2.12E+08	238.9	111.62	2.14	-	11	-1100.0	55880	99.80	1 FALSE	1.948 0.0	29 0.180	5 0.002	0.119	0.0763	0.0010	1098	25	1097	10	1097	12 1	0.00
FC-1 - 6	2.08E+08	373.1	225	1.66	2	6	460.0	43165	99.82	1 FALSE	1.962 0.0	26 0.180	5 0.002	0.436	0.0767	0.0007	1115	19	1102	6	1099	12	98.2
FC-1 - 7	2.06E+08	405.1	265.3	1.53	10	10	100.0	9325	99.87	1 FALSE	1.949 0.0	25 0.18.	7 0.002	0.253	0.0757	0.0008	1083	20	1098	6	1103	11	0.00
FC-1 - 8	2.09E+08	270.1	147.6	1.83	16	12	75.0	3926	99.79	1 FALSE	1.951 0.0	29 0.180	5 0.002	0.203	0.0762	0.000	1094	25	1099	10	1098	12	99.3
FC-1 - 9	2.06E+08	252.2	136.4	1.85	-	10	-1000.0	58040	99.75	1 FALSE	1.954 0.0	30 0.180	5 0.002	0.123	0.0760	0.0010	1095	24	1099	10	1097	12	99.3
FC-1 - 10	2.04E+08	181.6	73.3	2.48	12	6	78.3	3448	99.72	1 FALSE	1.945 0.0	36 0.180	5 0.002	0.244	0.0756	0.0012	1074	34	1095	12	1099	13 1	0.00
FC-1 - 12	2.01E+08	160.7	72	2.23	-7	6	-132.4	35790	69.66	1 FALSE	1.956 0.0	34 0.180	5 0.002	0.362	0.0763	0.0012	1100	30	1102	12	1097	12	6.66
FC-1 - 13	2.02E+08	324.3	193.2	1.68	8	11	137.5	9155	99.81	1 FALSE	1.952 0.0	27 0.18	7 0.002	0.288	0.0755	0.0009	1076	24	1098	6	1102	12 1	0.00
FC-1 - 14	2.02E+08	248.4	154.21	1.61	14	13	92.9	3977	99.71	1 FALSE	1.944 0.0	32 0.18	5 0.002	0.107	0.0762	0.0011	1092	30	1095	11	1094	12	99.3
FC-1 - 15	1.99E+08	123.71	51.7	2.39	4	11	157.1	3923	79.67	1 FALSE	1.963 0.0	36 0.187	7 0.002	0.205	0.0765	0.0013	1100	33	1101	12	1102	13 1	0.00
FC-1 - 17	1.99E+08	339.7	213.1	1.59	-2	10	-500.0	75500	99.82	1 FALSE	1.958 0.0	28 0.180	5 0.002	0.384	0.0764	0.0009	1101	22	1100	10	1102	12 1	0.00
FC-1 - 18	2.02E+08	162.8	90.1	1.81	6	11	122.2	4013	69.66	1 FALSE	1.927 0.0	33 0.18-	± 0.002	0.103	0.0762	0.0012	1095	31	1089	12	1601	12 1	0.00
FC-1 - 19	2.01E+08	255.6	153.8	1.66	-10	10	-100.0	57110	99.76	1 FALSE	1.963 0.0	31 0.18	7 0.002	0.20	0.0765	0.0011	1101	27	1102	11	1103	12 1	0.00
FC-1 - 20	2.23E+08	112.8	43.68	2.58	÷.	14	-466.7	26030	67.67	1 FALSE	1.953 0.0	53 0.18	5 0.003	0.149	0.0763	0.0019	1102	53	1098	18	1098	14	9.66
Plesovice1 - 1	2.19E+08	694.4	64.95	10.69	7	13	185.7	6663	16.66	1 FALSE	0.387 0.0	07 0.05-	100.01	0.175	0.0525	0.0009	315	39	332	ŝ	336	4	01.4
Plesovice1 - 2	2.03E+08	702.5	65.23	10.77	۲-	Π	-157.1	45880	99.84	1 FALSE	0.394 0.0	07 0.05-	100.01	0.211	0.0524	0.0008	297	35	337	Ŋ	340	4	01.0
Plesovice1 - 3	2.01E+08	480.4	38.94	12.34	6	11	122.2	3452	99.87	1 FALSE	0.396 0.0	09 0.05	100.01	0.33(0.0531	0.0010	321	43	338	9	340	4	9.00
Plesovice1 - 4	2.01E+08	751.2	75.86	9.90	-8	10	-125.0	48580	69.66	1 FALSE	0.403 0.0	07 0.05,	100.01	0.595	0.0545	0.0008	384	33	343	S	336	4	98.0
Plesovice1 - 5	1.97E+08	495.1	42.89	11.54	-	11	1100.0	31750	99.72	1 FALSE	0.405 0.0	08 0.05	± 0.001	0.207	0.0540	0.0010	364	43	345	9	341	4	98.7
Plesovice1 - 6	1.95E+08	767.3	75.02	10.23	<u>ئ</u>	12	-400.0	48120	99.79	1 FALSE	0.398 0.0	07 0.05	± 0.001	0.221	0.0532	0.0009	329	37	340	Ŋ	339	4	7.66
Plesovice1 - 7	1.94E+08	811.4	65.93	12.31	-	10	1000.0	50840	99.81	1 FALSE	0.396 0.0	07 0.05	± 0.001	0.274	0.0531	0.0008	323	35	339	Ŋ	338	4	6.66
Plesovice1 - 8	1.92E+08	776.4	78.12	9.94	1	11	1100.0	47770	99.80	1 FALSE	0.393 0.0	08 0.05	± 0.001	0.225	0.0534	0.0010	330	43	336	9	337	4	00.3
Plesovice1 - 9	1.91E+08	1254.1	114.67	10.94	-13	10	-76.9	76540	99.80	1 FALSE	0.397 0.0	06 0.05.	3 0.001	0.408	0.0542	0.0007	371	28	340	ŝ	335	4	98.7
Plesovice1 - 10	1.93E+08	527	47.1	11.19	4	11	275.0	8243	99.64	1 FALSE	0.409 0.0	10 0.05	£ 0.001	0.069	0.0549	0.0012	398	51	349	~	341	4	97.7
Plesovice2 - 1	2.25E+08	773.7	64.58	11.98	9	13	216.7	9037	06.66	1 FALSE	0.394 0.0	07 0.05	100.01	0.045	0.0529	0.000	313	38	337	S	341	4	01.3
Plesovice2 - 2	2.13E+08	932.8	103.59	9.00	4	10	250.0	15828	98.66	1 FALSE	0.398 0.0	06 0.05:	5 0.001	0.077	0.0529	0.0007	318	30	340	Ŋ	342	4	9.00
Plesovice2 - 3	2.04E+08	602.9	52.83	11.41	9	12	200.0	6623	99.82	1 FALSE	0.398 0.0	09 0.05:	5 0.001	0.027	0.0525	0.0011	310	46	340	9	342	4	00.7
Plesovice2 - 4	2.02E+08	730.5	67.7	10.79	-	10	-1000.0	48000	99.83	1 FALSE	0.400 0.0	08 0.05:	5 0.001	0.256	0.0534	0.0009	337	38	341	9	342	4	00.3
Plesovice2 - 5	2.04E+08	394.1	30.72	12.83	2	11	550.0	12805	99.75	1 FALSE	0.392 0.0	08 0.05	3 0.001	0.025	0.0536	0.0011	339	45	336	9	333	4	99.1
Plesovice2 - 6	2.00E+08	643.1	58.64	10.97	12	13	108.3	3482	99.72	1 FALSE	0.403 0.0	08 0.05	± 0.001	0.168	0.0538	0.0010	347	43	343	9	339	4	98.9
Plesovice2 - 7	1.96E+08	899.1	94.83	9.48	9-	11	-183.3	56260	99.76	1 FALSE	0.397 0.0	07 0.05	3 0.001	0.336	0.0537	0.008	348	33	339	Ŋ	335	4	98.8
Plesovice2 - 8	1.93E+08	446.3	36.21	12.33	Ŋ	6	200.0	5851	99.87	1 FALSE	0.390 0.0	09 0.05.	3 0.001	0.171	0.0525	0.0010	317	44	334	9	335	4	00.2
Plesovice2 - 9	1.97E+08	677.9	71.72	9.45	ŝ	10	333.3	14340	99.80	1 FALSE	0.395 0.0	08 0.05	± 0.001	0.245	0.0538	0.0009	352	38	338	9	336	4	99.5
NOTES: ¹ after	Hg correction	1; ² radiog	enic Pb; ^ŝ	³ error α	orrectio	n. Abbr	eviations:	cps = count:	s per se	cond.													

