

Lithogeochemistry, petrology, and the acid-generating potential of the Goldenville and Halifax groups and associated granitoid rocks in metropolitan Halifax Regional Municipality, Nova Scotia, Canada

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[Aller au sommaire du numéro](#)

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Résumé de l'article

Une cartographie géologique détaillée des groupes de Goldenville et de Halifax dans la municipalité régionale de Halifax, en Nouvelle-Écosse, a donné lieu à deux divisions de ces deux groupes. Pour ce qui est du groupe de Goldenville, les deux formations suivantes ont été établies: la Formation de Taylors Head, composée principalement de méta-grès (wacke feldspathique) et la Formation de Beaverbank sus-jacente, composée principalement de métasiltite riche en Mn. En ce qui concerne le groupe de Halifax, les deux formations que voici ont été identifiées: la Formation de Cunard, composée principalement d'ardoise riche en sulfures, et la Formation de Bluestone sus-jacente, composée surtout de métasiltite et de méta-grès. Les études lithogéochimiques et pétrologiques ont permis d'établir que les Formations de Beaverbank, de Cunard et de Bluestone étaient susceptibles d'agir comme précurseurs acides, la pyrrhotite étant le principal minéral à sulfure de fer présent. La présence d'un drainage rocheux acide (DRA) est régie par la composition chimique de la roche brute et l'association minérale. Pour vérifier la possibilité d'un DRA, de concert avec des analyses traditionnelles à fluorescence X et de bilan acide-base, un appareil portatif à fluorescence X a été utilisé pour recueillir de manière relativement économique des données sur la roche brute et les sulfures. Les résultats obtenus par les trois méthodes se comparent bien. L'utilisation d'un appareil à fluorescence X portatif permet de prédire de manière économique et rapidement la possibilité du DRA, mais cette technique ne saurait se substituer à l'analyse du bilan acide-base prévue par la réglementation.

[Traduit par la rédaction]

The Appledore Island pluton of the Rye Complex, coastal New Hampshire and Maine, USA: geochronological and chemical evidence for the affinity of an enigmatic terrane

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ABSTRACT

The Rye Complex of coastal New Hampshire and Maine is a peri-Gondwanan terrane that up to now had an uncertain origin. An offshore portion of the complex, Appledore Island of the Isles of Shoals, hosts a mainly dioritic intrusion that yielded an U-Pb zircon age of 361.09 ± 0.14 Ma, allowing comparison of its geochemical characteristics with mafic rocks of similar age across the northern Appalachian orogen. The Appledore Island diorite has similar major, trace, and isotopic compositions as continental rift tholeiite in the Narragansett Basin in southern New England and in the Maritimes Basin of Canada. These intraplate volcanic rocks range from 375 to 330 Ma, bracketing the age of the Appledore Island diorite. Their intraplate tectonic setting reflects regional extension during the Late Devonian to Early Carboniferous which produced successor basins after the Acadian orogeny. The geochemical and age similarities of the Appledore Island diorite and the mafic rocks of the successor basins suggest that the Rye Complex is a basement fragment of a successor basin block. Further evidence of the identity of the Rye Complex is provided by the isotopic composition of intermingled, comagmatic granitic rocks associated with the Appledore Island diorite. The granite has a Ganderian isotopic signature, suggesting that the Rye Complex is a Ganderian basement block that was transposed by movement along the Norumbega Fault System to its position adjacent to the Merrimack Trough of New Hampshire and Maine.

RÉSUMÉ

Le complexe Rye sur les côtes du New Hampshire et du Maine est un terrane périgondwanien dont l'origine était incertaine jusqu'à ce jour. On a découvert dans une partie du complexe en mer, l'île Appledore des îles de Shoals, une intrusion composée principalement de diorite qui a livré un âge par datation U-Pb sur zircon de $361,09 \pm 0,14$ Ma, permettant de comparer ses caractéristiques géochimiques à celles de roches mafiques du même âge dans l'orogène des Appalaches septentrionales. La diorite de l'île Appledore présente des compositions d'éléments majeurs, traces et isotopiques comparables à celles de la tholéiite du rift continental dans le bassin de Narragansett au sud de la Nouvelle-Angleterre et dans le bassin des Maritimes du Canada. L'âge de la diorite de l'île Appledore se situe dans la fourchette d'âges de ces roches volcaniques intraplaques (de 375 à 330 Ma). Leur cadre tectonique intraplaque correspond à l'extension régionale au cours de la période allant du Dévonien tardif au Carbonifère précoce, qui a produit des bassins successeurs après l'orogénèse acadienne. Les similitudes sur le plan de la géochimie et de l'âge de la diorite de l'île Appledore par rapport aux roches mafiques des bassins successeurs semblent indiquer que le complexe Rye constitue un fragment du socle continental d'un bassin successeur. La composition isotopique des roches granitiques comagmatiques entremêlées associées à la diorite de l'île Appledore fournit une autre preuve de la nature du complexe Rye. Le granit a une signature isotopique ganderienne, ce qui tend à indiquer que le complexe Rye est un socle continental ganderien qui a été transposé par le mouvement le long du système de failles de Norumbega à sa position adjacente à la Fosse de Merrimack dans le New Hampshire et le Maine.

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INTRODUCTION

Problems inherent in the interpretation of orogenic belts include understanding the precise timing and nature of their formation and what processes were involved. Many mountain belts do not result from simple, easily datable crustal plate collisions, but rather from a series of collisions of smaller microplates, volcanic arcs, and seamounts along the leading edge of continental plates. The Appalachian orogen, for example, consists of numerous accreted terranes, each with its own individual tectonic history, that were added to the Laurentian margin over time (Fig. 1 inset; Fig. 2). Temporally, these terranes preserve a record from the Grenville orogeny through a series of late Paleozoic orogenic events, including the Taconic (~550–440 Ma), Salinic (~450–420 Ma), Acadian (~416–356 Ma), Neocadian (~380–340 Ma), and Alleghanian (~325–280 Ma), that ended with the joining of Gondwana and Laurentia during the Alleghanian orogeny at ~280 Ma (van Staal *et al.* 2009, 2012).

In recent years considerable progress has been made in Canada in characterizing and subdividing broad lithotectonic zones into distinct terranes based on geochemical and other characteristics (e.g., Barr *et al.* 1994, 1998; van Staal *et al.* 1996, 1998, 2009, 2012; van Staal and Barr 2012; van Staal and Zagorevski 2012; Nance *et al.* 2008). Although several sub-terrane exist, the two larger terranes or domains that are interpreted to extend southwest through New England are Ganderia and Avalonia (Hibbard *et al.* 2006; Thompson *et al.* 1996, 2012; Aleinikoff *et al.* 2007; Walsh *et al.* 2007). Ganderia is a Late Neoproterozoic to Early Cambrian arc and subsequent passive margin terrane that rifted from Gondwana at about 505 Ma (e.g., van Staal and Barr 2012). Avalonia, an assembly of several fault-bounded, largely juvenile arc-related belts, may have rifted from Gondwana as much as 30 m.y. after Ganderia, and the two microcontinents had separate histories in the Early Paleozoic (van Staal *et al.* 2009, 2012). Docking of Ganderia with composite Laurentia is interpreted to have caused the Salinic orogeny, followed by Avalonia which caused the Acadian orogeny (van Staal *et al.* 2009, 2012).

Most researchers agree that Neoproterozoic rocks of the Boston and Narragansett Bay areas correlate with Avalonia of Maritime Canada (Socci *et al.* 1990; Skehan and Rast 1995). The Nashoba terrane and the Merrimack Trough, lying adjacent to Avalonia on the west, are now thought to have Ganderian affinities (Wintsch *et al.* 2007; Kuiper *et al.* 2013). However, some parts of New England remain poorly understood and their correlation with other Appalachian terranes is not yet determined. One of the more enigmatic of these New England terranes is the Rye Complex of coastal New Hampshire and Maine.

The Rye Complex of probable Ordovician or older age (Lyons *et al.* 1997; Hussey *et al.* 2010; Kane *et al.* 2014) lies outboard of the Merrimack Trough and is considered to be a peri-Gondwanan terrane (Fig. 1). Attempts to corre-

late the Rye Complex with other peri-Gondwanan terranes to the northeast along coastal Maine, however, have been tentative. Possible correlatives include the Cape Elizabeth and Passagassawaukeag formations and/or the Falmouth-Brunswick sequence, based on similar lithologies. Pelite, feldspathic wacke, calc-silicate rock, minor marble, and amphibolite are common to the Rye Complex and their potential correlatives (Bothner and Hussey 1999; Hussey *et al.* 2010). Uncertain age determinations and limited P-T and isotopic data from the Rye Complex, however, have hindered confirmation of these suggestions.

In this contribution, we present geochemistry and geochronology for the Appledore Island diorite that intruded variously deformed metasedimentary and metaigneous units of the Rye Complex. We use these data to compare these rocks to igneous rocks of similar age in coastal New England and Maritime Canada. We also present isotopic data for both the diorite and comagmatic granite that occurs along the margin of the dioritic pluton. The granite is interpreted to represent a partial melt of the host rocks, and as such, may represent the chemical signature of those rocks.

GEOLOGIC SETTING

The Rye Complex is located east of the Merrimack trough in the Rye anticline of southeastern New Hampshire. Metasedimentary and metaigneous rocks of the complex are variably mylonitized and cut by a number of complex brittle dextral strike-slip faults (Swanson 1995). The on-shore part of the complex occupies a relatively small area, spanning a length of ~25 km and only 7 km at its widest point (Lyons *et al.* 1997; Hussey *et al.* in press). It also crops out in numerous submerged ledges (Brooks 1986, 1990) and on the 9 small islands that compose the Isles of Shoals located ~16 km off the coast of New Hampshire and Maine, increasing its regional extent by well over 400%.

Katz (1917) referred to this crystalline area as the Algonkian(?) complex, and Wandke (1922) proposed the name Rye Gneiss. Later, Billings (1956) and Novotny (1963, 1969) used the name Rye Formation and subdivided the unit into an older metasedimentary and younger metavolcanic members. Hussey (1980) showed that the “metavolcanic” member consists of a series of felsic intrusive rocks that cross-cut the host at very shallow angles. He also documented the widespread mylonitization of the complex, which has several internal ultramylonite zones, the largest of which is named the Great Commons fault zone in New Hampshire (Carrigan 1984; Swanson 1988). The Portsmouth fault zone bounds the complex on the west and north, everywhere separating it from the adjacent Kittery Formation of the Merrimack belt. Both fault zones are right-lateral ductile shear zones within the Norumbega fault system (Bothner

