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Holocene emergence and shoreline delevelling, southern Eureka Sound, High Arctic Canada

L'émersion des terres et l'évolution du littoral à l'Holocène, dans la partie sud de l'Eureka Sound, Haut Arctique canadien

Auftauchen und Entwicklung der Küsten, südlicher Eureka Sound, kanadische Hocharktis

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Résumé de l'article

Cet article présente une reconstitution des changements du niveau marin relatif et de ces conséquences sur le littoral au postglaciaire. Les deltas marins soulevés, les plages et les limites de l'érosion par les vagues qui datent du début de l'Holocène attestent d'une émersion des terres de plus de 150 m dans la région. La limite marine est métachrone et s'est établie successivement au retrait glaciaire. Des contrastes forts dans les courbes du niveau marin relatif et le taux d'émrersion inital ont été enregistrés dans la région à l'étude. Dans le Blind Fiord, le niveau marin relatif s'est abaissé de façon continue après la déglaciation. L'émersion initiale s'élevait à ⩾ 5 m/siècle. Ce taux contraste avec les courbes de Starfish Bay et de Irene Bay, où le taux d'émersion initial était de ⩽ 1 m/siècle. Les iso- bases tracées sur le littoral de 8,5 ka au Eureka Sound montrent qu'une zone d'émersion plus élevée (⩾ 130 m) s'étend le long du chenal et se termine dans le voisinage de l'entrée de la Norwegian Bay. Ce modèle démontre qu'il y a eu une charge glaciaire distincte au-dessus du Eureka Sound durant le dernier maximum glaciaire, ce qui a aussi été démontré par l'entremise d'autres éléments de preuves d'ordre géologique.

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HOLOCENE EMERGENCE AND SHORELINE DELEVELLING, SOUTHERN EUREKA SOUND, HIGH ARCTIC CANADA

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ABSTRACT This paper is a reconstruction of postglacial relative sea level change and shoreline delevelling in southern Eureka Sound, High Arctic Canada. Postglacial emergence of up to 150 m is recorded in this area by raised marine deltas, beaches and washing limits that date from the early Holocene. Marine limit is metachronous and formed successively with glacier retreat. Marked contrasts in the form of relative sea level curves and rate of initial emergence are recorded from the study area. In Blind Fiord, relative sea level fell continuously following deglaciation. Initial emergence was characterised by rates of ≥5 m/century. This contrasts with curves from Starfish and Irene bays, where the rate of initial emergence was ≤1 m/century. Isobases drawn on the 8.5 ka shoreline for greater Eureka Sound demonstrate that a cell of highest emergence (≥130 m asl) extends along the length of the channel, and closes in the vicinity of the entrance to Norwegian Bay. This pattern confirms a distinct loading centre over Eureka Sound during the Last Glacial Maximum, and is compatible with independent glacial geological evidence indicating that the thickest ice was centred over the channel during the Late Wisconsinan.

RÉSUMÉ *L'émersion des terres et l'évolution du littoral à l'Holocène, dans la partie sud de l'Eureka Sound, Haut Arctique canadien.* Cet article présente une reconstitution des changements du niveau marin relatif et de ces conséquences sur le littoral au postglaciaire. Les deltas marins soulevés, les plages et les limites de l'érosion par les vagues qui datent du début de l'Holocène attestent d'une émersion des terres de plus de 150 m dans la région. La limite marine est métachrone et s'est établie successivement au retrait glaciaire. Des contrastes forts dans les courbes du niveau marin relatif et le taux d'émrersion inital ont été enregistrés dans la région à l'étude. Dans le Blind Fiord, le niveau marin relatif s'est abaissé de façon continue après la déglaciation. L'émersion initiale s'élevait à ≥ 5 m/siècle. Ce taux contraste avec les courbes de Starfish Bay et de Irene Bay, où le taux d'émersion initial était de ≤ 1 m/siècle. Les isobases tracées sur le littoral de 8,5 ka au Eureka Sound montrent qu'une zone d'émersion plus élevée (≥ 130 m) s'étend le long du chenal et se termine dans le voisinage de l'entrée de la Norwegian Bay. Ce modèle démontre qu'il y a eu une charge glaciaire distincte au-dessus du Eureka Sound durant le dernier maximum glaciaire, ce qui a aussi été démontré par l'entremise d'autres éléments de preuves d'ordre géologique.

ZUSAMMENFASSUNG *Auftauchen und Entwicklung der Küsten, südlicher Eureka Sound, kanadische Hocharktis.* Dieser Artikel präsentiert eine Rekonstruktion der Wechsel des postglazialen relativen Meeresniveaus und ihrer Auswirkungen auf die Küste im südlichen Eureka Sound, kanadische Hocharktis. Ein postglaziales Auftauchen bis zu 150 m wird in diesem Gebiet durch angehobene Meeresdeltas, Strände und Verwaschungsgrenzen dokumentiert, die aus dem frühen Holozän stammen. Die Meeresgrenze ist metachron und wurde allmählich während des Gletscherrückzugs gebildet. Starke Kontraste wurden in der Form der relativen Meeresniveaukurven und in der ursprünglichen Auftachrate im untersuchten Gebiet festgestellt. In Blind Fiord sank das relative Meeresniveau kontinuierlich nach der Enteisung. Das ursprüngliche Auftauchen geschah in Raten von ≥5 m/Jahrhundert. Dies kontrastiert mit Kurven von Starfish Bay und Irene Bay, wo die ursprüngliche Auftauchrate ≤1 m/Jahrhundert betrug. Die auf der 8.5 ka Küstenlinie gezogenen Isobasen im Eureka Sound zeigen, dass eine Zone höheren Auftauchens (≥130 m) sich über die Länge der Rinne erstreckt und in der Nähe des Eingangs zur Norwegian Bay endet. Dieses Modell bestätigt ein getrenntes Ladungszentrum über Eureka Sound während des letzten glazialen Maximums, und verträgt sich auch mit unabhängigen glazialen geologischen Anhaltspunkten, die zeigen, dass das dickste Eis sich über der Rinne während des späten Wisconsin konzentrierte.

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INTRODUCTION

This paper discusses the postglacial relative sea level history of southern Eureka Sound, High Arctic Canada (Fig. 1), focussing on initial emergence, pattern of shoreline delevelling and implications for former glacier loading. Blake (1970) proposed the existence of the Innuitian Ice Sheet in the Canadian High Arctic during the Late Wisconsinan on the basis of the pattern of differential postglacial rebound since 5 ka BP. He demonstrated that shorelines of this age in the Queen Elizabeth Islands were highest (>25 m asl) throughout a broad northeast/southwest oriented corridor extending from northern Eureka Sound to Bathurst Island, and he proposed that this emergence reflected a regional ice sheet over the Queen Elizabeth Islands during the Last Glacial Maximum (*cf.* Walcott, 1972; Tushingham, 1991). In contrast, a markedly different reconstruction, of a restricted Late Wisconsinan glaciation for the same region, was proposed by England (1976a and b) on the basis of glacial geologic data from northeastern Ellesmere Island and an alternative interpretation of the postglacial emergence. These contrasting reconstructions formed the end members in an ensuing debate concerning the extent of ice during the Last Glacial Maximum in the Queen Elizabeth Islands (*e.g.*, Blake, 1992a and b, 1993; Blake *et al.,* 1992; de Freitas, 1990; Tushingham, 1991; England, 1987, 1990, 1996, England *et al.*, 1991; Lemmen, 1989; Bell, 1996).

Previous work on the postglacial emergence of northern Eureka Sound and Greely Fiord reported evidence for a period of stable relative sea level at marine limit (England, 1992). Similar sea level curves were presented for northeastern and eastern Ellesmere Island which show either an interval of sea level arrest subsequent to marine limit formation (England, 1983), or slow (1-2 m/century) initial emergence (England, 1997; see also Retelle *et al.*, 1989). These curves contrast markedly with others reported from elsewhere in the region (Blake, 1975, 1992a; Hodgson *et al.*, 1991; Lemmen *et al.*, 1994; Bednarski, 1995; Dyke, 1998) which show most rapid initial emergence (4.5-7 m/century), and are thus similar to curves from areas of the Canadian Arctic which were formerly covered by the Laurentide Ice Sheet (Andrews, 1970; Dyke, 1984; Dyke *et al.,* 1991).