APPENDIX B

Location and sample information, organic-walled microfossils and trace fossils. Sample coordinates are in Grid Zone 21T (WGS84). Locations are indicated also on Figure 2 and shown in a stratigraphic column (Fig. B1).



Figure B1. Stratigraphic interpretation of the southern limb of the southernmost syncline on the eastern shore of Scatarie Island (Fig. 2). Members 3 and 4 were logged in August 2017. Members 1 and 2 are schematic and based on notes made in 1991 when they were partially exposed. Abbreviation: NW Cove Fm = Northwest Cove Formation.

Scatarie Island, Bengal Road Fm. Coastal section of upper part of Bengal Road Formation

SC17-5. 0289970, 5099472. Abundant dispersed organic matter and very scarce organic-walled microfossils that are limited to fragments without recognizable forms, except for a small filament of possible cyanobacteria (Fig. 6a).

SC17-6. 0289994, 5099484. Small amount of dispersed organic matter and very scarce organic-walled microfossils without recognizable forms, except for an acritarch assigned to *Polygonium* sp. (Fig. 6b).

SC17-7. 0289995, 5099517. Small amount of dispersed organic matter. Possible black carbonaceous fragments without recognizable forms.

Hay Island, MacCodrum? Formation

1. Section on southern Hay Island at 0291878, 5099931. Teichichnid trace fossils (Fig. 7). Sample SC17-2, with small amount of dispersed organic matter, possible black carbonaceous fragments without recognizable morphologies. From same area at 0291885, 5099911, sample SC17-3, with small amount of dispersed organic matter, possible black carbonaceous fragments without recognizable morphologies. Among the maceration-resistant minerals are some idiomorphic zircon crystals. Also sample SC17-4, with no dispersed organic matter. Possible black carbonaceous fragments without recognizable shapes.

2. Section on northeastern Hay Island at 0292071, 5100170, cleaved dark grey siltstone with carbonate nodules. Sample SC17-8, with small amounts of organic matter. Possible black carbonaceous fragments without recognizable forms.

3. Section on northern Hay Island at 0292230, 5100309. *Gyrolithes, Teichichnus* and other trace fossils (Fig. 8). Sample SC17-9, with no dispersed organic material. Possible black carbon fragments without recognizable forms.