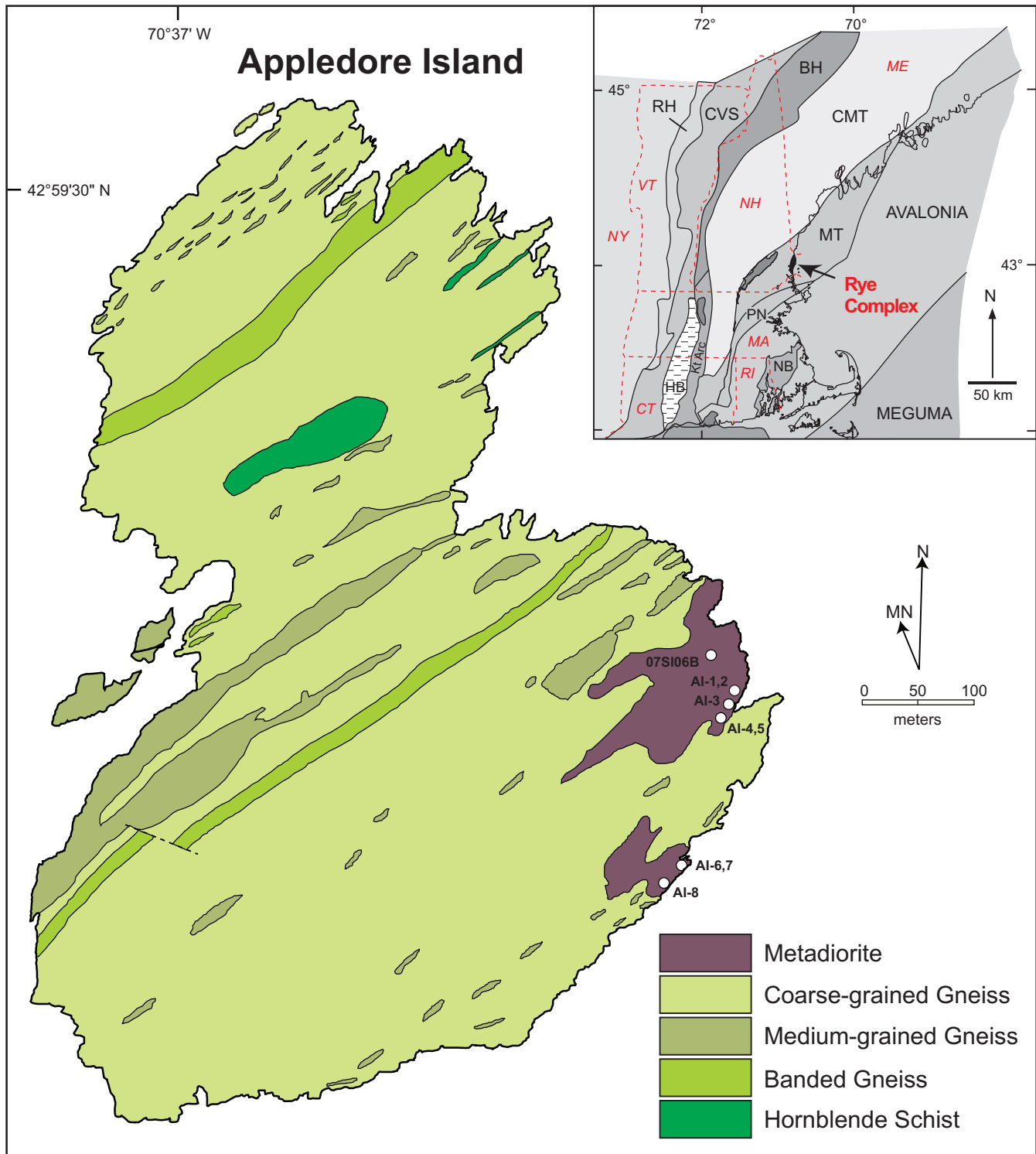


Figure 1. Geologic map of Appledore Island of the Isles of Shoals, Maine (after Blomshield 1975) located approximately at the tip of the Rye Complex arrow (inset). The Appledore Island diorite is a small intrusion located along the southeastern portion of the island. Sample locations shown by white dots. Inset is a simplified geologic map of New England and northeastern Canada showing the terranes and the location of the Rye Complex. Offshore terrane boundaries are broadly defined geophysically (Brooks 1990; Bothner and Hussey 1999). RH = Rowe Hawley, CVS = Connecticut Valley synclinorium, BH = Bronson Hill, CMT = Central Maine trough, MT = Merrimack trough, PN = Putnum-Nashoba, HB = Hartford Basin, NB = Narragansett Basin. Merrimack Trough and Putnam-Nashoba are thought to be Ganderian.

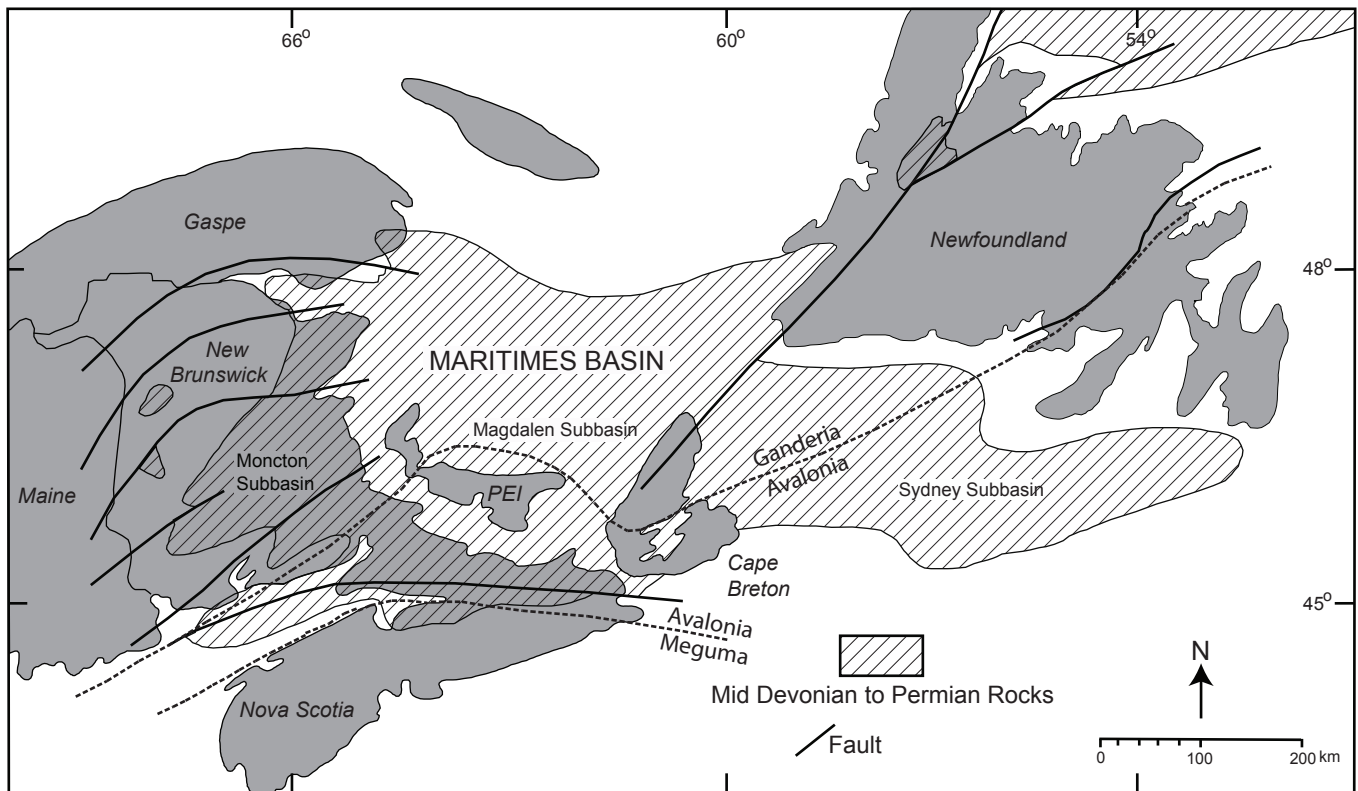


Figure 2. Generalized map showing the distribution of Paleozoic sedimentary rocks in Atlantic Canada and the locations of post-Acadian successor basins (after Fyffe and Barr 1986). Boundaries between Meguma and Avalonia and Avalonia and Ganderia are from Hibbard *et al.* (2006).

and Hussey 1999; Swanson 1999). Felsic plutonic bodies and abundant associated pegmatite, both variably sheared into near-parallelism with highly mylonitized host rocks, led Lyons *et al.* (1997) to coin the name Rye Complex. The host complex contains highly sheared quartzofeldspathic gneiss, pelitic schist, and amphibolite (Hussey 1980; Carrigan, 1984; Swanson 2007), all metamorphosed to amphibolite facies. Interlayered metapelite contains garnet, andalusite, and sillimanite, and calcareous rocks contain diopside (Hussey 1980; Carrigan 1984). Rocks within the Great Commons fault system show the most intense mylonitization (Novotny 1969; Carrigan 1984; Swanson 2007).

Because mylonitization is commonly accompanied by increased fluid flow, alteration was expected in both host and intrusive rocks. For this reason, we looked for outcrops as far as possible from the zones of intense mylonitization to minimize the effects of open system behavior and to obtain the freshest samples possible. The Isles of Shoals, located approximately 20 km from the Great Commons fault zone, host a relatively low-grade diorite on Appledore Island. The diorite seems to be the best candidate in the area to provide an original igneous chemical composition.

GEOLOGY AND PETROGRAPHY OF THE APPLIEDORE ISLAND DIORITE

The bedrock geology of the Isles of Shoals was first mapped by Fowler-Billings (1959) and later by Blomshield (1975). Appledore Island contains the most varied lithologies of the isles including the only diorite. Fowler-Billings (1959) associated the diorite with the larger Exeter Diorite on the New Hampshire mainland. The Appledore Island diorite occupies approximately 0.2 km² on the eastern shore of Appledore Island in two mappable bodies (Fig. 1). Both are amoeboid in shape with tentacle-like offshoots that exploit the dominant northeast-trending foliation in the host coarse-grained quartzofeldspathic gneiss. The diorite is medium-grained, weakly foliated, and intermingled with grey muscovite-biotite granite (Fig. 3). These rocks are cut by Mesozoic diabase dykes.

The Appledore Island diorite was termed metadiorite by Bothner *et al.* (2009) because of its weak foliation. It contains pleochroic green hornblende which is generally unzoned, although it shows some opaque phase exsolution along cleavage planes and titanite is commonly developed at grain boundaries with biotite. Some amphibole grains show

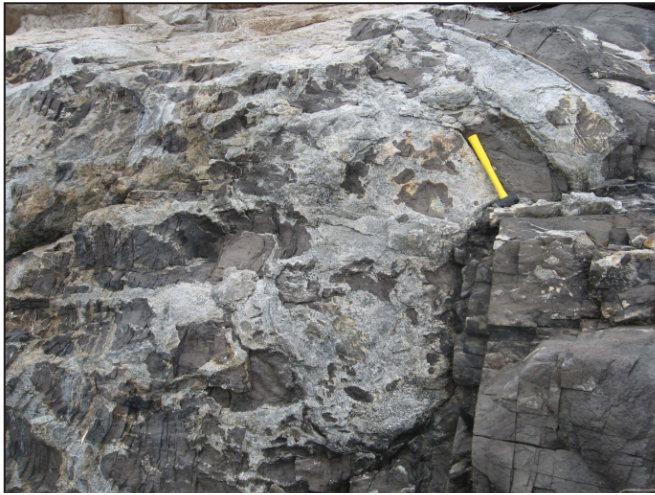


Figure 3. Photograph showing mingling of felsic and mafic magmas at the western margin of the Appledore Island pluton. The granitic rocks are of minor abundance along the contact between the pluton and its country rocks.

weak zoning from faint brown cores to green mantles and rims (Fig. 4). Weaker pleochroism is characteristic of small grains and in some places the rims of larger grains. Plagioclase is typically quite fresh and preserves strong zoning, a feature that would have been eradicated under intense metamorphism and deformation like that experienced by plutons in the Rye Complex onshore. Minor amounts of sericite are present in some of the more calcic cores.

Abundant enclaves of the diorite are present along the southeastern granitic margin of the southern lobe of the pluton. The rounded nature and lack of evidence of brittle morphologies suggests that the two magma types were contemporaneous (Fig. 3). The granite that is mingled with the diorite contains muscovite and biotite of similar size. In places, they are intergrown, suggesting that the muscovite is magmatic rather than secondary (Miller *et al.* 1981).

ANALYTICAL METHODS

Amphibole analyses were obtained using a Cameca SX50 electron microprobe at Brigham Young University (BYU). Bulk-rock major and trace element XRF analyses were done with a Siemens SRS 303 at BYU using fused disks for major elements and pressed powder pellets for trace elements. Additional trace elements were determined by inductively coupled plasma mass spectrometry (ICP-MS) at ALS Chemex, Reno, Nevada.

Isotopic compositions of Nd, Sr and Pb were obtained at Memorial University of Newfoundland, using a multicollector Finnigan Mat 262 mass spectrometer in static mode. Nd and Sr isotopic ratios were normalized to $^{146}\text{Nd}/^{144}\text{Nd} = 0.721903$ and $^{88}\text{Sr}/^{86}\text{Sr} = 8.375209$, respectively. The reported $^{143}\text{Nd}/^{144}\text{Nd}$ and $^{87}\text{Sr}/^{86}\text{Sr}$ ratios were corrected for the deviation from repeated duplicates of the JNdi-1 ($^{143}\text{Nd}/^{144}\text{Nd}$

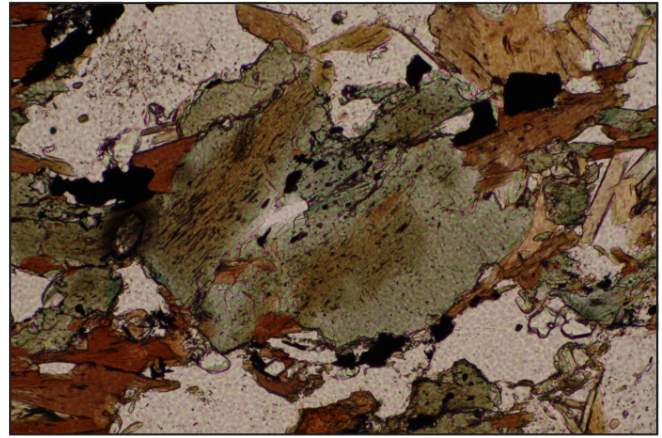


Figure 4. Photomicrograph of zoned amphibole in the Appledore Island diorite in plane polarized light. Horizontal field of view is 2 mm.