Isobases drawn by England (1992, 1997) on the 8 ka shoreline in Greely Fiord, Ellesmere Island (Fig. 1), exhibit a narrow plunging ridge of maximum emergence, parallel to the regional geological structure. England contrasted this isobase pattern with the broad cells of uplift documented elsewhere in Arctic Canada which record postglacial unloading following removal of the Laurentide Ice Sheet (*cf.* Andrews, 1970; Dyke, 1984). He also argued that the isobase ridge was difficult to reconcile with a soley glacioisostatic explanation, as it did not conform to the distribution of the last ice load in the region which was inferred to be restricted (England, 1987, 1990; Bell, 1992, 1996). He therefore proposed a possible neotectonic contribution to Holocene emergence for western Ellesmere Island (England, 1992, 1997).

Both the isobase pattern and slow rate of initial emergence, coupled to independent glacial geological evidence suggesting a restricted ice-cover during the Last Glacial Maximum, were interpreted as being incompatible with a Late Wisconsinan Innuitian Ice Sheet (England, 1983, 1992, 1997; Bell, 1996). However, more recent fieldwork validates an extensive Late Wisconsinan glacier cover for at least the eastern and southeastern sectors of the Queen Elizabeth Islands (Ellesmere, Axel Heiberg and Devon islands) (Hättestrand and Stroeven, 1996; Bednarski, 1998; England, 1998, 1999; Ó Cofaigh *et al.*, 1998; Ó Cofaigh, 1999; Dyke, 1999; see also Funder, 1989; Funder and Hansen, 1996). Integration of this new glacial geologic evidence with the associated postglacial relative sea level histories of these areas has only recently commenced (Dyke, 1998; England and Ó Cofaigh, 1998).

Glacial geologic and chronologic evidence indicates that Late Wisconsinan glaciation in southern Eureka Sound was characterised by an extensive ice cover (Ó Cofaigh, 1998, 1999; Ó Cofaigh *et al.*, 1998, in press). This paper presents the postglacial relative sea level history associated with removal of that ice load, and it provides new data on the magnitude, timing and pattern of postglacial emergence (Fig. 2). It has three principal objectives: (1) to reconstruct initial postglacial emergence at several sites where the best chronological control is available; (2) to reconstruct the pattern of shoreline delevelling in southern Eureka Sound and to link this data with previously published work to the north and south (England, 1976b, 1992; Bell, 1996; Dyke, 1998, 1999); and (3) to assess implications for former ice sheet loading in the region.

STUDY AREA

Eureka Sound is the inter-island channel, 300 km long and 10-28 km wide, which separates Ellesmere and Axel Heiberg islands (Figs. 1 and 2). Geologically, the study area is dominated by north-northeast striking sedimentary rocks, although igneous rocks outcrop locally (Trettin, 1991). The geological structure dictates a topographic grain of ridges and valleys. Uplands reaching >1000 m asl are dissected by steep-sided fiords and valleys aligned both parallel to bedrock structure (*e.g.*, Blind Fiord) and cross-cutting it (*e.g.*, Bay Fiord). Contemporary glaciers are limited to small, upland ice-caps, although the region is bordered immediately to the east and west by extensive ice-fields on central Axel Heiberg and Ellesmere islands (Fig. 2).

LATE WISCONSINAN GLACIATION OF SOUTHERN EUREKA SOUND

During the Late Wisconsinan, southern Eureka Sound supported extensive glaciation, consisting of expanded ice-caps which were coalescent along the length of the channel. Ice-divides were located along the highlands of central Ellesmere and Axel Heiberg islands, from which ice flowed east and west into Eureka Sound, with development of preferential flow along the axes of major fiords (Ó Cofaigh, 1998, 1999; Ó Cofaigh *et al.*, 1998, in press). In Eureka Sound, trunk ice flowed north towards Nansen Sound (*cf.* Fyles, *in* Jenness 1962; Bell, 1992; Bednarski, 1998) and south towards Norwegian Bay. Raanes Peninsula supported a local ice-dome which was coalescent with trunk ice in Eureka Sound. Deglaciation of southern Eureka Sound commenced ≥9.2 ka BP

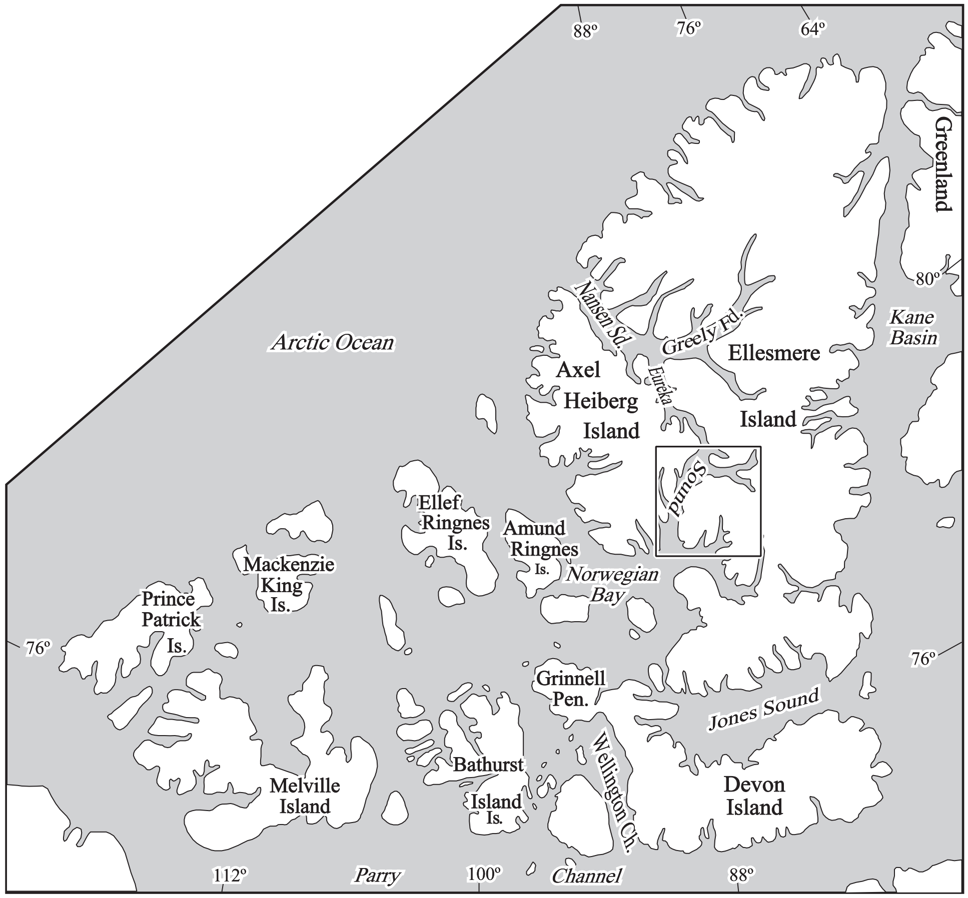


FIGURE 1. Queen Elizabeth Islands, Canada, and location of the study area.

Localisation de l'archipel de la Reine-Elisabeth et de la région à l'étude.

[9.9 ka calender years BP] (Ó Cofaigh, 1998, 1999) and was characterised by initial break-up of ice in the channel with subsequent retreat east and west to the former ice-divides. Thus marine limit throughout the study area is time-transgressive and records sequential entry of the sea with ice retreat.

METHODOLOGY

SURVEYING TECHNIQUE AND DEFINITION OF MARINE LIMIT

The altitude of raised marine features was determined using a Wallace and Tiernan micro-altimeter (accuracy ±2 m). Readings were corrected for fluctuations in atmospheric pressure and site specific temperature. High tide level, commonly demarcated by a well-defined kelp line, was used as the reference datum for sea level. Radiocarbon dates on marine shells and driftwood provide chronological control on the establishment of marine limit and subsequent emergence.