$= 0.512115$; Tanaka *et al.* 2000) and NBS 987 ($^{87}\text{Sr}/^{86}\text{Sr} = 0.710340$) standards. During the course of data acquisition replicates of the standards give a mean value of $^{143}\text{Nd}/^{144}\text{Nd} = 0.512137 \pm 19$ (Std dev, $n=112$) for JNdi-1 and $^{87}\text{Sr}/^{86}\text{Sr} = 0.710291 \pm 25$ (Std dev, $n=46$) for NBS 987. Pb isotopic ratios are reported corrected for mass fractionation. A correction factor of 0.107% per amu was obtained by measuring the deviation from repeated ($n=12$) analyses of the NBS 981 standard. In-run precisions on all isotopic ratios are given at 95% confidence level. Errors on Nd and Sr isotopic compositions are $<0.002\%$ and errors on the $^{147}\text{Sm}/^{144}\text{Nd}$ ratio are estimated to be less than 0.1%.

The ϵ_{Nd} values are calculated using $^{147}\text{Sm}/^{144}\text{Nd} = 0.1967$ and $^{143}\text{Nd}/^{144}\text{Nd} = 0.512638$ values for the present day chondrite uniform reservoir (CHUR). ^{147}Sm decay constant is $6.54 \cdot 10^{-12} \text{ y}^{-1}$ (Steiger and Jäger 1977). T_{DM} is calculated using a linear evolution for a mantle separated from the CHUR at 4.55Ga and having a present day ϵ_{Nd} value of +10 (Goldstein *et al.* 1984). Following Stern (2002), samples with $^{147}\text{Sm}/^{144}\text{Nd} > 0.165$ yield unreliable model ages: two of our samples with these high $^{147}\text{Sm}/^{144}\text{Nd}$ are not included in the discussion here.

Zircon grains were separated from bulk rock samples by standard crushing, heavy liquid, and magnetic separation techniques, and subsequently handpicked under the binocular microscope based on clarity and crystal morphology. To minimize the effects of Pb loss, the grains were subjected to a version of the thermal annealing and acid leaching (also known as chemical abrasion or CA-TIMS) technique of Mattinson (2005) prior to isotope dilution thermal ionization mass-spectrometry (ID-TIMS) analyses using a mixed ^{205}Pb - ^{233}U - ^{235}U tracer solution (spike). Details of zircon pretreatment, dissolution and U and Pb chemical extraction procedures are described in Ramezani *et al.* (2007).

U and Pb isotopic measurements were performed on a VG Sector-54 multicollector thermal ionization mass spectrometer at MIT, Cambridge, MA. Pb and U were loaded together on a single Re filament in a silica-gel/phosphoric acid mixture (Gerstenberger and Haase 1997). Pb isotopes were measured by peak-hopping using a single Daly photomultiplier detector and U isotopic measurements were made in static mode using multiple Faraday collectors. Data reduction, age calculation, and the generation of concordia plots were carried out using the method of McLean *et al.* (2011), and the statistical reduction and plotting program REDUX (Bowring *et al.* 2011). Unless otherwise noted, U-Pb errors on analyses from this study are reported as: $\pm X/Y/Z$, where X is the internal error in absence of all systematic errors, Y includes the tracer calibration error, and Z includes both tracer calibration and decay constant errors of Jaffey *et al.* (1971)

RESULTS OF U-PB ZIRCON GEOCHRONOLOGY

Zircons extracted from the sample 07SI06B comprise a homogeneous population of clear long-prismatic grains that are interpreted as magmatic in origin. Some of the grains show inclusions but they were not selected for dating. All six analyzed fractions yielded a concordant cluster (Fig. 5) with a weighted $^{206}\text{Pb}/^{238}\text{U}$ mean date of $361.09 \pm 0.14/0.21/0.44$ Ma (MSWD = 0.39; Table 1).

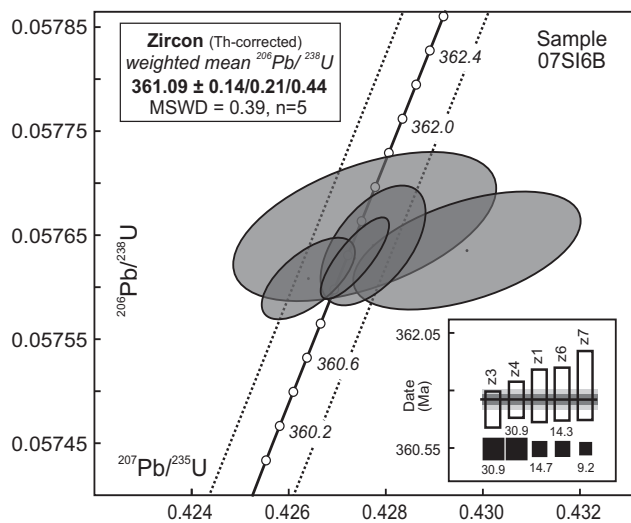


Figure 5. CA-TIMS concordia diagram for Appledore diorite sample 07SI06B. Analytical data are presented in Table 1.

AMPHIBOLE CHEMISTRY

Representative amphibole analyses are listed in Table 2. Compositions range from magnesio-hornblende to actinolitic hornblende (Hawthorne *et al.* 2012). Some analyses of amphibole rims are actinolite. One sample (AI-11) has anomalous amphibole compositions compared to the other analyzed samples. They are also magnesio-hornblende like the majority of the amphibole from the Appledore diorite, but have higher Ti and moderately higher Al than other amphiboles at equivalent $\text{Fe}/(\text{Fe}+\text{Mg})$ values (Table 2).

Amphibole thermobarometric results (Zenk and Schulz 2004) are plotted in Fig. 6. Temperatures range from 410 to 610° C and pressures from 0.2 to 0.6 GPa. The anomalous amphiboles of sample AI-11 also fall within this range, regardless of their higher Ti and Al contents. This temperature range is subsolidus; the decline in temperature with pressure indicates a clockwise, retrograde path for the pluton. The low temperatures indicate that no igneous amphibole has been preserved; hence the pressure at which the pluton was emplaced is unknown. However, the preservation of retrograde amphibole at 0.6 GPa shows that the emplacement pressure was, at minimum, this high.

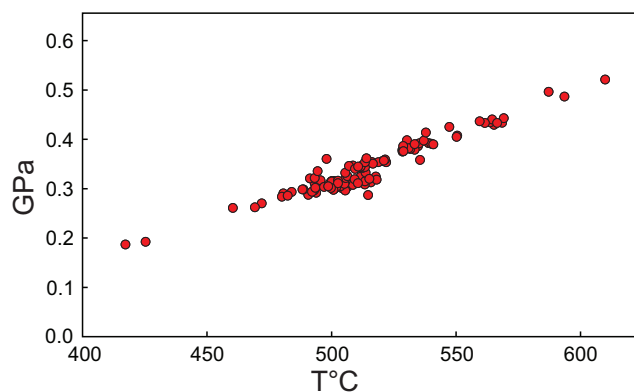


Figure 6. Pressure - temperature plot of amphibole compositions (after Zenk and Schulz 2004), showing that the data indicate subsolidus temperatures extending down to ~ 420°C and pressures ranging from 0.2 to 0.6 GPa.

Table 1. U-Th-Pb data for zircon from Appledore Island diorite. Sample 07SI06B.

Fracture	Composition		Isotopic Ratios										Dates [Ma]				
	Th/U ^(a)	Pb [*] /Pbc ^(b)	^(d) 206Pb/ ^(c) 204Pb	^(e) 208Pb/ ^(c) 206Pb	^(f) 206Pb/ ^(g) 238U	⁽ⁱ⁾ ±2σ	^(h) 207Pb/ ^(g) 235U	^(j) ±2σ	^(k) 206Pb/ ^(g) 238U	^(l) ±2σ	^(m) 207Pb/ ^(h) 235U	⁽ⁿ⁾ ±2σ	^(o) 207Pb/ ^(h) 235U	^(p) ±2σ	^(q) 206Pb/ ^(h) 238U	^(r) ±2σ	Corr.
	[pg]				[%]	[%]	[%]	[%]	[%]	[%]	[abs.]	[abs.]	[abs.]	[abs.]	[abs.]	[abs.]	coef.
z1	0.18	0.4	19.75	1310.7	0.056	0.057617	0.11	0.42925	0.72	0.054057	0.697	0.378	362.641	2.20	372.41	15.7	0.30
z3	0.90	0.3	91.43	4975.5	0.284	0.057595	0.07	0.42638	0.26	0.053717	0.227	0.242	360.606	0.78	358.18	5.1	0.54
z4	0.67	0.4	98.91	5686.7	0.212	0.057616	0.07	0.42745	0.20	0.053831	0.160	0.241	361.113	0.60	362.96	3.6	0.65
z6	0.87	0.4	73.14	4009.4	0.275	0.057627	0.10	0.42772	0.30	0.053855	0.265	0.354	361.556	0.90	364.00	6.0	0.46
z7	0.76	0.6	27.24	1547.8	0.238	0.057639	0.13	0.42730	0.75	0.053791	0.710	0.455	361.257	2.28	361.31	16.0	0.39
z8	0.87	0.9	38.40	2116.4	0.273	0.057704	0.17	0.42633	0.59	0.053608	0.487	0.600	360.565	1.78	353.60	11.0	0.68

Notes: Blank and Oxygen composition: ²⁰⁶Pb/²⁰⁴Pb = 18.42 ± 0.35; ²⁰⁷Pb/²⁰⁴Pb = 15.35 ± 0.23; ²⁰⁸Pb/²⁰⁴Pb = 37.46 ± 0.74; ¹⁸O/¹⁶O = 0.00205 ± 0.00002. ^(a) Th contents calculated from radiogenic ²⁰⁸Pb and the ²⁰⁷Pb/²⁰⁶Pb date of the sample, assuming concordance between U-Th and Pb systems. ^(b) Total mass of radiogenic Pb. ^(c) Total mass of common Pb. ^(d) Ratio of radiogenic Pb (including ²⁰⁸Pb) to common Pb. ^(e) Measured ratio corrected for fractionation and spike contribution only. ^(f) Measured ratios corrected for fractionation, tracer, blank and initial common Pb. ^(g) Corrected for Initial Th/U disequilibrium using radiogenic ²⁰⁸Pb and Th/U [magma] = 2.8. ^(h) Isotopic dates calculated using the decay constants λ238 = 1.55125E-10, λ235 = 9.8485E-10 [Jaffey *et al.* 1971], and for the ²³⁸U/²³⁵U = 137.818 + 0.045 (Hiess *et al.* 2012).

Table 2. Representative amphibole compositions from Appledore Island diorite.

Sample	weight %																cations calculated on the basis of 23 oxygen										
	SiO ₂	TiO ₂	Al ₂ O ₃	MgO	CaO	MnO	FeO	Na ₂ O	K ₂ O	F	Cl	Total	O=F	O=Cl	TOTAL	Si	Ti	Al	Mg	Ca	Mn	Fe	Na	K	F	Cl	Total
11 amph-1	48.35	1.19	6.47	12.92	11.50	0.27	15.31	0.80	0.28	0.15	0.02	97.24	-0.06	0.00	97.18	7.14	0.13	1.13	2.84	1.82	0.03	1.89	0.23	0.05	-0.06	0.00	15.20
11 amph-2	48.72	1.24	5.76	13.34	10.36	0.38	16.05	0.69	0.38	0.18	0.01	97.09	-0.08	0.00	97.02	7.21	0.14	1.01	2.94	1.64	0.05	1.99	0.20	0.07	-0.08	0.00	15.16
11 amph-2	48.82	1.18	5.75	13.35	11.14	0.26	15.31	0.66	0.35	0.21	0.03	97.06	-0.09	-0.01	96.96	7.21	0.13	1.00	2.94	1.76	0.03	1.89	0.19	0.07	-0.09	-0.01	15.14
11 amph-2	48.13	1.05	6.59	13.13	11.42	0.20	15.33	0.75	0.34	0.13	0.01	97.09	-0.05	0.00	97.03	7.12	0.12	1.15	2.90	1.81	0.02	1.90	0.22	0.06	-0.05	0.00	15.24
3 amph-1	47.78	0.48	7.73	10.81	12.03	0.29	17.60	0.78	0.27	0.14	0.03	97.92	-0.06	-0.01	97.85	7.09	0.05	1.35	2.39	1.91	0.04	2.18	0.22	0.05	-0.06	-0.01	15.22
3 amph-1	48.60	0.47	6.58	11.43	12.21	0.21	16.58	0.64	0.29	0.16	0.02	97.19	-0.07	-0.01	97.12	7.22	0.05	1.15	2.53	1.94	0.03	2.06	0.19	0.06	-0.07	-0.01	15.16
3 amph-2	48.95	0.32	6.50	11.94	11.68	0.27	16.60	0.49	0.49	0.16	0.04	97.43	-0.07	-0.01	97.36	7.25	0.04	1.13	2.63	1.85	0.03	2.06	0.14	0.09	-0.07	-0.01	15.15
3 amph-3	47.81	0.48	7.49	10.86	11.67	0.34	17.27	0.70	0.27	0.16	0.02	97.06	-0.07	0.00	96.99	7.14	0.05	1.32	2.42	1.87	0.04	2.16	0.20	0.05	-0.07	0.00	15.17
6 Amph-1	46.98	0.50	8.69	10.38	11.89	0.29	17.44	0.73	0.36	0.17	0.02	97.44	-0.07	0.00	97.37	7.00	0.06	1.53	2.31	1.90	0.04	2.17	0.21	0.07	-0.07	0.00	15.20
6 Amph-1	47.16	0.52	8.51	10.44	11.87	0.40	17.48	0.88	0.35	0.14	0.02	97.75	-0.06	-0.01	97.69	7.01	0.06	1.49	2.31	1.89	0.05	2.17	0.25	0.07	-0.06	-0.01	15.25
6 Amph-1	51.08	0.30	5.01	12.65	11.94	0.34	16.30	0.45	0.16	0.19	0.02	98.43	-0.08	0.00	98.34	7.44	0.03	0.86	2.75	1.86	0.04	1.99	0.13	0.03	-0.08	0.00	15.05
6 Amph-2	45.38	0.93	10.15	9.64	11.73	0.29	17.90	1.04	0.58	0.18	0.04	97.86	-0.08	-0.01	97.78	6.78	0.11	1.79	2.15	1.88	0.04	2.24	0.30	0.11	-0.08	-0.01	15.30
8 Amph-1	43.86	1.02	11.87	8.72	11.96	0.29	18.11	0.96	0.63	0.18	0.04	97.64	-0.08	-0.01	97.55	6.59	0.12	2.10	1.95	1.93	0.04	2.28	0.28	0.12	-0.08	-0.01	15.31
8 Amph-1	46.99	0.58	9.06	10.23	11.86	0.35	17.19	0.82	0.35	0.25	0.02	97.70	-0.11	-0.01	97.59	6.98	0.06	1.59	2.27	1.89	0.04	2.13	0.24	0.07	-0.11	-0.01	15.15
8 Amph-1	48.13	0.47	8.24	10.60	11.81	0.39	16.86	0.78	0.29	0.22	0.02	97.82	-0.09	0.00	97.72	7.11	0.05	1.44	2.33	1.87	0.05	2.08	0.22	0.06	-0.09	0.00	15.11
8 Amph-3	46.28	0.79	10.29	9.63	11.74	0.35	16.87	0.84	0.46	0.13	0.04	97.40	-0.06	-0.01	97.34	6.88	0.09	1.80	2.14	1.87	0.04	2.10	0.24	0.09	-0.06	-0.01	15.19

Note: Samples have Al- prefix.

WHOLE-ROCK GEOCHEMISTRY

The Appledore Island pluton is bimodal but mostly dioritic. The five analyzed dioritic samples contain between 50 to 52 wt. % SiO_2 , whereas the three granitic samples contain between 72 to 74 wt. % SiO_2 (Table 3, Fig. 7). The CIPW normative mineralogy of the diorite samples contains normative hypersthene, and the samples range from olivine tholeiite to quartz tholeiite, indicating subalkaline compositions. The granitic samples are slightly peraluminous with 1–2 % normative corundum.

On both the Zr versus Zr/Y and Zr/Y versus Ti/Y diagrams (Figs. 8a, b), the Appledore Island diorite samples plot in the within-plate basalt field, and are transitional within-plate basalts on the Ti/Y versus Nb/Y diagram (Fig. 8c). They plot at the edge of the within-plate basalt field on the Ti-Zr-Y diagram (Fig. 9a) and in the within-plate alkalic basalt field on the Nb-Zr-Y diagram (Fig. 9b). Chondrite-normalized REE patterns show slight enrichment in

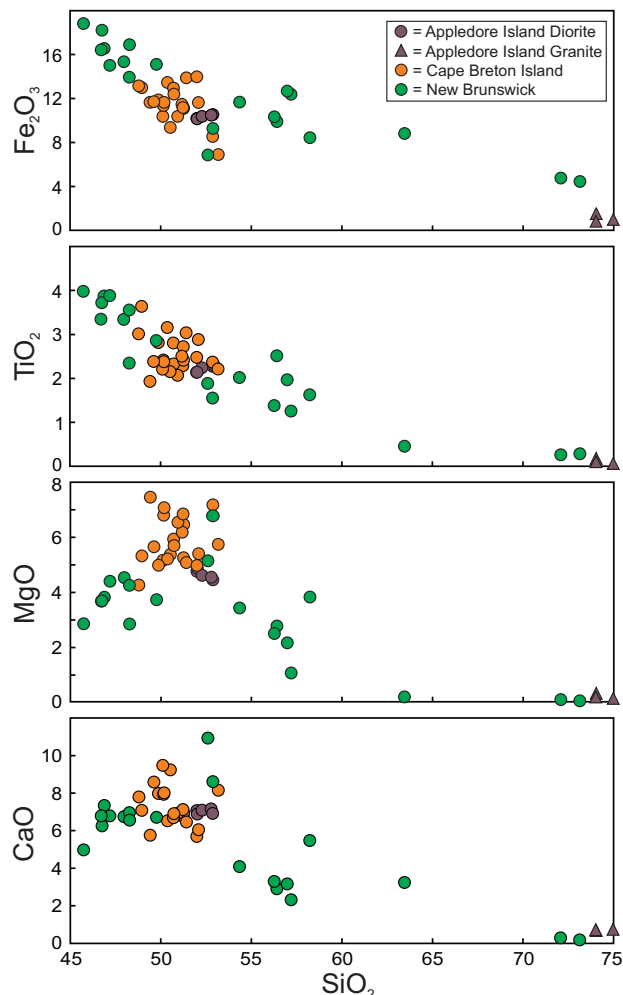


Figure 7. Major element variation diagrams showing the compositions of samples from the Appledore Island diorite and granite. Also shown for comparison are analyses of selected mafic volcanic and plutonic rocks of Devonian-Carboniferous age from Cape Breton Island (Barr *et al.* 1994) and New Brunswick (Fyffe and Barr 1986).

Table 3. Chemical analyses of the Appledore Island diorite.

Sample	AI-1	AI-2	AI-3	AI-4	AI-5	AI-6	AI-7	AI-8
SiO_2 (wt.%)	51.17	51.07	51.44	72.95	72.73	51.88	74.73	51.96
TiO_2	2.10	2.10	2.21	0.17	0.10	2.23	0.06	2.23
Al_2O_3	16.44	16.28	16.22	14.35	14.27	16.15	14.31	16.23
Fe_2O_3	10.02	9.97	10.21	1.51	0.82	10.34	0.99	10.31
MnO	0.13	0.17	0.12	0.00	0.00	0.15	0.00	0.14
MgO	4.69	4.79	4.54	0.30	0.17	4.37	0.11	4.48
CaO	6.96	6.76	6.98	0.67	0.71	6.80	0.74	7.04
Na_2O	3.73	3.65	3.47	3.15	4.04	2.82	4.45	2.60
K_2O	2.19	2.40	2.19	5.24	5.27	2.34	4.15	2.40
P_2O_5	0.98	0.99	1.03	0.20	0.16	1.03	0.13	1.04
LOI	0.61	0.69	0.65	0.62	0.29	0.74	0.50	0.71
Total	99.02	98.87	99.06	99.16	98.56	98.85	100.17	99.14
Rb (ppm)	48.1	75.1	50.5	281	224	77.8	310	107.5
Sr	524	588	543	65	67.8	538	120	603
Ba	466	747	524	173	197	546	284	587
Cs	1.71	3.4	1.29	3.85	3.1	3.79	5.09	6.56
La	39.4	48.2	49.4	20.5	11.9	47.4	9.1	58.4
Ce	89.5	108.5	109	49.3	27.6	106.5	19.1	126
Pr	12.4	14.95	14.85	7	3.8	15.1	2.34	17.45
Nd	49.9	59.9	59.8	27.3	14.4	61.6	8.4	70.6
Sm	9.44	10.9	11.1	7.68	3.91	11.45	1.81	12.75
Eu	2.59	2.97	2.92	0.45	0.45	3.05	0.39	3.5
Gd	8.88	10.55	10.55	6.91	3.74	10.95	1.9	12.25
Tb	1.11	1.27	1.3	0.88	0.5	1.35	0.34	1.52
Dy	5.79	6.67	6.6	3.89	2.37	6.97	2.13	7.89
Ho	1.12	1.24	1.21	0.56	0.38	1.31	0.4	1.48
Er	2.89	3.4	3.29	1.3	0.91	3.49	1.2	3.91
Tm	0.38	0.44	0.44	0.15	0.12	0.45	0.19	0.52
Yb	2.31	2.71	2.62	0.96	0.76	2.69	1.21	3.14
Lu	0.35	0.41	0.38	0.14	0.11	0.42	0.18	0.48
Zr	275	334	326	69	45	343	52	408
Y	23.1	27.2	26.5	13.4	8.8	27.9	10.2	32.4
Hf	7	8.2	8	2.9	1.8	8.6	2.4	9.8
Nb	29.4	34.4	34	14	7.3	35.5	15.7	39.9
Ta	1.6	1.9	1.8	1.3	0.7	2	1.9	2.1
Th	3.74	4.38	4.25	15	8.17	4.76	5.01	4.66
U	1.18	1.58	1.45	6.07	3.63	1.78	6.94	1.91
Pb	10	11	13	37	29	13	30	12
V	115	138	136	<5	<5	137	<5	157
Cr	70	80	70	<10	<10	70	<10	80
Co	25.2	29.9	26.3	15.4	23.7	28.5	20.7	30.3
Ni	30	37	28	<5	<5	22	<5	28
Cu	16	21	18	<5	<5	18	<5	17
Ga	18.5	21.5	20.8	21.6	15.7	22.1	20.5	24.4
Zn	117	145	133	39	23	137	33	150
Sn	2	3	2	2	2	2	4	3
W	68	71	57	216	346	107	286	80