Marine limit is the maximum elevation attained by the sea along a glacioisostatically depressed coastline. Its elevation at a site reflects distance from the former ice margin (which is an indication of ice thickness over the site), date of deglaciation and eustatic sea level rise (Andrews, 1970). Throughout the study area, marine limit was taken to be either: (1) the highest raised marine delta or beach; or (2) the lowest undisturbed till or felsenmeer (washing limits) as commonly marked by a notch cut in till with a well sorted sediment veneer or bedrock below, or by an abrupt textural transition between poorly sorted till/felsenmeer and sorted sediment below; or (3) the highest elevation at which well preserved marine shells were found, which provides a minimum estimate on marine limit.

MARINE LIMIT: ELEVATION AND PATTERN

The highest marine limit observed in the study area occurs on the north coast of Stor Island at 145-151 m asl (Fig. 3). Marine limits of >140 m asl also occur along the south coast of Raanes Peninsula between Eureka Sound and Trold Fiord (Fig. 3). Blind Fiord, Trold Fiord, Starfish Bay and Jaeger Bay all exhibit a progressive decline in marine limit from mouth to head. For example, in Trold Fiord, marine limit falls from 143 m asl at the mouth to 98 m asl at the head, whereas in Starfish Bay marine limit decreases from 113 m to 80 m asl.

Along Eureka Sound, north of Hare Bay, marine limit is recorded by deltas at the mouths of several valleys. In inner Trapper's Cove, ice-contact deltas grade to relative sea levels at 118-120 m asl (Fig. 3). This contrasts with outer Trapper's Cove and the Eureka Sound coast, where marine limit is defined by poorly-preserved raised beaches at 83 m asl (minimum), and deltas immediately to the south at 99 m asl (Fig. 3). The north coast of Raanes Peninsula is characterised by variable marine limit elevations which range from 76 to 120 m asl in Eureka Sound and central Bay Fiord, respectively, before

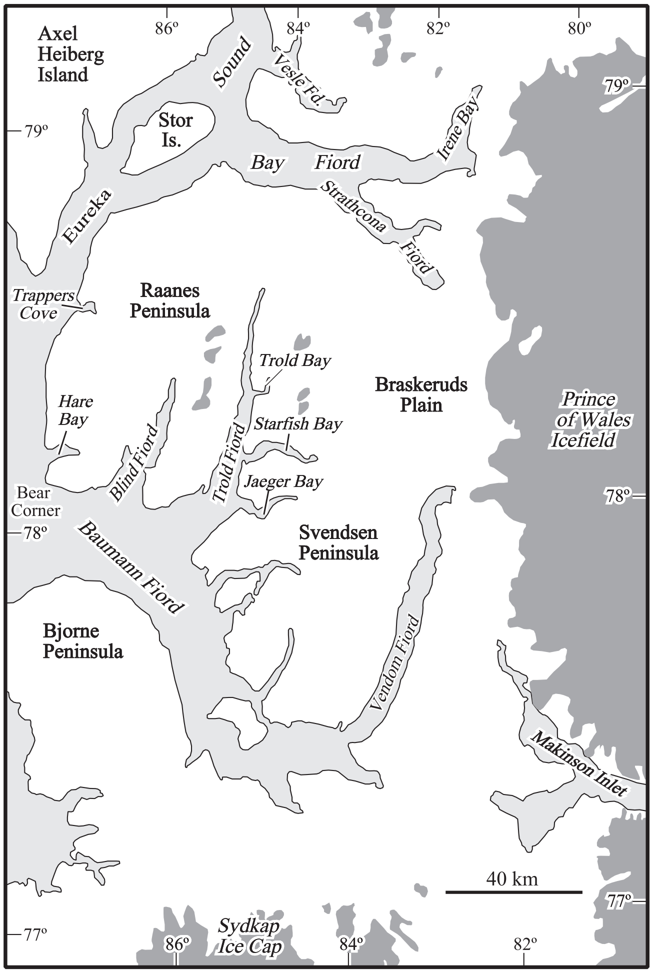


FIGURE 2. Southern Eureka Sound with contemporary ice-cover (dark shading).

Le couvert de glace actuel sur la partie sud du Eureka Sound (trame plus foncée).

falling progressively to 65-87 m asl at the heads of Strathcona Fiord and Irene Bay (Fig. 3). Regionally, therefore, marine limit exhibits an overall decrease in elevation eastwards from Eureka Sound to the fiord heads. However, this decrease is variable over short distances, a pattern inferred to reflect the metachronous age of marine limit occasioned by ice retreat.

RELATIVE SEA LEVEL CURVES

Emergence data are presented for three sites, Blind Fiord, Starfish Bay and Irene Bay (Fig. 4). At each site, the elevation and age of radiocarbon-dated samples and their associated relative sea levels are given. Figure 5 shows relative sea level curves for the three sites. All radiocarbon dates (including calibrated ages, reported as “cal BP”[1]) are listed in Table I.

1. All dates were calibrated using CALIB 3.0 (Stuiver and Reimer, 1993), and the date range reported here is that which yields 100 % probability when 2σ is used.

The three sites were inundated by ice during the Last Glacial Maximum (Ó Cofaigh, 1998, 1999). Blind Fiord was fed by ice from a local dome centred over Raanes Peninsula. This local ice-dome was coalescent with westerly-flowing, regional ice emanating from a divide under the present-day Prince of Wales Icefield. Starfish and Irene bays were infilled by this regional ice.