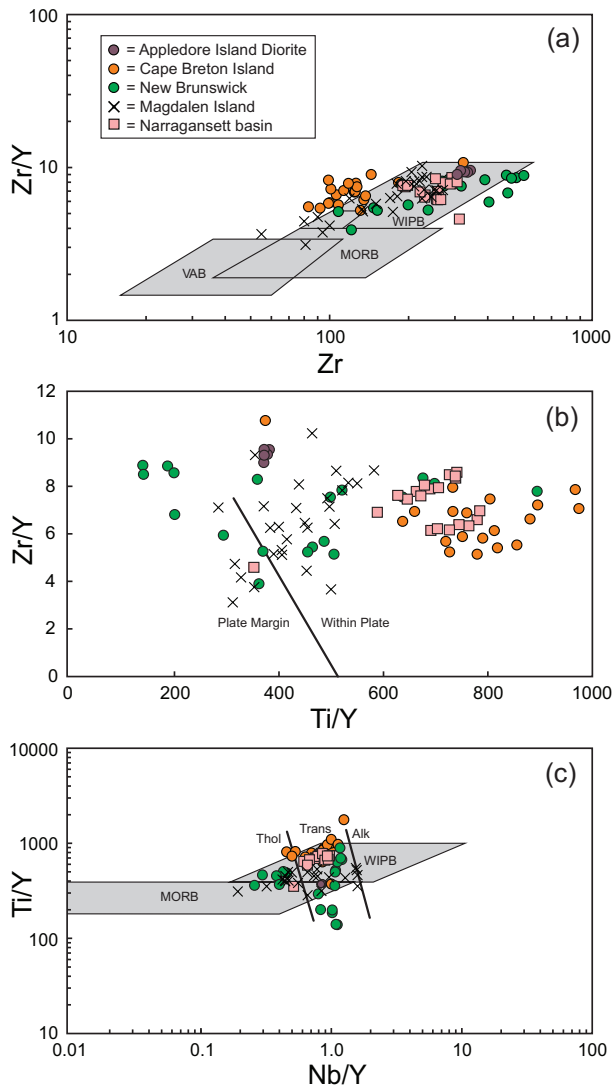


Figure 8. Chemical discrimination diagrams for Appledore Island diorite analyses. (a) Zr/Y versus Zr diagram with fields from Pearce and Norry (1979). VAB = volcanic arc basalts, MORB = mid oceanic ridge basalts, WIPB = within plate basalts. (b) Zr/Y versus Ti/Y diagram with fields from Pearce and Gale (1977). (c) Ti/Y versus Nb/Y diagram with fields from Pearce (1982). Abbreviations: Thol = tholeiitic basalt, Trans = transitional basalt, Alk = alkalic basalt, MORB = mid oceanic ridge basalt, WIPB = within plate basalt. Also shown for comparison are data from mafic volcanic and plutonic rocks in Cape Breton Island (Barr *et al.* 1994), New Brunswick (Fyffe and Barr 1986; Park *et al.* 2014), the Magdalen Islands (Barr *et al.* 1985; La Flèche *et al.* 1998), and Narragansett Basin (Maria and Hermes 2001).

the LREE, ranging between 100 and 250 times chondrites (Fig. 10a). The HREE patterns are relatively flat at about 15 to 20 times chondrite, giving $(La/Yb)_N$ values of 11.5–12.7. Eu anomalies are either absent or very small. Multi-element

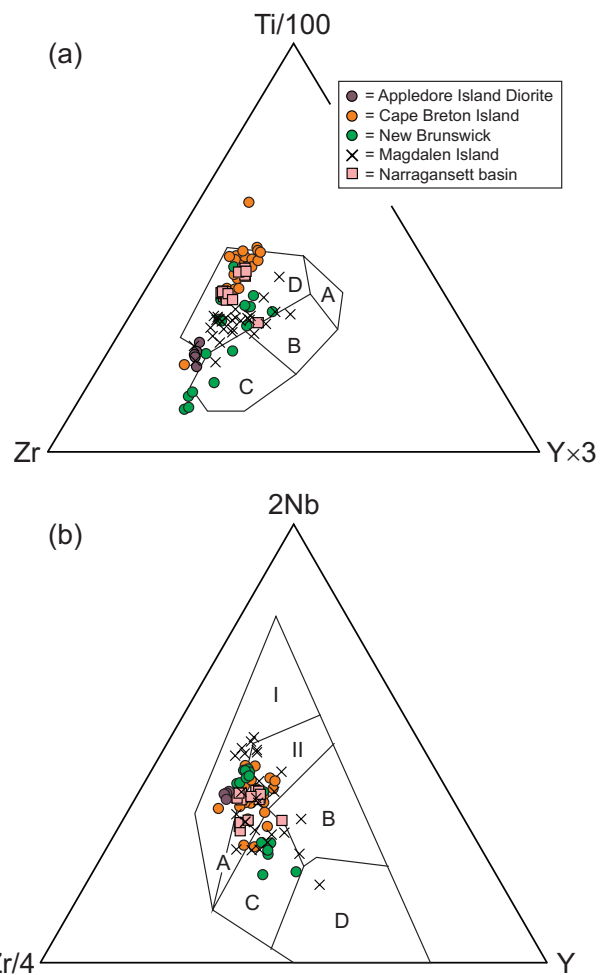


Figure 9. Chemical discrimination diagrams for Appledore Island diorite analyses. (a) Ti - Zr - Y diagram of Pearce and Cann (1973). Field A = island arc tholeiite; B = MORB, island arc tholeiite, and calc-alkaline basalts; C = calc-alkaline basalt; D = within plate basalt. (b) Nb - Zr - Y diagram (after Meschede 1986). Also shown for comparison are data from mafic volcanic and plutonic rocks in Cape Breton Island (Barr *et al.* 1994), New Brunswick (Fyffe and Barr 1986; Park *et al.* 2014), the Magdalen Islands (Barr *et al.* 1985; La Flèche *et al.* 1998), and Narragansett Basin (Maria and Hermes 2001).

normalized diagrams show moderate enrichment in the most incompatible elements with values near 100 times primitive mantle (Fig. 10b). Compared to neighboring Th and U, Nb is not depleted; no significant negative anomalies are evident. From La to Lu, the patterns show gradual decreasing normalized values.

Three analyses of granitic rocks that intruded and mingled with the diorite of Appledore Island are plotted in tectonic discrimination diagrams in Figure 11. The granite samples plot in the syn-collisional granite fields.

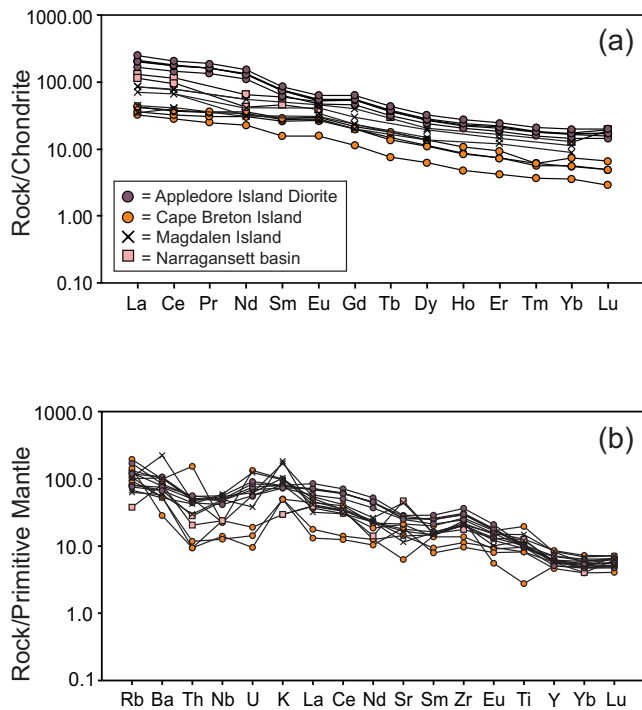


Figure 10. (a) Chondrite-normalized REE patterns for Appledore Island diorite samples and rocks from Cape Breton Island, New Brunswick, and Narragansett Basin (Barr *et al.* 1985; 1994; Fyffe and Barr, 1986; Maria and Hermes 2001). (b) Primitive mantle-normalized incompatible trace element patterns for Appledore Island diorite samples and mafic rocks from Narragansett Basin (Maria and Hermes 2001), Magdalen Islands (La Flèche *et al.* 1998), Cape Breton Island (Pe-Piper and Piper 1998), and New Brunswick (Park *et al.* 2014).

ISOTOPIC COMPOSITIONS

The $\epsilon_{\text{Nd}(361 \text{ Ma})}$ values in four samples from the Appledore Island diorite are somewhat depleted, plotting at about +2 (Fig. 12; Table 4). Initial $^{87}\text{Sr}/^{86}\text{Sr}$ values are near bulk silicate Earth, ranging from 0.704085 to 0.704420. In contrast, the granitic samples have negative $\epsilon_{\text{Nd}(361 \text{ Ma})}$ ranging from -1.3 to -6.1, with corresponding initial Sr ratios ranging from 0.705596 to 0.706155, considerably higher than the initial Sr ratios of the dioritic samples.

Age-corrected $^{207}\text{Pb}/^{204}\text{Pb}$ data from diorite and granite samples are plotted against $\epsilon_{\text{Nd}(361 \text{ Ma})}$ on Figure 13. For reference, the figure also shows fields for Laurentian, Ganderian, and other non-North American (including Avalonian) crust as compiled by Tomascak *et al.* (2005), and a field for Mesozoic tholeiitic basalt from the Central Atlantic Magmatic Province (Pegram 1990). Data from the

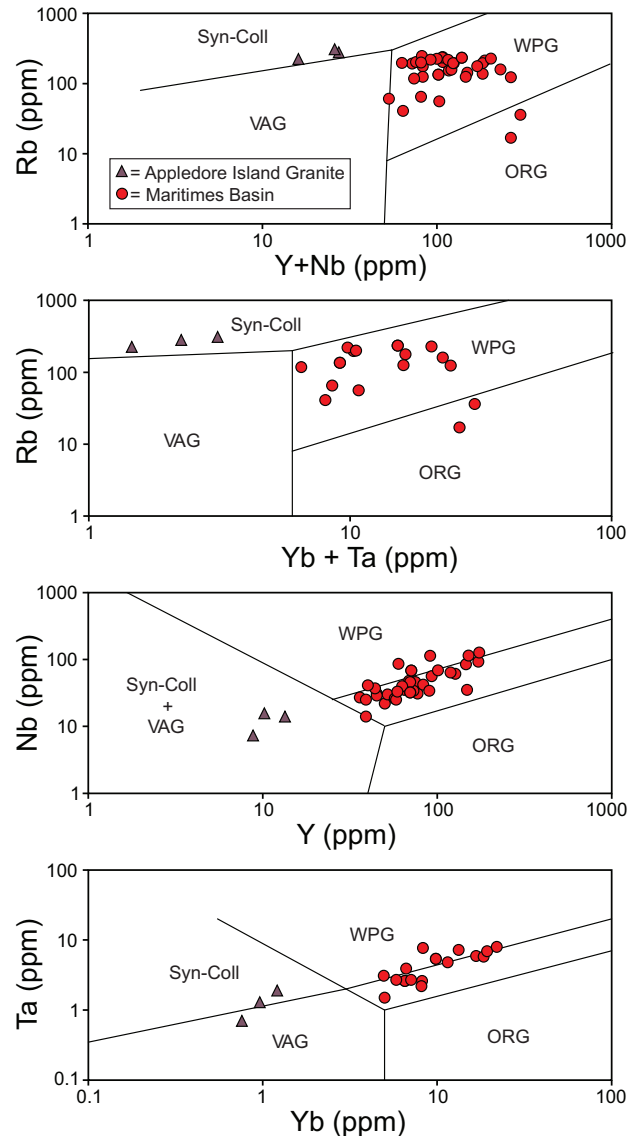


Figure 11. Appledore Island granite analyses plotted on chemical discrimination diagrams for felsic rocks with fields from Pearce *et al.* (1984): (a) Rb versus Yb+Ta; (b) Rb versus Y+Nb; (c) Nb versus Y; (d) Ta versus Yb. VAG = volcanic arc granites, WPG = within-plate granites, ORG = ocean ridge granites, Syn-Coll = syncollisional granites. Maritime Canada samples are from Pe-Piper *et al.* (1989, 1991); Piper *et al.* (1996); and Park *et al.* (2014).

Appledore Island diorite samples plot in the Mesozoic tholeiite field. The granitic samples from Appledore Island have $^{207}\text{Pb}/^{206}\text{Pb}$ values similar to those from the associated dioritic rocks, but have lower $\epsilon_{\text{Nd}(361 \text{ Ma})}$ values. Two of the granitic samples plot adjacent to the Ganderian field and one plots in the area of overlap between the Ganderian and Avalonian fields.

Table 4. Nd, Sr, and Pb isotopic compositions of the Appledore Island diorite.