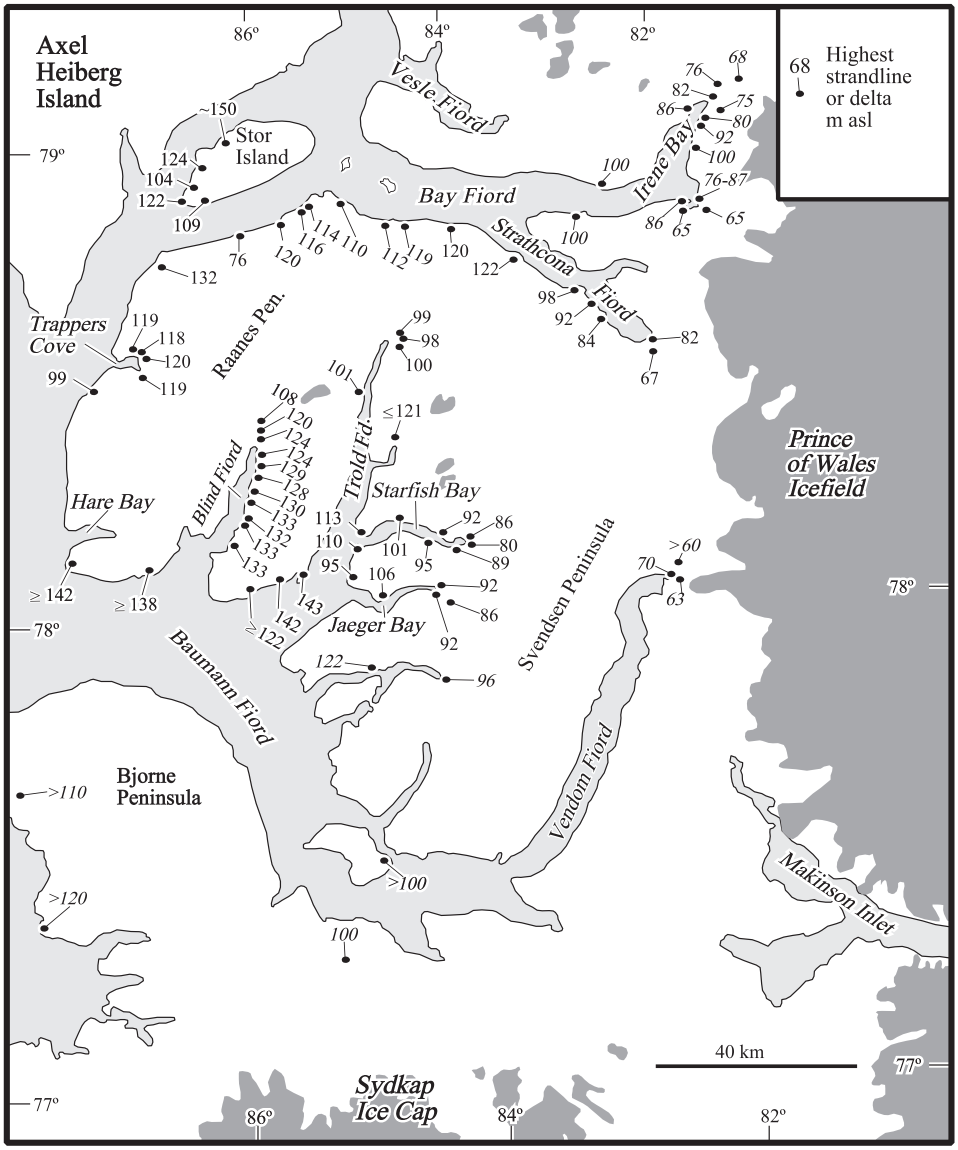
BLIND FIORD

At the mouth of Blind Fiord, marine limit is defined by the uppermost raised beach at 138 m asl. A fragment of *Mya truncata* collected from a beach surface at 128 m asl yielded an Accelerator Mass Spectrometry (AMS) date of 8590±70 BP [9460-9060 cal BP] (TO-5862; Site 1, Fig. 4 and Table I), and provides a minimum age estimate on the 138 m marine limit. In the central fiord, marine limit falls to 133 m asl and two samples provide minimum dates on its formation. A surface fragment of *M. truncata* from a raised beach at 127 m asl dated 8510±80 BP [9400-8970 cal BP](TO-5612; Site 2, Fig. 4 and Table I). Immediately up-fiord, whole valves and fragments of *Hiatella arctica* and *M. truncata* from a raised beach at 119 m asl dated 8550±80 BP [9420-9000 cal BP](GSC-6047; Site 3, Fig. 4 and Table I). Both dates provide minimum age estimates for the 133 m asl marine limit. The standard errors of these three dates overlap and thus indicate rapid ice-retreat and formation of marine limit through the outer and central fiord. The dates also indicate that at least 5 m of emergence occurred in <100 (^{14}C and calendar years) years assuming that the samples relate to their respective marine limits (138 m and 133 m). If the samples date relative sea levels at their elevations, then emergence could have been as much as 9 m in <100 years (sample elevations at 128 and 119 m asl). Although both GSC-6047 and TO-5612 yielded similar ages for the 133 m marine limit in central Blind Fiord, the former is a bulk date, and thus could contain a mixture of different aged shells. If this sample is excluded from the reconstruction and the emergence rate based on the two AMS dates (TO-5862 and TO-5612), initial emergence is still 5 m in <100 years.

Control on subsequent emergence in Blind Fiord is provided by four dates. In the inner fiord, marine limit is defined by a gravel beach berm at 129 m asl (Fig. 4). A surface fragment of *H. arctica* from a raised beach at 123 m asl dated 8310±80 BP [9210-8630 cal BP] (TO-5608; Site 4, Fig. 4 and Table I) and provides a minimum age estimate on the 129 m marine limit. North of the fiord head, whole valves and fragments of *H. arctica* and *M. truncata* from 107 m asl dated 8220±100 BP [9100-8470 cal BP] (GSC-6054; Site 5, Fig. 4 and Table I), which is also a minimum date on local marine limit at 124 m asl. Up-fiord of this site, a marine limit delta at 120 m asl occurs at the mouth of a lateral meltwater channel. Single valves and fragments dominated by *M. truncata* collected at 95 m asl on the delta foreslope gave a radiocarbon date of 8090 ±110 BP [8950-8370 cal BP] (GSC-5896; Site 6, Fig. 4 and Table I), which is a minimum age for the 120 m marine limit. Finally, paired valves of *Astarte borealis* and *H. arctica* from 31 m asl in silt immediately underlying a delta at 39 m asl in the central fiord dated 5640±110 [6360-5870 cal

FIGURE 3. Marine limit elevations (m asl) marked by the uppermost delta, washing limit or raised beach, southern Eureka Sound. Italicized marine limit elevations are from Hodgson (1985).

Les limites du niveau marin (en m) telles qu'enregistrées par les deltas les plus élevés, par les limites de l'érosion par les vagues ou par les plages soulevées, dans la partie sud de l'Eureka Sound. Les limites du niveau marin données en italique sont de Hodgson (1985).



BP] (GSC-6102; Site 7, Fig. 4 and Table I). It is important to note that although sites at the fiord head are separated from sites in the outer fiord by a distance of ~38 km, the fiord *parallels* the regional isobases at 8.5 ka BP (see Fig. 6 below), and hence differential postglacial rebound does not compromise treating the dates as a single relative sea level curve. The relative sea level curve for Blind Fiord (Fig. 5A) demonstrates continuous emergence from 8.6 ka BP to present. Initial emergence was ≥5 m/century.

STARFISH BAY

Deglaciation of Starfish Bay followed the retreat of ice from outer Trold Fiord. At the mouth of Starfish Bay, ice-contact deltas occur at 113 m asl (Fig. 3) and mark a stillstand during retreat of ice into the outer part of the bay. Marine limit drops to 101 m asl along the north shore of the bay where it is defined by a prominent washing limit. A sample of whole valves of *H. arctica* and *M. truncata* collected from glaciomarine silt at 78 m asl dated 7740±90 BP [8410-8000 cal BP] (GSC-6037; Site 8, Fig. 4 and Table I). This is a minimum age for deglaciation and marine limit.

Marine limit in the inner fiord is marked by a bench cut in till at 89 m asl (Fig. 4). At the fiord head, well-developed deltas fronted by extensive glaciomarine silt occur at 86 and 80 m asl (Fig. 4). These deposits mark a major stillstand of the trunk glacier during deglaciation. A sample of *Portlandia arctica* collected from silt at 68 m asl, 3 km west of the fiord head, dated 8710±120 BP [9730-9060 cal BP] (GSC-2719; Site 9, Fig. 4 and Table I) (Hodgson, 1985). Hodgson inferred that this date provided an approximate age for the fiord head deltas. However, because the sample site occurs 3 km west of the deltas, the validity of this inferred relationship is uncertain, and the shells could alternatively correlate with the 89 m bench in the inner fiord. A final sample consisting of a single valve of *H. arctica,* was collected from silt at 71 m asl east (inland) of these deltas. This dated 7240±80

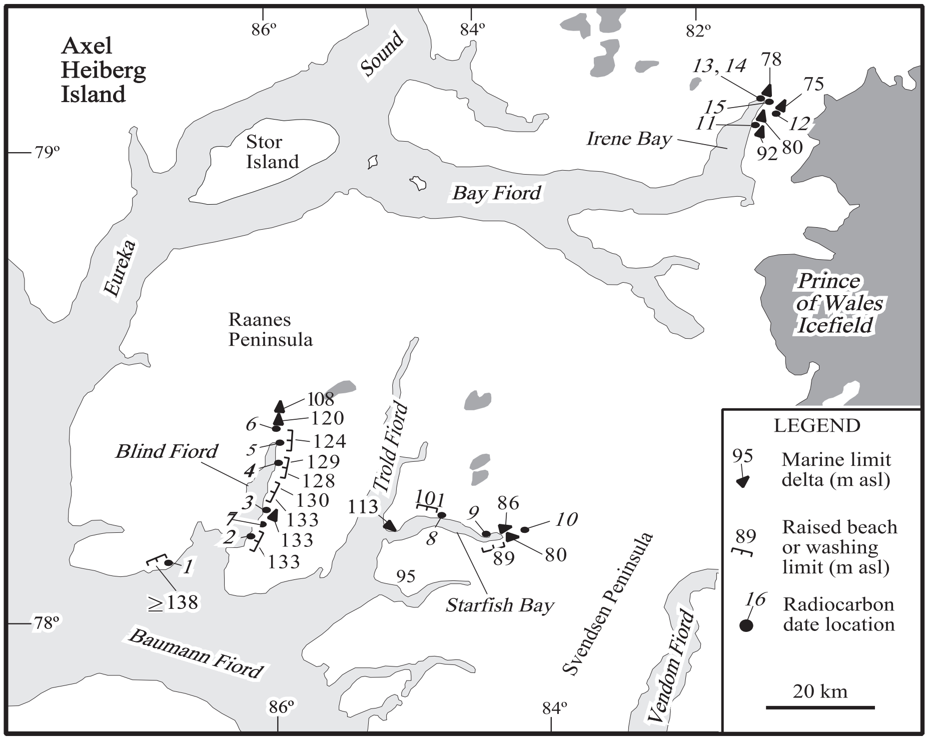


FIGURE 4. Marine limit landforms, elevations (m asl) and location of radiocarbon dates in Blind Fiord, Starfish Bay and Irene Bay as discussed in text. Corresponding site numbers of radiocarbon dates are listed in Table I.

Les formes relatives à la limite du niveau marin, les altitudes et la localisation des datations au radiocarbone du Blind Fiord et des Starfish Bay et Irene Bay. Les numéros de sites des datations sont donnés au tableau 1

BP [7910-7570 cal BP] (TO-5596; Site 10, Fig. 4, Table I) and provides a minimum age for deglaciation and formation of the fiord head deltas.

Therefore, at the fiord head between 8710±120 BP [9730-9060 cal BP] and 7240±80 BP [7910-7570 cal BP], sea level could have fallen by as little as 9 m, based on the assumption that the 8.7 ka BP [9730-9060 cal BP] date relates to the 89 m marine limit in the inner fiord, and the 7.2 ka BP [7910-7570 cal BP] date relates to a relative sea level at 80 m asl (Fig. 5B, “minimum”). This is equivalent to an emergence rate of only 0.6 m/century (based on the calibrated ages). If the 7.2 ka BP [7910-7570 cal BP] shells are related to a relative sea level at 71 m asl (the sample elevation), then emergence increases to 18 m (1 m/century) (Fig. 5B, “maximum”).

IRENE BAY

Prominent marine limit deltas and thick raised marine silt record ice-marginal stabilisation and deposition at the fiord head during deglaciation. Hodgson (1985) reported several radiocarbon dates from this area and these are discussed below. Ice-contact deltas on the south side of inner Irene Bay are graded to relative sea levels of 80 and 92 m asl (Figs. 3 and 4). Whole valves of *P. arctica* were collected at 70-74 m asl from glaciomarine rhythmites capped by the 80 m delta. This sample dated 8820±90 BP [9810-9340 cal BP] (GSC-1978; Site 11, Fig. 4 and Table I). Immediately east of this site, ice-contact deltas with thick pro-delta silt grade to 75 m asl. Whole valves of *H. arctica* and *M. truncata* from 66-70 m asl in this silt dated 7340±170 BP [8170-7500 cal BP] (GSC-3397; Site 12, Fig. 4 and Table I) and provide a minimum age for the 75 m delta. These ages indicate ~5 m of emergence between 8.8 ka BP [9810-9340 cal BP] and 7.3 ka BP [8170-7500 cal BP], equivalent to an emergence rate of only 0.3 m/century, based on the calibrated ages (Fig. 5C, “minimum”). If the standard errors of the dates are considered, the 8.8 ka BP [9810-9340 cal BP] date related to the 92 m delta (the highest marine limit in this part of the fiord) and the 7.3 ka BP [8170-7500 cal BP] date related to a sea level at 70 m asl (the sample elevation), this results in a maximum initial emergence rate of 2 m/century (Fig. 5C, “maximum”).

Additional evidence suggests that emergence remained slow until at least 5.2 ka BP (5.6 ka cal BP). Immediately up-fiord from the 8.8 ka BP site, a marine limit delta is graded to a former relative sea level at 78 m asl. A sample of paired valves of *A. borealis* and *H. arctica* collected by A. Podor from bedded sand on the delta foreslope at 55 m asl dated 5200±70 BP [5840-5460 cal BP] (GSC-5897; Site 13, Fig. 4 and Table I). This provides a minimum age estimate on delta formation. Thus, assuming no elevation measurement error, sea level at 5.2 ka BP [5840-5460 cal BP] must have been at least as high as 55 m asl (Fig. 5C). This indicates that between 7.3 ka BP [8170-7500 cal BP] and 5.2 ka BP [5840-5460 cal BP] a *maximum* of 20 m of emergence occurred, equivalent to ~1 m/century based on the calibrated ages. A second sample consisting of a piece of drift-wood was recovered from 55 m asl in a raised beach which off-laps the same delta. This dated 6360±100 BP [7400-7010 cal BP] (GSC-5966; Site 14, Fig. 4 and Table I). Given the presence of GSC-5897 from the same elevation and

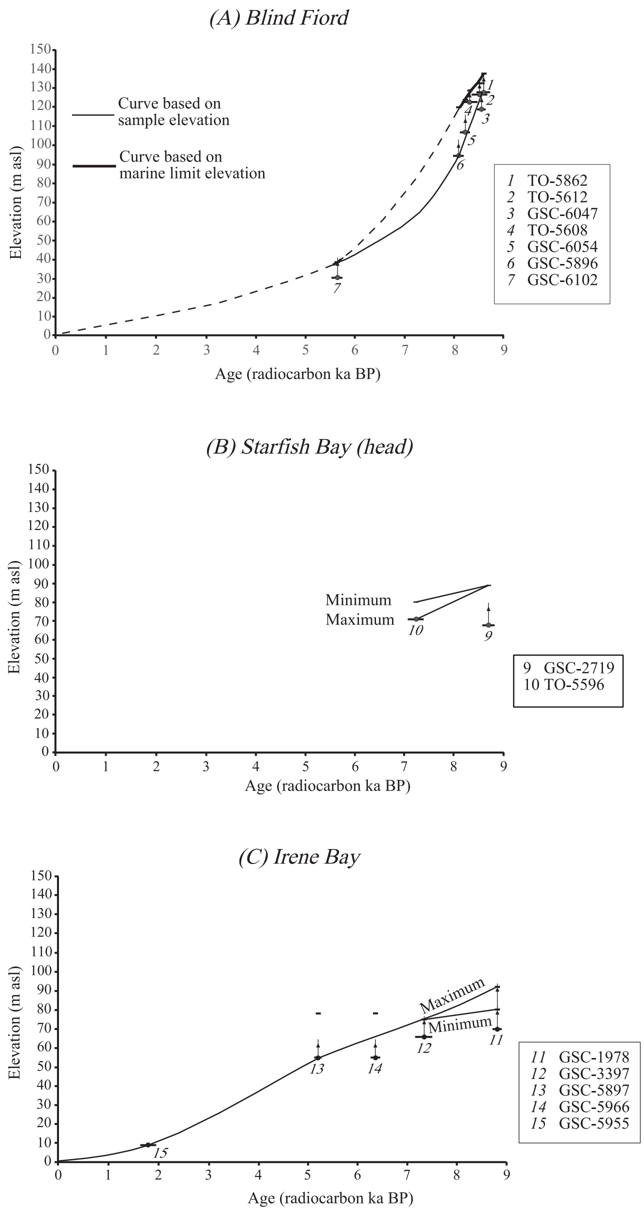


FIGURE 5. Emergence curves (radiocarbon years) from Ellesmere Island side of southern Eureka Sound. (A) Blind Fiord. (B) Starfish Bay (head). (C) Irene Bay. Dashed parts of curves are approximate. Site numbers of control points refer to Table I.

Courbes d'émersion (années au radiocarbone) de l'île d'Ellesmere (partie sud du Eureka Sound) (A) Blind Fiord. (B) Starfish Bay (fond). (C) Irene Bay. Les parties en pointillé sont approximatives. Les numéros de sites correspondent à ceux du tableau I.

location which dated >1000 years younger, this driftwood date is considered to be a maximum age on the 55 m sea level. A second driftwood sample recovered from 9 m asl in a raised beach at the fiord head dated 1790±160 [2050-1340 cal BP] (GSC-5955; Site 15, Fig. 4 and Table I).

POSTGLACIAL ISOBASES 8.5 KA BP

Regionally across western Ellesmere and Axel Heiberg islands, many radiocarbon dates of 8500±150 BP [9420-8650 cal BP] are available. In southern Eureka Sound, more dates fall into this interval than any other, and hence it was selected for assessing differential emergence in the study area. These dates form the control points for the isobase pattern in Figure 6 which integrates new shoreline and radiocarbon data from southern Eureka Sound with previously published information from northwestern Ellesmere Island (Bednarski, 1995), northern Eureka Sound (England, 1992; Bell, 1996), western and northeastern Axel Heiberg Island (Lemmen *et al.*, 1994; Bednarski, 1998) and Norwegian Bay (Hodgson, 1985; Hodgson in McNeely, 1989). This isobase reconstruction (Fig. 6) will likely be refined with additional radiocarbon and relative sea level data, particularly from Norwegian Bay.