Sample	Nd	Sm	$^{147}\text{Sm}/^{144}\text{Nd}$	$^{143}\text{Nd}/^{144}\text{Nd}$	$\epsilon_{(0)}$	$\epsilon_{(361\text{ Ma})}$	$T_{\text{(DM)}}$	Rb	Sr	$^{87}\text{Sr}/^{86}\text{Sr}_m$	$^{87}\text{Sr}/^{86}\text{Sr}_i$	$^{208}\text{Pb}/^{204}\text{Pb}$	$^{207}\text{Pb}/^{204}\text{Pb}$	$^{206}\text{Pb}/^{204}\text{Pb}$
AI-1	53.35	9.94	0.1126	0.512518	-2.3	1.5	954 Ma	48.1	524	0.705803	0.704420	38.57	15.66	18.76
AI-3	59.37	10.76	0.1096	0.512516	-2.4	1.6	934 Ma	50.5	543	0.705573	0.704194	38.53	15.65	18.74
AI-4	27.3	7.68	0.1702	0.512265	-7.3	-6.1	-	281	65	0.770084	0.705596	38.69	15.70	19.33
AI-5	14.4	3.91	0.1643	0.512278	-7.0	-5.5	-	224	67.8	0.755112	0.705900	38.50	15.68	19.17
AI-6	57.06	10.61	0.1125	0.512521	-2.3	1.6	948 Ma	77.8	538	0.706278	0.704134	38.56	15.65	18.85
AI-7	8.4	1.81	0.1304	0.512414	-4.4	-1.3	1.35 Ga	310	120	0.744596	0.706155	38.35	15.70	19.60
AI-8	61.97	11.17	0.1089	0.512510	-2.5	1.5	932 Ma	107.5	603	0.706728	0.704085	38.66	15.68	18.96

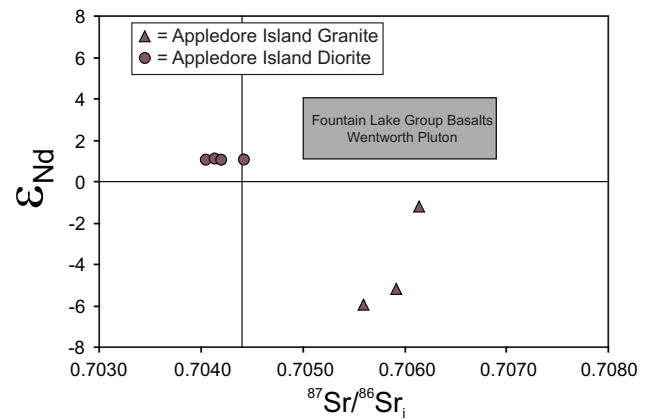


Figure 12. Epsilon Nd $_{(361\text{ Ma})}$ versus Sr isotope ratio for four samples of the Appledore Island diorite and three samples of the Appledore Island granite. The grey rectangle shows the range of values for the Fountain Lake Group basalt and the Wentworth Pluton (Dessureau *et al.* 2000; Pe-Piper and Piper, 1998).

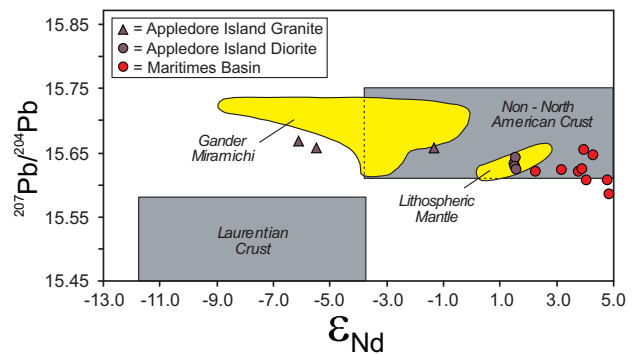


Figure 13. Age-corrected $^{207}\text{Pb}/^{204}\text{Pb}$ and Nd isotopic values for Appledore Island diorite and granite samples and basalt from the Magdalen Basin (Pe-Piper and Piper 1998). Also plotted are reference fields for Laurentia, Ganderia, and Avalonian and related terranes from Tomascak *et al.* (2005) and for Mesozoic flood basalts of eastern North America (Pegram 1990).

DISCUSSION

To identify the tectonic affinity of the Rye Complex, we compare the composition of the Appledore Island diorite to mafic rocks of the same age to the northeast in Maritime Canada and the Narragansett Basin to the south. As discussed below, the Appledore Island diorite is the plutonic equivalent of rocks of continental tholeiite derivation across the northern Appalachians whose compositions are indicative of Late Devonian to Early Carboniferous extension. In this contribution, we limit our comparisons to lavas and plutons in coastal New England and Maritime Canada of these same compositional signatures and tectonic settings that were erupted/emplaced in post-Acadian succes-

sor basins (Blanchard *et al.* 1984; Durling and Marillier 1993; Pe-Piper and Piper 2002). These include the ~ 340 Ma basalts of Cape Breton Island (Blanchard *et al.* 1984; Barr *et al.* 1994), Carboniferous basalts of New Brunswick (Fyffe and Barr 1986; Park *et al.* 2014), 385 Ma rocks of the Guysborough area, Nova Scotia (Cormier *et al.* 1995), 355–362 Ma basalts of the Magdalen Basin (Barr *et al.* 1985; Pe-Piper and Piper 1998, La Flèche *et al.* 1998), 330 to 375 Ma Fountain Lake Group basalts of Nova Scotia (Dessureau *et al.* 2000), and ~ 370 Ma Wamsutta Formation in the Narragansett Basin of southern New England (Thompson and Hermes 2003).

The Maritimes Basin of Atlantic Canada contains basalt, rhyolite, and related gabbro-granite plutons which range in age from ca. 375 Ma to 330 Ma (Dunning *et al.* 2002), thus overlapping with the age of the Appledore Island diorite. On Harker and tectonic discrimination diagrams, the mafic rocks have within-plate tholeiitic characteristics, and are very similar to the Appledore Island diorite (Figs. 7, 8, 9). The majority of Magdalen Basin samples have lower REE concentrations compared to the Appledore Island diorite but REE patterns are generally parallel to those of the Appledore Island diorite samples (Fig. 10a). Magdalen Island and Cape Breton Island samples have similar normalized multi-element patterns to the Appledore Island diorite (Fig. 10b).

The Wamsutta Formation in the Narragansett Basin contains bimodal volcanic rocks (Maria and Hermes 2001) including a 373 ± 2 Ma rhyolite (best estimate from upper Concordia intercept and weighted mean $^{206}\text{Pb}/^{238}\text{U}$ and $^{207}\text{Pb}/^{235}\text{U}$ ages of Thompson and Hermes 2003). Associated Wamsutta Formation basalts are thus probably somewhat older than the Appledore Island diorite. However, like the diorite they are within-plate tholeiites (Figs. 7, 8, 9). Their LREE concentrations approach those of the Appledore Island diorite samples, but with slightly lower $(\text{La}/\text{Yb})_{\text{N}}$ values of 7.4–9.1 (Fig. 10a). Two basaltic samples display normalized multi-element patterns similar to those of the Appledore Island diorite samples (Fig. 10b).

Basalt of the Fountain Lake Group and gabbro from the Wentworth Pluton from the Cobequid Highlands in the Maritimes Basin yielded $\epsilon_{\text{Nd}(361 \text{ Ma})}$ ranging from 1 to 6 (Dessureau *et al.* 2000; Pe-Piper and Piper, 1998), overlapping with the data from the Appledore Island diorite. Few Sr data are available for mafic rocks from the Maritimes Basin, but initial $^{87}\text{Sr}/^{86}\text{Sr}$ values for quartz diorite from the Gilbert-Wyvern pluton of the western Cobequid Highlands range from 0.7069 to 0.7076 (Pe-Piper and Piper 1998). Age-corrected $^{207}\text{Pb}/^{204}\text{Pb}$ and Nd isotopic data from the diorite from Appledore Island and volcanic rocks from the Maritimes Basin define a trend extending from the Mesozoic tholeiite field to higher $\epsilon_{\text{Nd}(361 \text{ Ma})}$ values of up to 5. All samples of both Appledore Island and the Maritimes Basin have age-corrected $^{207}\text{Pb}/^{206}\text{Pb}$ values of ~ 15.6 to 15.65.

The similarities in major element, trace element and isotopic compositions and age between the Appledore Island

diorite and continental rift tholeiites of Maritime Canada and the Narragansett Basin indicate that the Appledore Island diorite is a component in the extended magmatic episode that was expressed from Rhode Island to Cape Breton Island. Its age is especially similar to that of basalt, rhyolite, and equivalent plutonic rocks in the Cobequid Highlands which have yielded ages of ca. 360–356 Ma (Dunning *et al.* 2002). Correlation of the Appledore Island diorite with the mafic rocks of the post-Acadian successor basins implies that the rocks of the Rye Complex which it intrudes likely represents the basement of a post-Acadian successor basin, intruded by mafic and felsic magmas related to basin formation. The other basins are filled with sediments and intercalated volcanic rocks, all deposited on peri-Gondwanan basements (Durling and Marillier 1993). But unlike the Late Devonian to Permian successions in the Maritimes Basins where sediments unconformably overlie either Ganderian or Avalonian basement, depending on the location of the subbasin (Hibbard *et al.* 2006), the Appledore Island diorite intruded a crystalline block that had experienced amphibolite-facies metamorphic conditions (Blomshield 1975). We interpret those rocks as basement; sedimentary and volcanic cover, if ever present, have long been removed. The major question regarding the Rye Complex, then, is this peri-Gondwanan basement Ganderian or Avalonian?

With the exception of the peraluminous Black Brook Granitic Suite (Yaowanoioyothin and Barr 1991), the Maritimes Basin contains 360–340 Ma A-type granites (Pe-Piper *et al.* 1989, 1991; Piper *et al.* 1996; Park *et al.* 2014). In contrast, the granitic rocks associated with the Appledore Island diorite are mildly peraluminous with syn-collisional trace element characteristics (Fig. 11). While these data seem to be at odds with the within-plate signature of the comagmatic diorite, the granite signature may not provide an accurate indication of the tectonic setting of the Appledore Island pluton. As shown by Förster *et al.* (1997), granite geochemistry may instead reflect the tectonic setting of the source rocks. Partial melting in the contact aureole probably occurred by mica dehydration melting reactions, and because both muscovite and biotite have large partition coefficients for Rb (Rollinson 1993), mica breakdown partitions large amounts of Rb into the melt, accounting for the high Rb/Sr ratios and high Rb concentrations with the compositions of syn-collisional granites (Harris *et al.* 1993). In contrast, partial melting influenced by high water contents causes higher degrees of melting of quartz and feldspar with minimal mica melting, and because feldspars have high Sr partition coefficients and lower Rb partition coefficients than micas, the melt has lower Rb/Sr ratios than melts derived by mica dehydration melting. Thus the Rb/Sr ratios and Rb concentrations are defined by the nature of the melting reaction, which in syn-collisional settings, is dominantly mica dehydration melting. Hence, although the granites of the Appledore Island pluton plot in the syn-collisional fields of Figure 11, they very well may result from mica dehydration melting in the contact aureole rather than

a particular tectonic setting. The metasedimentary rocks of the Rye Complex clearly have experienced amphibolite-facies metamorphism at pressures at least as high as those recorded by the retrograde amphibole in Appledore Island diorite, where pressures reached as high as 0.6 GPa. Hence, we interpret the granites of the Appledore pluton to represent partial melts of the Rye Complex basement.

The Nd and Pb isotopic data (Fig. 13) suggest that the granites are partial melts of Ganderian sources. One granitic sample, (AI-7), plots in the overlap zone between Ganderia and Avalonia, but the other two samples have $\epsilon_{\text{Nd}(361 \text{ Ma})}$ values that are more negative than typical Avalonian basement. Some Avalonian cover sediments have similar low ϵ_{Nd} values (Samson *et al.* 2000; Sakoski *et al.* 2007), but they are unmetamorphosed sedimentary rocks overlying Avalonian basement. Therefore, the peraluminous granite of the Appledore Island granite is interpreted to be a partial melt of the Rye Complex basement and that basement has isotopic compositions that are more indicative of Ganderia than Avalonia.

CONCLUSIONS

The 361.09 ± 0.14 Ma diorite on Appledore Island is compositionally similar to other rift-related, plutonic and volcanic continental tholeiites from the Maritimes Basin of Canada and the Wamsutta Formation of the Narragansett Basin of Rhode Island. The tholeiites are thought to be generated during regional extension after the Acadian orogeny, producing post-Acadian successor basins. Slightly peraluminous granitic rocks occur along the margin of the Appledore Island pluton, and the isotopic composition of the granites indicates that they are partial melts of Ganderian basement rocks. Thus the Rye Complex correlates with other peri-Gondwanan terranes that underwent extension and intrusion of continental rift tholeiites, and the Ganderian isotopic signature of the Appledore Island granitic rocks indicates that the Rye Complex basement is Ganderian.