Isobases drawn on the 8.5 ka shoreline rise in elevation towards Eureka Sound and form an elongate ridge oriented crudely north/south. The highest value (130 m asl) forms a closed cell of maximum emergence centred along the axis of the channel. At its northern end, the ridge extends north-eastwards into Greely Fiord with closure of the 120 m and 130 m isobases (*cf.* England, 1992). The northwestern extent of the 130 m isobase cell in Nansen Sound is uncertain, as Bednarski (1998) reports high marine limits (150-160 m asl) in this region, but associated dating control is poor. The 8.5 ka shoreline falls to ≤110 m asl in central Hare Fiord (Fig. 6) (Bednarski, 1995).

At the southern end, the 130 m isobase extends at least as far south as the mouth of Blind Fiord/Bear Corner, and may extend further south onto Bjorne Peninsula and the entrance to Norwegian Bay. Further southwest, on the north coast of Grinnell Peninsula, Devon Island, the 8.5 ka shoreline is <130 m asl (Dyke 1998, 1999). This demonstrates that the 130 m isobase closes to the northwest, and supports the above interpretation of closure in the vicinity of northern Bjorne Peninsula. It also indicates that the 120 m isobase either closes in Norwegian Bay in the vicinity of Graham and eastern Cornwall islands, or continues southwestwards onto Grinnell Peninsula. Currently, this cannot be resolved given the lack of emergence data from much of Norwegian Bay, and hence the 120 and 110 m isobases are left open to the south.

DISCUSSION

INITIAL POSTGLACIAL EMERGENCE

Marked contrasts in initial postglacial emergence are recorded in the study area. In Blind Fiord, initial emergence was rapid and characterised by rates of ≥5 m/century (cf. Blake, 1975, 1992a; Lemmen *et al*., 1994). At this site, relative sea level exhibits a continuous fall since deglaciation and marine limit formation (*cf.* Type A curve of Quinlan and Beaumont, 1981, and Zone 1 of Clark *et al*., 1978). The emergence history is thus broadly similar to that from other areas of eastern Arctic Canada formerly covered by the Laurentide Ice Sheet. Postglacial rebound in such areas typically exhibits continuous emergence since deglaciation with initial emer-

TABLE I

Holocene radiocarbon dates, greater Eureka Sound [a]

Site	Location	Laboratory dating No.[b]	Material	Age (years BP)	Enclosing material	Sample elev.(m asl)	Related RSL (m asl)	Calibrated age (cal BP)[c]
1	Ellesmere Island Baumann Fiord 78°08'N, 86°40'W	TO-5862	*Mya truncata* fragment	8590±70	Surface	128	≥128 - ≤138	9460-9060
2	Ellesmere Island Blind Fiord 78°11'N, 86°04'W	TO-5612	*Mya truncata* fragment	8510±80	Surface	127	≥127 - ≤133	9400-8970
3	Ellesmere Island Blind Fiord 78°14'N, 85°57'W	GSC-6047	*Hiatella arctica, Mya truncata*	8550±80	Gravel	119	≥119 - ≤133	9420-9000
4	Ellesmere Island Blind Fiord 78°22'N, 85°50'W	TO-5608	*Hiatella arctica* fragment	8310±80	Surface	123	≥123 - ≤129	9210-8630
5	Ellesmere Island Blind Fiord 78°22'N, 85°48'W	GSC-6054	*Mya truncata, Hiatella arctica*	8220±100	Surface	107	≥107 - ≤124	9100-8470
6	Ellesmere Island Blind Fiord 78°23'N, 85°40'W	GSC-5896	*Mya truncata*, *Hiatella arctica,* fragments	8090±110	Surface	95	≥95 - ≤120	8950-8370
7	Ellesmere Island Blind Fiord 78°13'N, 86°04'W	GSC-6102	*Hiatella arctica*, *Astarte borealis*	5640±110	Silt	31	≥31 - ≤39	6360-5870
8	Ellesmere Island Starfish Bay, 78°13'N, 84°34'W	GSC-6037	*Hiatella arctica*, *Mya truncata*	7740±90	Silt	78	>78 - ≤101	8410-8000
9	Ellesmere Island Starfish Bay 78°11'N, 84°08'W	GSC-2719	*Portlandia arctica*	8710±120	Silt	68	>68 - ≤89	9730-9060
10	Ellesmere Island Starfish Bay 78°12'N, 84°00'W	TO-5596	*Hiatella arctica*	7240±80	Silt	71	>71	7910-7570
11	Ellesmere Island Irene Bay 79°01'N, 81°31'W	GSC-1978	*Portlandia arctica*	8820±90	Silt	70-74	≥80 (≤92)	9810-9340
12	Ellesmere Island Irene Bay 79°01'N, 81°28'W	GSC-3397	*Hiatella arctica*	7340±170	Surface	66-70	≥70 - ≤75	8170-7500
13	Ellesmere Island Irene Bay 78°03'N, 81°28'W	GSC-5897	*Astarte borealis, Hiatella arctica*	5200±70	Sand and silt	55	≥55 - ≤78	5840-5460
14	Ellesmere Island Irene Bay 78°03'N, 81°28'W	GSC-5966	Driftwood	6360±100	Sand	55	≥55 - ≤78	7400-7010
15	Ellesmere Island Irene Bay 78°03'N, 81°28'W	GSC-5955	Driftwood	1790±160	Sand	9	9	2050-1340
16	Ellesmere Island Eureka Sound 79°14'N, 85°30'W	TO-2245	*Hiatella arctica*	8430±70	Silt	101	>101	9350-8920
17	Ellesmere Island Cañon Fiord 79°35'N, 80°31'W	TO-2339	*Mya truncata*	8380±80	Sand	98	>98 - ≤114	9310-8740
18	Ellesmere Island Fosheim Peninsula 79°48'N, 86°18'W	GSC-5156	*Hiatella arctica*	8680±80	Silt	132	>132 - ≤150	9560-9150
19	Ellesmere Island Fosheim Peninsula 79°50'N, 85°13'W	TO-2241	*Mya truncata*	8480±80	Silt	93	>93	9380-8950
20	Ellesmere Island Fosheim Peninsula 79°58'N, 85°22'W	GSC-4708	*Mya truncata*	8520±80	Silt	100	>100 - ≤146	9400-8970

TABLE I (cont.)

Holocene radiocarbon dates, greater Eureka Sound [a]

Site	Location	Laboratory dating No.[b]	Material	Age (years BP)	Enclosing material	Sample elev.(m asl)	Related RSL (m asl)	Calibrated age (cal BP)[c]
21	Ellesmere Island Fosheim Peninsula 79°58'N, 84°26'W	GSC-5155	*Mya truncata*	8570±120	Silt	100	>100 - ≤145	9500-8950
22	Ellesmere Island Fosheim Peninsula 80°11'N, 86°33'W	TO-2233	*Mya truncata*	8440±80	Surface	94	≥94	9370-8900
23	Ellesmere Island Fosheim Peninsula 80°15'N, 85°02'W	TO-2229	*Mya truncata*	8450±80	Silt	94	>94 - ≤149	9370-8920
24	Ellesmere Island Greely Fiord 80°24'N, 81°30'W	GSC-2369	*Mya truncata*	8450±100	Surface	127	≥127 - ≤146	9400-8820
25	Ellesmere Island d'Iberville Fiord 80°43'N, 80°35'W	S-2640	*Hiatella arctica*	8415±130	Silt	88	>88 - ≤139	9390-8660
26	Ellesmere Island Greely Fiord 80°43'N, 80°35'W	S-2645	*Mya truncata*	8465±130	Sand	95	>95 - ≤124	9430-8730
27	Ellesmere Island Hare Fiord 81°05'N, 85°14'W	S-2641	*Hiatella arctica*	8590±130	Silt	88	>88 - ≤110	9540-8940
28	Ellesmere Island Nansen Sound 81°06'N, 90°10'W	S-2639	*Mya truncata*	8370±130	Silt	103-105	>105	9360-8600
29	Axel Heiberg Is. Strand Fiord 91°47'N, 79°10'W	GSC-5408	*Mya truncata, Hiatella arctica*	8390±100	Sand and gravel	37-63	>63 - ≤105	9340-8700
30	Axel Heiberg Is. Strand Fiord 90°17'N, 79°15'W	GSC-5411	*Mya truncata*	8430±80	Surface	84-93	≥93 - ≤124	9360-8850
31	Ellesmere Island Baumann Fiord 78°06'N, 85°52'W	GSC-244	*Hiatella arctica*, *Mya truncata*, fragments	8480±140 [d]	Surface	116	≥116 - ≤122	
32	Ellesmere Island Trold Fiord 78°06'N, 85° 27W	AA-23591	*Hiatella arctica* fragment	8645±60	Gravel	131	≥131 - ≤143	9500-9180
33	Ellesmere Island Norwegian Bay 77°07'N, 87°42'	GSC-840	*Hiatella arctica, Mya truncata,* fragments	8590±150 [d]	Surface	107	≥107	
34	Graham Island 77°16'N, 89°57'W	GSC-2253	*Mya truncata*	8420±160	Sand	102	>102	9430-8600
35	Axel Heiberg Island Nansen Sound 80°22'N, 88°33'W	TO-2280	*Mya truncata*	8450±70	Surface	88	≥88	9350-8940

[a] Sources (including this paper): Dyck *et al.*, 1965; Blake, 1970; Hodgson, 1985; Hodgson *in* McNeely, 1989; England, 1990, 1992; Lemmen *et al.*, 1994; Bednarski, 1995, 1998; Bell, 1996.

[b] Laboratory designations: GSC = Geological Survey of Canada; TO = IsoTrace Laboratory, University of Toronto; S = Saskatchewan Research Council; AA = University of Arizona. TO and AA samples were dated by accelerator mass spectrometry. These samples were corrected for isotopic fractionation to a base of $\delta^{13}C$ = -25 ‰; a reservoir correction of 410 years was then applied, which is equivalent to correction to a base of $\delta^{13}C$ = 0 ‰; GSC and S samples were dated conventionally and corrected for fractionation to a base of ^{13}C = 0 ‰. GSC terrestrial organic samples were dated conventionally and corrected for fractionation to a base of $\delta^{13}C$ =-25 ‰.

[c] Dates were calibrated using CALIB 3.0 (Stuiver and Reimer, 1993) and the calibrated date range reported here is that which yields 100 % probability at 2σ.

[d] 1960's GSC uncorrected dates (Hodgson, 1985). These dates have not been corrected for isotopic fractionation or a marine reservoir effect. Approximate corrections could be made for isotopic fractionation to a base of $\delta^{13}C$ = -25 ‰ by adding 400-410 years to this uncorrected age (R. McNeely, unpublished communication to GSC clientele, 1991). A similar amount could then be subtracted to account for the marine reservoir effect. However, such a correction has not been applied as the result would be approximately the same as the uncorrected raw date reported here. GSC dates obtained during the course of this study (1990's) typically show differences between raw and corrected (to a base of $\delta^{13}C$ = 0‰) ages which are well within the reported standard errors of the individual dates.

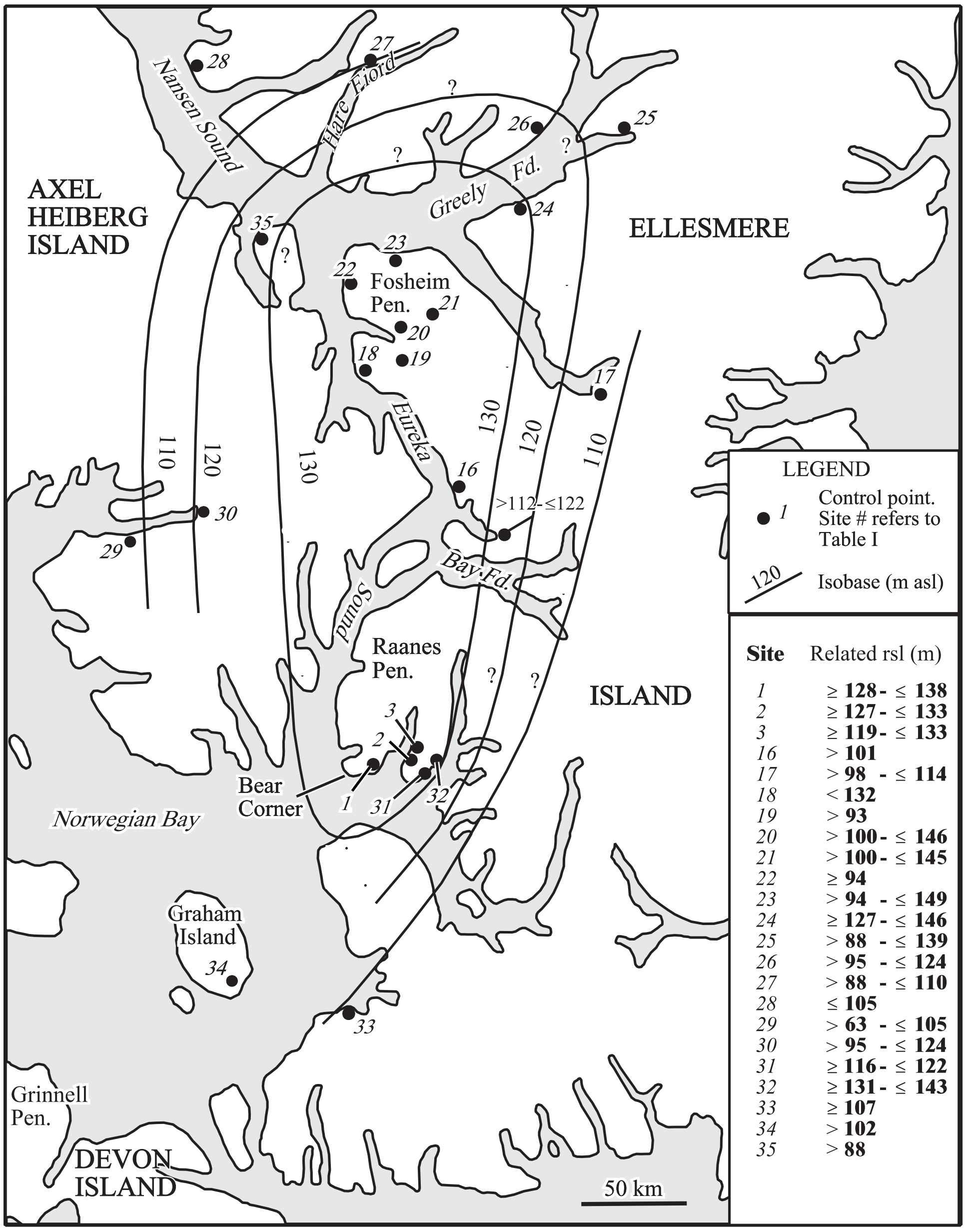


FIGURE 6. Postglacial isobases drawn on the 8.5 ka BP shoreline in greater Eureka Sound. The reconstruction utilizes radiocarbon dates from previous Quaternary studies in this region (Hodgson, 1985; England, 1992; Bednarski, 1995, 1998; Bell, 1996) and builds upon isobases drawn previously for northern Eureka Sound (England, 1992, 1997; Bell, 1996). Italicized site numbers refer to Table I. Control point without site numbers ("≥112 m-≤ 122 m") is based on a personal communication from J. England (1998).

Les isobases postglaciaires établies à partir du littoral de 8,5 ka BP (Eureka Sound). La reconstitution est fondée sur les datations au radiocarbone provenant d'études précédentes (Hodgson, 1985 ; England, 1992 ; Bednarski, 1995, 1998 ; Bell, 1996) et sur les isobases déjà établies pour la partie nord du Eureka Sound (England, 1992, 1997 ; Bell, 1996). Les numéros de sites correspondent à ceux du tableau I. Le point coté sans numéro de site (« ≥112 m-≤ 122 m ») nous a été communiqué par J. England (comm. pers., 1998).

gence of 7-30 m/century gradually decreasing to present (Andrews, 1970; Dyke, 1983, 1984; Dyke *et al.,* 1991). The relative sea level history of Blind Fiord is therefore compatible with glacial geological evidence that indicates extensive glaciation of southern Eureka Sound during the Last Glacial Maximum (Ó Cofaigh, 1998, 1999; Ó Cofaigh *et al.*, 1998, in press).