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REFERENCES

- Aleinikoff, J.N., Wintsch, R.P., Tollo, R.P., Unruh, D.M., Fanning, C.M., and Schmitz, M.D. 2007. Ages and origins of rocks of the Killingworth dome, south-central Connecticut: Implications for the tectonic evolution of southern New England. *American Journal of Science*, 307, pp. 63–118. <http://dx.doi.org/10.2475/01.2007.04>
- Barr, S.M., Brisebois, D., and MacDonald, A.S. 1985. Carboniferous volcanic rocks of the Magdalen Islands, Gulf of St. Lawrence. *Canadian Journal of Earth Sciences*, 22, pp. 1679–1688. <http://dx.doi.org/10.1139/e85-176>
- Barr, S.M., Grammatikopoulos, A.L., and Dunning, G.R. 1994. Early Carboniferous gabbro and basalt in the St. Peters area, southern Cape Breton Island, Nova Scotia. *Atlantic Geology*, 30, pp. 247–258.
- Barr, S.M., Raeside, R.P., and White, C.E. 1998. Geological correlations between Cape Breton Island and Newfoundland, northern Appalachian Orogen. *Canadian Journal of Earth Sciences*, 35, pp. 1252–1270. <http://dx.doi.org/10.1139/e98-016>
- Billings, M.P. 1956. *Geology of New Hampshire: Part II, Bedrock Geology*. New Hampshire Planning and Development Commission, Concord, New Hampshire, 203 p.
- Blanchard, M-C, Jamieson, R., and More, E. 1984. Late Devonian - Early Carboniferous volcanism in western Cape Breton Island, Nova Scotia. *Canadian Journal of Earth Sciences*, 21, pp. 762–774. <http://dx.doi.org/10.1139/e84-083>
- Blomshield, R.J. 1975. Superposed deformation on the Isles of Shoals, Maine - New Hampshire. Unpublished M.S. thesis, University of New Hampshire, Durham, New Hampshire, 54 p.
- Bothner, W.A. and Hussey, A.M. II. 1999. Norumbega connections: Casco Bay, Maine to Massachusetts? *In* Norumbega Fault System of the Northern Appalachians. *Edited by* Ludman, A. and West, D.P., Jr. *Geological Society of America Special Paper*, 331, pp. 59–72. <http://dx.doi.org/10.1130/0-8137-2331-0.59>
- Bothner, W.A., Blackburn, T., Bowring, S., Buchwaldt, R., and Hussey, A.M. II. 2009. Temporal constraints on Paleozoic plutonism in southwestern Maine and southeastern New Hampshire: Revisions and implications. *Geological Society of America Abstracts with Programs*, 41, p. 32.
- Bowring, J.F., McLean, N.M., and Bowring, S.A. 2011. Engineering cyber infrastructure for U-Pb geochronology; Tripoli and U-Pb Redux. *Geochemistry, Geophysics, Geosystems*, 12, QOAA19.
- Brooks, J. A. 1986. *Geology of New Hampshire's Inner Continental Shelf*. Unpublished M.S. thesis, University of New Hampshire, Durham, New Hampshire, 137 p.
- Brooks, J. A. 1990. *The Petrogenesis of the Agamenticus Complex and Late Paleozoic and Mesozoic Tectonics in New England*. Unpublished Ph.D dissertation, University of New Hampshire, Durham, New Hampshire, 317 p.

- Carrigan, J.A. 1984. Geology of the Rye Formation: New Castle Island and adjacent areas of Portsmouth Harbor, New Hampshire and Maine. Unpublished M.S. thesis, University of New Hampshire, Durham, New Hampshire, 128 p.
- Cormier, C.F.M, Barr, S.M., and Dunning, G.R. 1995. Geological setting and petrochemistry of early Middle Devonian volcanic and gabbroic rocks in the Guysborough area, Nova Scotia. *Atlantic Geology*, 31, pp. 153–166.
- Dessureau, G., Piper, D.J.W., and Pe-Piper, G. 2000. Geochemical evolution of earliest Carboniferous continental tholeiitic basalts along a crustal-scale shear zone, southwestern Maritimes basin, eastern Canada. *Lithos*, 50, pp. 27–50. [http://dx.doi.org/10.1016/S0024-4937\(99\)00042-0](http://dx.doi.org/10.1016/S0024-4937(99)00042-0)
- Dunning, G.R., Barr, S.M., Giles, P.S., McGregor, D.C., Pe-Piper, G., and Piper, D.J.W. 2002. Chronology of Devonian to early Carboniferous rifting and igneous activity in southern Magdalen Basin based on U-Pb (zircon) dating. *Canadian Journal of Earth Sciences*, 39, pp. 1219–1237. <http://dx.doi.org/10.1139/e02-037>
- Durling, P. and Marillier, F. 1993. Tectonic setting of Middle Devonian to Lower Carboniferous rocks in the Magdalen Basin. *Atlantic Geology*, 29, pp. 199–217.
- Förster, H.-J., Tischendorf, G., and Trumbull, R.B. 1997. An evaluation of the Rb vs. (Y+Nb) discrimination diagram to infer tectonic setting of silicic igneous rocks. *Lithos*, 40, pp. 261–293. [http://dx.doi.org/10.1016/S0024-4937\(97\)00032-7](http://dx.doi.org/10.1016/S0024-4937(97)00032-7)
- Fowler-Billings, K. 1959. Geology of the Isles of Shoals. State of New Hampshire Department of Resources and Economic Development, Concord, New Hampshire, 47 p.
- Fyffe, L.R. and Barr, S.M. 1986. Petrochemistry and tectonic significance of Carboniferous volcanic rocks in New Brunswick. *Canadian Journal of Earth Sciences*, 23, pp. 1243–1256. <http://dx.doi.org/10.1139/e86-121>
- Gerstenberger, H. and Haase, G. 1997. A highly effective emitter substance for mass spectrometric Pb isotope ratio determinations. *Chemical Geology*, 136, pp. 309–312. [http://dx.doi.org/10.1016/S0009-2541\(96\)00033-2](http://dx.doi.org/10.1016/S0009-2541(96)00033-2)
- Goldstein, S.L., O’Nions, R.K., Keith, R., and Hamilton, P.J. 1984. A Sm-Nd isotopic study of atmospheric dusts and particulates from major river systems. *Earth and Planetary Science Letters*, 70, pp. 221–236. [http://dx.doi.org/10.1016/0012-821X\(84\)90007-4](http://dx.doi.org/10.1016/0012-821X(84)90007-4)
- Harris, N., Inger, S., and Massey, J. 1993. The role of fluids in the formation of High Himalayan leucogranites. *In Himalayan Tectonics. Edited by P.J. Treloar and M.P. Searle*, M.P. Geological Society of London Special Publications 74, pp. 391–400.
- Hawthorne, F.C., Oberti, R., Harlow, G.E. Maresch, W.V., Martin, R.F., Schumacher, J.C., and Welch, M.D. 2012. Nomenclature of the amphibole supergroup. *American Mineralogist*, 97, pp. 2031–2048. <http://dx.doi.org/10.2138/am.2012.4276>
- Hibbard, J.P., van Staal, C.R., Rankin, D.W., and Williams, H. 2006. Lithotectonic map of the Appalachian orogen, Canada and United States of America. Geological Survey of Canada, Map 2096A, scale 1:1 500 000.
- Hiess, J. Condon, D.J., McLean, N., and Noble, S.R. 2012. $^{238}\text{U}/^{235}\text{U}$ systematics in terrestrial uranium-bearing minerals. *Science*, 335, 1610–1614. <http://dx.doi.org/10.1126/science.1215507>
- Hussey, A.M. 1980. The Rye Formation of Gerrish Island, Kittery, Maine. *The Maine Geologist*, 7, pp 2–3.
- Hussey, A.M.II, Bothner, W.A., and Aleinikoff, J. 2010. The tectono-stratigraphic framework and evolution of southwestern Maine and southeastern New Hampshire. *In From Rodinia to Pangea: The Lithotectonic Record of the Appalachian Region. Edited by R.P. Tolo, M.J. Bartholomew, J.P. Hibbard, and P.M. Karabinos*. Geological Society of America Memoir, 206, pp. 205–230.
- Hussey, A.M. II, Bothner, W.A., and Thompson, P.J. in press. Bedrock Geology of the Kittery 100 000 Quadrangle, Maine and New Hampshire: Maine Geological Survey Bulletin.
- Jaffey, A.H., Flynn, K.F., Glendenin, L.E., Bentley, W.C., and Essling, A.M. 1971. Precision measurement of half-lives and specific activities of ^{235}U and ^{238}U . *Physical Review*, 4, pp. 1889–1906.
- Kane, P., Hepburn, J.C., and Bothner, W.A. 2014. Age of the Rye Complex, SE New Hampshire: A detrital zircon study. *Geological Society of America Abstracts with Programs*, 46, p. 118.
- Katz, F.J. 1917. Stratigraphy in southwestern Maine and southeastern New Hampshire. U. S. Geological Survey Professional Paper 108-I, pp. 165–177.
- Kuiper, Y.D., Hepburn, J.C., Kay, A., Loan, M.L., Sorota, K.J., Reynolds, E.C., Dabrowski, D.R., Stroud, M.M., and Markwort, R.J. 2013. Origin and evolution of the Nashoba and Merrimack terranes in southeastern New England: What have we learned in the past decade? *Geological Society of America Abstracts with Programs*, 45, p. 64.
- La Flèche, M.R., Camine, G., and Jenner, G.A. 1998. Geochemistry of post-Acadian, Carboniferous continental intraplate basalts from the Maritimes Basin, Magdalen Islands, Québec, Canada. *Chemical Geology*, 148, pp. 115–136. [http://dx.doi.org/10.1016/S0009-2541\(98\)00002-3](http://dx.doi.org/10.1016/S0009-2541(98)00002-3)
- Lyons, J.B., Bothner, W.A., Moench, R.H., and Thompson, J.B. 1997. Bedrock geologic map of New Hampshire. United States Geological Survey, scale 1: 500 000.
- Maria, A. and Hermes, O.D. 2001. Volcanic rocks in the Narragansett Basin, southeastern New England: Petrology and significance to early basin formation. *American Journal of Science*, 301, pp. 286–312. <http://dx.doi.org/10.2475/ajs.301.3.286>
- Mattinson, J.M. 2005. Zircon U/Pb chemical abrasion (CA-TIMS) method; combined annealing and multi-step partial dissolution analysis for improved precision and accuracy of zircon ages. *Chemical Geology*, 220, pp. 47–66. <http://dx.doi.org/10.1016/j.chemgeo.2005.03.011>