However, the Blind Fiord curve contrasts with those from Irene and Starfish bays, where initial emergence rates of ≤1 m/century are recorded. It is important to note that the timing of initial emergence at all three sites is similar. Slow initial emergence in Irene and Starfish bays commenced at 8.8 ka BP [9810-9340 cal BP] and 8.7 ka BP [9730-9060 cal BP] respectively, extending to at least 7.3 ka BP [8170-7500 cal BP] and 7.2 ka BP [7910-7570 cal BP]. However, in Blind Fiord, emergence during this *same* interval (~8.6-8.0 ka BP; 9460-9060 cal BP to 8950-8370 cal BP) was characterised by rates of ≥5 m/century.

Sea level curves similar to those from Irene and Starfish bays have been presented for several other areas on Ellesmere Island (England, 1983, 1992, 1997). For example, in Greely Fiord, England (1992) documented a period of relative sea level stability at marine limit from 8.8 to 7.8 ka (^{14}C) BP, after which emergence proceeded slowly (2 m/century) until 7.2 ka BP when it increased to 13 m/century. He attributed this arrest in sea level following marine limit formation to a balance between glacioisostatic uplift and eustatic sea level rise. The lack of initial rapid emergence at these sites was considered to be a consequence of a limited Late Wisconsinan glacial cover as inferred from independent glacial geologic evidence (England, 1978, 1990, 1996).

Emergence curves showing a period of relative sea level stability at marine limit have also been presented from the area of the former Barents Sea Ice Sheet, from Spitsbergen (Forman, 1990) and Franz Josef Land (Forman *et al*., 1996, 1997), and this stability is also attributed to a balance between glacioisostasy and eustasy. These sites are inferred to have sustained a thinner glacial load than at the former ice sheet centre over the northern and western Barents Sea where Type A (Quinlan and Beaumont, 1981) sea level curves are reported (Forman, 1990; Bondevik *et al*., 1995). It should be noted, however, that ice thicknesses over the sites which exhibit slow initial unloading may have been as much as 1500 m (Forman *et al.,* 1995, 1996).

Thus the slow initial emergence recorded in Irene and Starfish bays may not necessarily be incompatible with an extensive ice cover over these sites during the Late Wisconsinan. A minimum estimate of ice thickness at the Last Glacial Maximum in southern Eureka Sound based on glacial geological evidence is 1200 m (Ó Cofaigh, 1999), and maximum estimates obtained from glaciological modelling are 1500 m or 2000 m (Reeh, 1984). The three emergence curves would therefore imply marked spatial variations in the form and rate of initial emergence, even between sites in close proximity; Blind Fiord and inner Starfish Bay are <45 km apart. Slow initial emergence at the fiord heads of Irene and Starfish bays might therefore reflect a major ice-marginal stillstand during retreat, as marked by the “drift belt” (Hodgson, 1985; Ó Cofaigh, 1998; Ó Cofaigh *et al*., in press), which locally restrained rebound at both fiord heads.

However, radiocarbon dates indicate that while rapid initial emergence (≥5 m/century) was proceeding in Blind Fiord, Starfish and Irene bays were experiencing slow initial emergence (≤1 m/century). Because of the proximity in the timing of deglaciation and marine limit formation between the three sites, the validity of such marked spatial variations in initial emergence is questionable, as the associated unloading would presumably have been integrated over a wide area (Walcott, 1970; Andrews, 1970). Thus, the argument that the slow emergence in Starfish and Irene bays simply represents the later part of the local emergence history, and that an earlier phase of rapid emergence is not recorded at these sites on account of their later deglaciation (*vs*. sites closer to Eureka Sound such as Blind Fiord), is negated by the presence of early Holocene radiocarbon dates at both fiord heads.

Obviously, the legitimacy of slow initial emergence in Starfish and Irene bays is dependent upon the validity of two fiord head dates of 8.8 ka BP [9810-9340 cal BP] and 8.7 ka BP [9730-9060 cal BP] on *P. arctica* which are associated with relative sea levels of 80-92 m asl. This reconstruction would be invalid if the samples relate to higher relative sea levels not observed in the field, or if the true ages of the original samples were actually younger than the reported ages. No evidence was found at either fiord head for higher shorelines above the surveyed marine limit which is well defined by raised deltas and wave-cut benches.

With respect to the dates being erroneously old, Forman and Polyak (1997) have demonstrated from radiocarbon dating of pre-bomb *P. arctica* that this species can have marine reservoir values as high as 764 years, possibly reflecting either the incorporation of old carbon from surrounding deposits and porewater, or from freshwater inputs by streams or glacial meltwater. A variable reservoir effect for this genus, possibly as high as 700-800 years, could therefore imply that some dates on *P. arctica* may be too old. Currently the reservoir effect for *P. arctica* in Eureka Sound is unknown. However, early Holocene radiocarbon dates on *P. arctica* from eastern Ellesmere and western Axel Heiberg islands (Blake, 1992a; Lemmen *et al*., 1994) compare closely to dates obtained on other genera from the same sites and units, indicating that not all *P. arctica* dates are anomalously old. Resolution of this issue can be made through radiocarbon dating of pre-bomb *P. arctica* from the Canadian High Arctic, and this work is currently in progress (J. England, personal communication 1998).

POSTGLACIAL ISOBASES

Isobases drawn on the 8.5 ka shoreline demonstrate an elongate ridge of emergence, oriented crudely parallel with the axis of Eureka Sound, extending from Greely Fiord in the north to the entrance to Norwegian Bay in the south. This extends previous reconstructions of postglacial shoreline delevelling in the region (Blake, 1970; England, 1976b, 1992, 1997; Bell, 1996), and is significant in that it demonstrates: (a) that the highest emergence values form a cell over the length of Eureka Sound; and (b) that this highest cell does not appear to extend southwestwards across Norwegian Bay to Grinnell Peninsula on Devon Island (*cf*. Blake, 1970), but rather closes in the vicinity of the entrance to Norwegian Bay.

Along the Eureka Sound/Nansen Sound fiord system, glacial geologic and chronologic evidence indicates extensive Late Wisconsinan glaciation, which inundated fiords and inter-island channels (Ó Cofaigh, 1998, 1999; Ó Cofaigh *et al*., 1998, in press; Bednarski, 1998). A similar reconstruction, advocating extensive ice during the Last Glacial Maximum, has recently been presented for Norwegian Bay, Wellington Channel and Devon Island (Hättestrand and Stroeven, 1996; Dyke, 1998, 1999). This glacial geological evidence negates an explanation of this isobase pattern in terms of overlapping peripheral depressions from separate ice masses on Ellesmere and Axel Heiberg islands (cf. England, 1976a). The pattern of shoreline delevelling recorded by the isobases is inferred to be the result of a glacioisostatic response to the unloading which accompanied early Holocene deglaciation (*cf*. Blake, 1970). This supports earlier reconstructions of maximum Late Wisconsinan loading along Eureka Sound/Nansen Sound (*e.g.*, Blake, 1970; Walcott, 1972).

Modelling of an Innuitian Ice Sheet over the Canadian High Arctic during the Last Glacial Maximum (Reeh, 1984) results in *maximum* ice thicknesses of 1500 and 2000 m being located *west* of the main ice-divide, in the vicinity of Fosheim Peninsula/Eureka Sound. In contrast, ice at the modelled main ice-divide, which is located along the highland rim of eastern Ellesmere Island, is ~700-1000 m thick (Reeh, 1984) (current thickness ~500-800 m, Koerner, 1989). Koerner *et al.* (1987) also propose that the ice at the main divide over the Agassiz Ice Cap during the Wisconsinan was only 200 m thicker than today. Glacial geological data from Eureka Sound indicate a *minimum* ice thickness of ~1200 m in the channel during the Last Glacial Maximum (Ó Cofaigh, 1999). This suggests that the 8.5 ka BP isobase pattern which shows maximum emergence along Eureka Sound reflects the thickest ice load being located there during the Last Glacial Maximum (cf. Blake, 1970). Closure of the highest values to the southwest points to a possible saddle connecting the Eureka Sound loading centre to that proposed in the western part of Norwegian Bay (Hättestrand and Stroeven, 1996; Dyke 1998, 1999).

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