- McLean, N.M., Bowring, J.F., and Bowring, S.A. 2011. An algorithm for U-Pb isotope dilution data reduction and uncertainty propagation. *Geochemistry, Geophysics, Geosystems*, 12, QOAA18.
- Meschede, M. 1986. A method of discriminating between different types of mid-ocean ridge basalts and continental tholeiites with the Nb-Zr-Y diagram. *Chemical Geology*, 56, pp. 207–218. [http://dx.doi.org/10.1016/0009-2541\(86\)90004-5](http://dx.doi.org/10.1016/0009-2541(86)90004-5)
- Miller, C.F., Stoddard, E.F., Bradfish, L.J., and Dollase, W.A. 1981. Composition of plutonic muscovite: genetic implications. *Canadian Mineralogist*, 19, pp. 25–34.
- Nance, R.D., Murphy, J.B., Strachan, R.A., Keppie, J.D., Gutierrez-Alonso, G., Fernandez-Suarez, J., Quesada, C., Linnemann, U., D'Lemos, R., and Pisarevsky, S.A. 2008. Neoproterozoic-early Palaeozoic tectonostratigraphy and palaeogeography of the peri-Gondwanan terranes: Amazonian versus West African connections. *Geological Society of London Special Publication* 297, pp. 345–383. <http://dx.doi.org/10.1144/SP297.17>
- Novotny, R.F. 1963. Bedrock geology of the Dover-Exeter-Portsmouth region, New Hampshire. Unpublished Ph.D. thesis, Ohio State University, Columbus, Ohio, 128 p.
- Novotny, R.F. 1969. The geology of the seacoast region of New Hampshire. *Edited by* T.R. Meyers. New Hampshire Department of Resources and Economic Development, Concord, New Hampshire, 46 p.
- Park, A.F., Treat, R.L., Barr, S.M., White, C.E., Miller, B.V., Reynolds, P.H., and Hamilton, M.A. 2014. Structural setting and age of the Partridge Island block, southern New Brunswick, Canada: a link to the Cobequid highlands of northern mainland Nova Scotia. *Canadian Journal of Earth Sciences*, 51, pp. 1–24. <http://dx.doi.org/10.1139/cjes-2013-0120>
- Pearce, J.A. 1982. Trace element characteristics of lavas from destructive plate boundaries. *In* *Andesites*. *Edited by* R.S. Thorpe, Wiley, Chichester, pp. 525–548.
- Pearce, J.A. and Cann, J.R. 1973. Tectonic setting of basic volcanic rocks determined using trace element analysis. *Earth and Planetary Science Letters*, 19, pp. 290–300. [http://dx.doi.org/10.1016/0012-821X\(73\)90129-5](http://dx.doi.org/10.1016/0012-821X(73)90129-5)
- Pearce, J.A. and Gale, J.R. 1977. Identification of ore-deposition environment from trace element geochemistry of associated igneous host rocks. *Geological Society Special Paper* 7, pp. 14–24. <http://dx.doi.org/10.1144/GSL.SP.1977.007.01.03>
- Pearce, J.A. and Norry, M.J. 1979. Petrogenetic implications of Ti, Zr, Y, and Nb variations in volcanic rocks. *Contributions to Mineralogy and Petrology*, 69, pp. 33–47. <http://dx.doi.org/10.1007/BF00375192>
- Pearce, J.A., Harris, N.B.W., and Tindle, A.G. 1984. Trace element discrimination diagrams for the tectonic interpretation of granitic rocks. *Journal of Petrology*, 25, pp. 956–983. <http://dx.doi.org/10.1093/petrology/25.4.956>
- Pegram, W.J. 1990. Development of continental lithospheric mantle as reflected in the chemistry of the Mesozoic Appalachian tholeiites, U.S.A. *Earth and Planetary Science Letters*, 97, pp. 316–331. [http://dx.doi.org/10.1016/0012-821X\(90\)90049-4](http://dx.doi.org/10.1016/0012-821X(90)90049-4)
- Pe-Piper, G. and Piper, D.J.W. 1998. Geochemical evolution of Devonian-Carboniferous igneous rocks of the Magdalen basin, Eastern Canada: Pb- and Nd-isotopic evidence for mantle and lower crustal sources. *Canadian Journal of Earth Sciences*, 35, pp. 201–221. <http://dx.doi.org/10.1139/e97-106>
- Pe-Piper, G. and Piper, D.J.W. 2002. A synopsis of the geology of the Cobequid Highlands, Nova Scotia. *Atlantic Geology*, 38, pp. 145–160.
- Pe-Piper, G., Cormier, R.F., and Piper, D.J.W. 1989. The age and significance of Carboniferous plutons of the western Cobequid Highlands, Nova Scotia. *Canadian Journal of Earth Sciences*, 26, pp. 1297–1307. <http://dx.doi.org/10.1139/e89-109>
- Pe-Piper, G., Piper, D.J.W., and Clerk, S.B. 1991. Persistent mafic igneous activity in an A-type granite pluton, Cobequid Highlands, Nova Scotia. *Canadian Journal of Earth Sciences*, 28, pp. 1058–1072. <http://dx.doi.org/10.1139/e91-096>
- Piper, J.W.D., Pe-Piper, G., and Pass, D. 1996. The stratigraphy and geochemistry of late Devonian to early Carboniferous volcanic rocks of the northern Chignecto peninsula, Cobequid Highlands, Nova Scotia. *Atlantic Geology*, 32, pp. 39–52.
- Ramezani, J., Schitz, M.D., Davydov, V.I., Bowring, S.A., and Snyder, W.S. 2007. High-precision U-Pb zircon age constrains on the Carboniferous - Permian boundary in the southern Urals stratotype. *Earth and Planetary Science Letters*, 256, pp. 244–257. <http://dx.doi.org/10.1016/j.epsl.2007.01.032>
- Rollinson, H. 1993. *Using Geochemical Data: Evaluation, Presentation, Interpretation*. Prentice Hall, Harlow, England, 352 p.
- Satkoski, A., Barr, S., and Samson, S. 2007. Sm-Nd isotopic and whole rock chemical compositions of Late Neoproterozoic and Cambrian sedimentary and metasedimentary rocks of the Caledonian highlands (Avalonia), southern New Brunswick. *Geological Society of America Abstracts with Programs*, 39, p. 95.
- Samson, S.D., Barr, S.M., and White, C.E. 2000. Nd isotopic characteristics of terranes within the Avalon Zone, southern New Brunswick. *Canadian Journal of Earth Sciences*, 37, pp. 1039–1052. <http://dx.doi.org/10.1139/e00-015>
- Skehan, J.W. and Rast, N. 1995. Late Proterozoic to Cambrian evolution of the Boston Avalon terrane. *In* *Current perspectives in the Appalachian-Caledonian Orogen*. *Edited by* J.P. Hibbard, C.R. van Staal, and P.A. Cawood. Geological Association of Canada Special Paper 41, pp. 207–225.
- Socci, A.D., Skehan, J.W., and Smith, G.W. 1990. Geology of the composite Avalon terrane of southern New England. *Geological Society of America Special Paper* 245, 254 p.

- Steiger, R.H. and Jager, E. 1977. Subcommittee on geochronology: convention of the use of decay constants in geo- and cosmochronology. *Earth and Planetary Science Letters*, 36, pp. 359–362. [http://dx.doi.org/10.1016/0012-821X\(77\)90060-7](http://dx.doi.org/10.1016/0012-821X(77)90060-7)
- Stern, R.J. 2002. Crustal evolution in East African orogen: a neodymium isotopic perspective. *Journal of African Earth Sciences*, 34, pp. 109–117. [http://dx.doi.org/10.1016/S0899-5362\(02\)00012-X](http://dx.doi.org/10.1016/S0899-5362(02)00012-X)
- Swanson, M.T. 1988. Pseudotachylite-bearing strike-slip duplex structures in the Fort Foster brittle zone, southern Maine. *Journal of Structural Geology*, 10, pp. 813–828. [http://dx.doi.org/10.1016/0191-8141\(88\)90097-1](http://dx.doi.org/10.1016/0191-8141(88)90097-1)
- Swanson, M. T. 1995. Distributed ductile dextral shear strain throughout the Casco Bay area. *In* Guidebook to field trips in southern Maine and adjacent New Hampshire. *Edited by* A.M. Hussey II and R.A. Johnston. New England Intercollegiate Geological Conference, 85th Annual Meeting, pp. 1–13.
- Swanson, M.T. 2007. Structure of Late Paleozoic brittle dextral strike-slip faults in coastal Maine exposures. *In* Guidebook to field trips in New Hampshire, adjacent Maine and Massachusetts. *Edited by* P. Thompson. Northeast Section, Geological Society of America, Durham, New Hampshire, Trip F-2, pp. 3–18.
- Swanson, M.T. 1999. Dextral transpression at the Casco Bay restraining bend, Norumbega fault zone, coastal Maine. *Geological Society of America Special Paper* 331. Pp. 85–104.
- Tanaka, T., Togashi, S., Kamioka, H., Amakawa, H., Kagami, H., Hamamoto, T., Yuhara, M., Orihashi, Y., Yoneda, S., Shimizu, H., Kunimaru, T., Takahashi, K., Yanagi, T., Nakano, T., Fujimaki, H., Shinjo, R., Asahara, Y., Tanimizu, M., and Dragusanu, C. 2000. JNdi-1; a neodymium isotopic reference in consistency with LaJolla neodymium. *Chemical Geology*, 168, pp. 279–281. [http://dx.doi.org/10.1016/S0009-2541\(00\)00198-4](http://dx.doi.org/10.1016/S0009-2541(00)00198-4)
- Thompson, M.D. and Hermes, O.D. 2003. Early rifting in the Narragansett basin, Massachusetts-Rhode Island: Evidence from Late Devonian bimodal volcanic rocks. *Journal of Geology*, 111, pp. 597–604. <http://dx.doi.org/10.1086/376768>
- Thompson, M.D., Hermes, O.D., Bowring, S.A., Isachsen, C.E., Besancon, J.R., and Kelly, K.L. 1996. Tectonostratigraphic implications of Late Proterozoic U-Pb zircon ages in the Avalon zone of southeastern New England. *In* Avalonian and Related Peri-Gondwanan Terranes of the Circum-North Atlantic. *Edited by* R.D. Nance and M.A. Thompson. Geological Society of America Special Paper 304, pp. 179–191. <http://dx.doi.org/10.1130/0-8137-2304-3.179>
- Thompson, M.D., Barr, S.M., and Grunow, A.M. 2012. Avalonian perspectives on Neoproterozoic paleogeography: Evidence from Sm-Nd isotope geochemistry and detrital zircon geochronology in SE New England, USA. *Geological Society of America Bulletin*, 124, pp. 517–531. <http://dx.doi.org/10.1130/B30529.1>
- Tomascak, P.B. Brown, M., Solar, G.S., Becker, H.J., Centorbi, T.L., and Tian, J. 2005. Source contributions to Devonian granite magmatism near the Laurentian border, New Hampshire and western Maine, USA. *Lithos*, 80, pp. 75–90. <http://dx.doi.org/10.1016/j.lithos.2004.04.059>
- van Staal, C. R. and Barr, S. 2012. Lithospheric Architecture and Tectonic Evolution of the Canadian Appalachians and Associated Atlantic Margin. *In* Tectonic Styles in Canada: the LITHOPROBE Perspective. *Edited by* J. Percival, F. Cook, and R.M. Clowes. Geological Association of Canada Special Paper 49, pp. 41–95.
- van Staal, C. R. and Zagorevski, A. 2012. Accreted terranes of the Appalachian orogen in Newfoundland: in the footsteps of Hank Williams. Field trip Guidebook –A1, Geological Association of Canada, Mineralogical Association of Canada joint meeting, 99 pp.
- van Staal, C., Sullivan, R.W., and Whalen, J.B. 1996. Provenance and tectonic history of the Gander margin in the Caledonian/Appalachian Orogen: implications for the origin and assembly of Avalonia. *In* Avalonian and Related Peri-Gondwanan Terranes of the Circum-North Atlantic. *Edited by* R.D. Nance and M.A. Thompson. Geological Society of America Special Paper 304, pp. 347–367. <http://dx.doi.org/10.1130/0-8137-2304-3.347>
- van Staal, C., Dewey, J.F., MacNiocail, C., and McKerrow, S. 1998. The Cambrian-Silurian tectonic evolution of the northern Appalachians: history of a complex, southwest Pacific-type segment of Iapetus. *In* Lyell: the Past is the Key to the Present. *Edited by* D.L. Blundell and A.C. Scott. Geological Society Special Publication 143, pp. 199–242.
- van Staal, C., Whalen, J.B., Valverde-Vaquero, P., Zagorevski, A., and Rogers, N. 2009. Pre-Carboniferous, episodic accretion-related, orogenesis along the Laurentian margin of the northern Appalachians. *In* Ancient Orogens and Modern Analogues. *Edited by* J.B. Murphy, J.D. Kerppe, and A.J. Hynes. Geological Society, London, Special Publications 327, pp. 271–316.
- van Staal, C., Barr, S.M., and Murphy, J.B. 2012. Provenance and tectonic evolution of Ganderia: Constraints on the evolution of the Iapetus and Rheic oceans. *Geology*, 40, pp. 987–990. <http://dx.doi.org/10.1130/G33302.1>
- Walsh, G.J., Aleinikoff, J.N., and Wintsch, R.P. 2007. Origin of the Lyme dome and implications for the timing of multiple Alleghanian deformational and intrusive events in southern Connecticut. *American Journal of Science*, 307, pp. 168–215. <http://dx.doi.org/10.2475/06.2007.06>
- Wandke, A. 1922. Intrusive rocks of the Portsmouth basin, Maine and New Hampshire. *American Journal of Science*, 4, pp. 139–158. <http://dx.doi.org/10.2475/ajs.s5-4.20.139>
- Wintsch, R.P., Aleinikoff, J.N., Walsh, G.J., Bothner, W.A., Hussey, A.M. II, and Fanning, C.M. 2007. SHRIMP U-Pb evidence for a Late Silurian age of metasedimentary rocks in the Merrimack and Putnum-Nashoba terranes, eastern New England. *American Journal of Science*, 307, pp. 119–167. <http://dx.doi.org/10.2475/01.2007.05>

Yaowanoyothin, W. and Barr, S.M. 1991. Petrology of the Black Brook granitic suite, Cape Breton Island, Nova Scotia. *Canadian Mineralogist*, 29, pp. 499–515.

Zenk, M. and Schulz, B. 2004. Zoned Ca-amphiboles and related P-T evolution in metabasites from the classical Barrovian metamorphic zones in Scotland. *Mineralogical Magazine*, 68, pp. 769–786. <http://dx.doi.org/10.1180/0026461046850218>

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